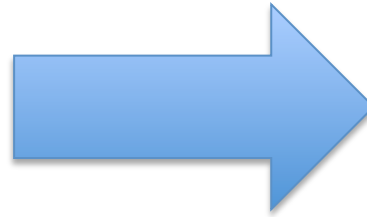
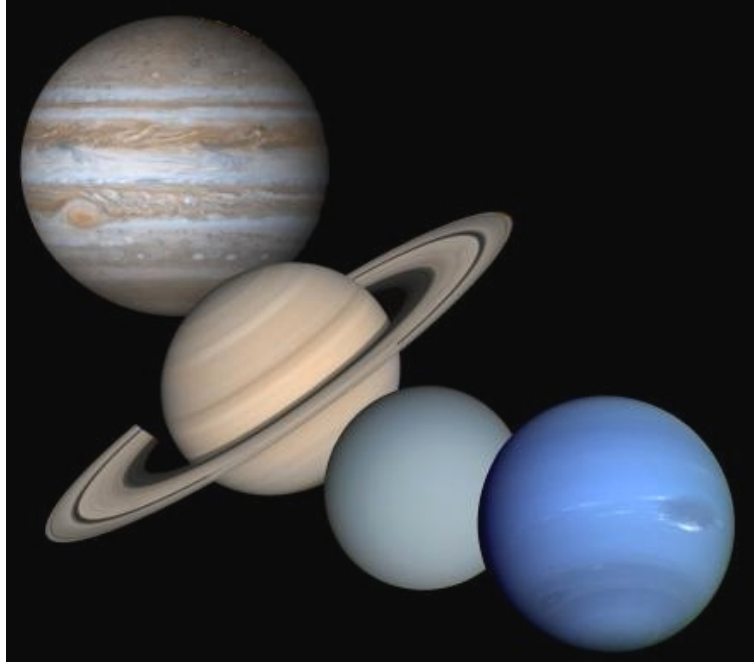


[NASA]

**Planetary outer atmospheres:  
interfaces from atmospheres  
to magnetospheres**

*Marina Galand  
Imperial College London, UK*

# Why should we care about giant planets?



**Our Solar System is the only piece of the Universe that we can examine *in situ*.**

**To understand the many extrasolar giant planets being discovered, we need to study giant planets in our own neighbourhood.**

# Outline

## **1. Setting the scene**

- Thermosphere/Ionosphere
- Energy sources

## **2. Ionosphere at the giant planets**

- What? How? Where? How much?

## **3. Auroral emissions**

## **4. Magnetosphere-ionosphere-thermosphere coupling**

- Energy crisis at the giant planets

## **5. Future prospect**

Setting the scene

# Atmospheres in the solar system

## Atmosphere composition

### N<sub>2</sub> atmospheres

- Earth
- Titan
- Triton
- Pluto

### CO<sub>2</sub> atmospheres

- Venus
- Mars
- Callisto (exosphere)

### He/H/H<sub>2</sub> atmospheres

- Jupiter: P10/P11/V1/V2/Ulysses/Cassini, Galileo
- Saturn: P11/V1/V2, Cassini-Huygens
- Uranus: V2
- Neptune: V2

### H<sub>2</sub>O atmosphere (exosphere)

- Comets
- Enceladus

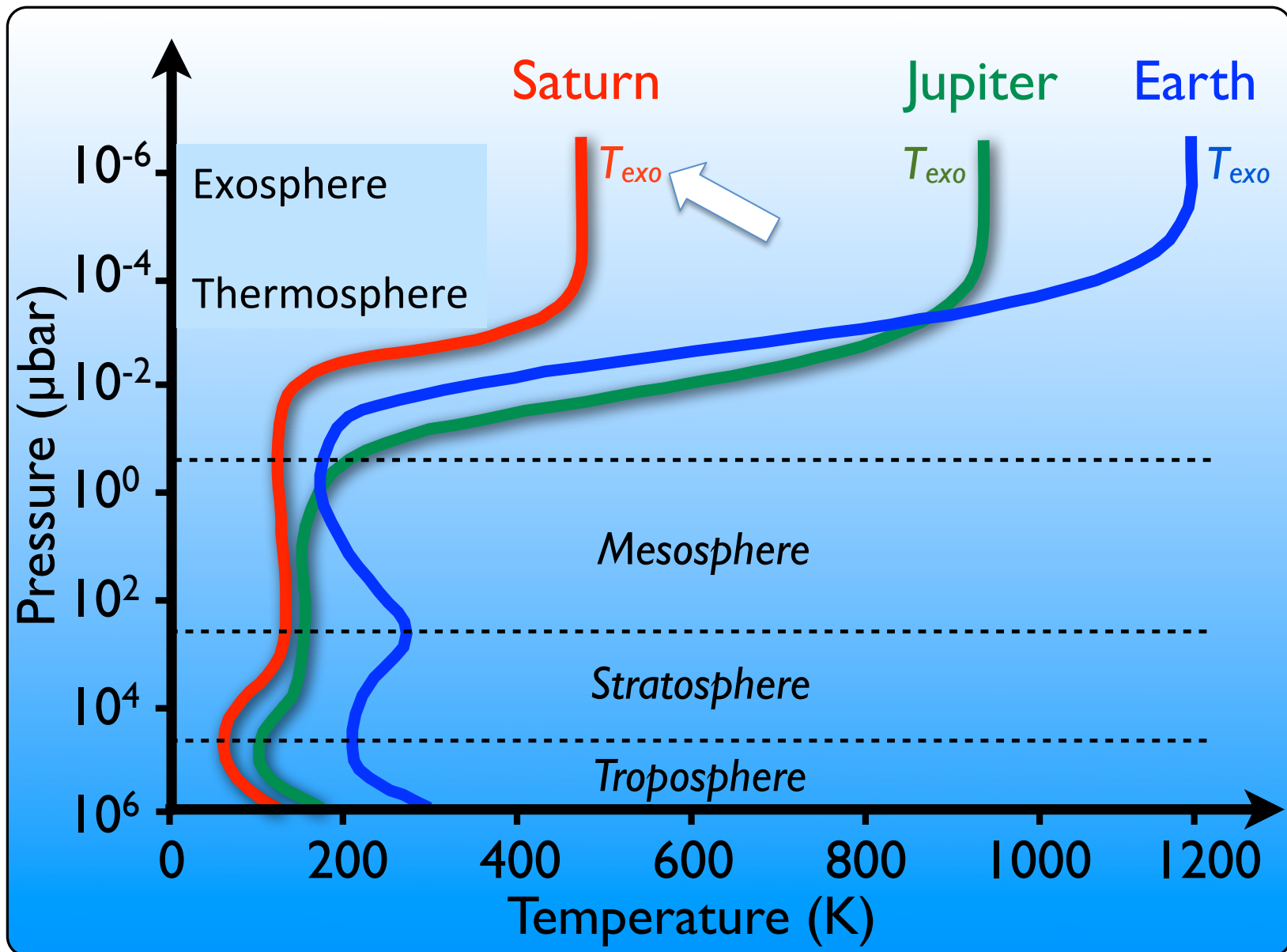
### SO<sub>2</sub> atmosphere (exosphere)

- Io

### O/O<sub>2</sub>/H<sub>2</sub>O atmosphere (exosphere)

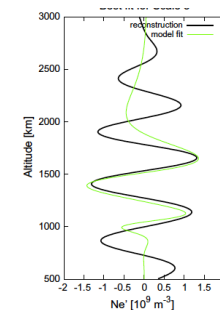
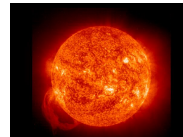
- Ganymede
- Europa

# Thermal Profiles



# Outer (Upper) Atmosphere

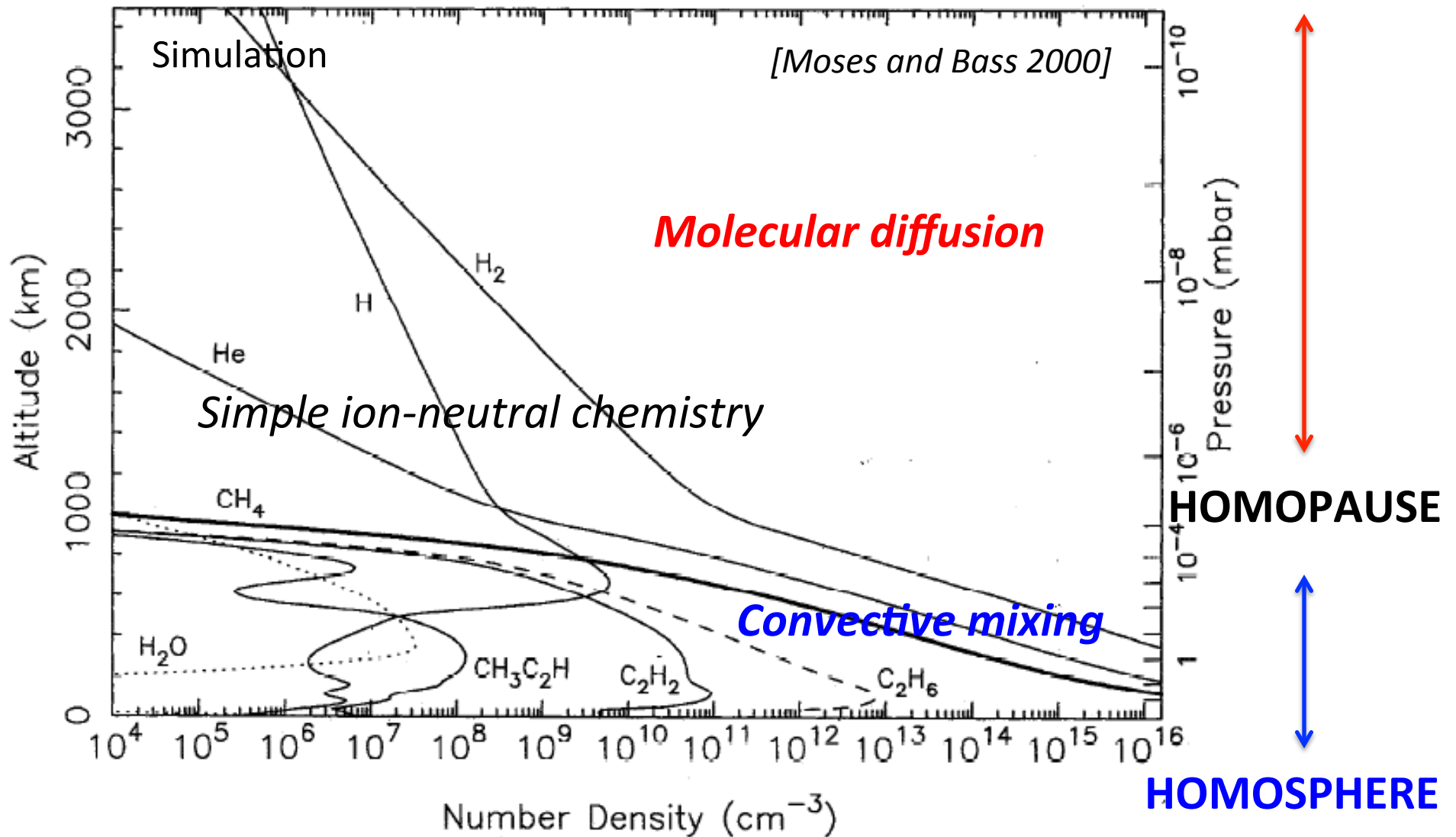
- **Gravitationally bound**
- **Key transition region** between the lower atmosphere and the magnetosphere/space environment
- **Energy and momentum sources:**
  - EUV/FUV solar radiation
  - Energetic particles from the planetary magnetosphere
  - Forcing from below (e.g., gravity waves)
- **Exist at all bodies** with “dense-enough” atmosphere



# Thermosphere of Saturn



Reference altitude (0 km) = 1 bar level

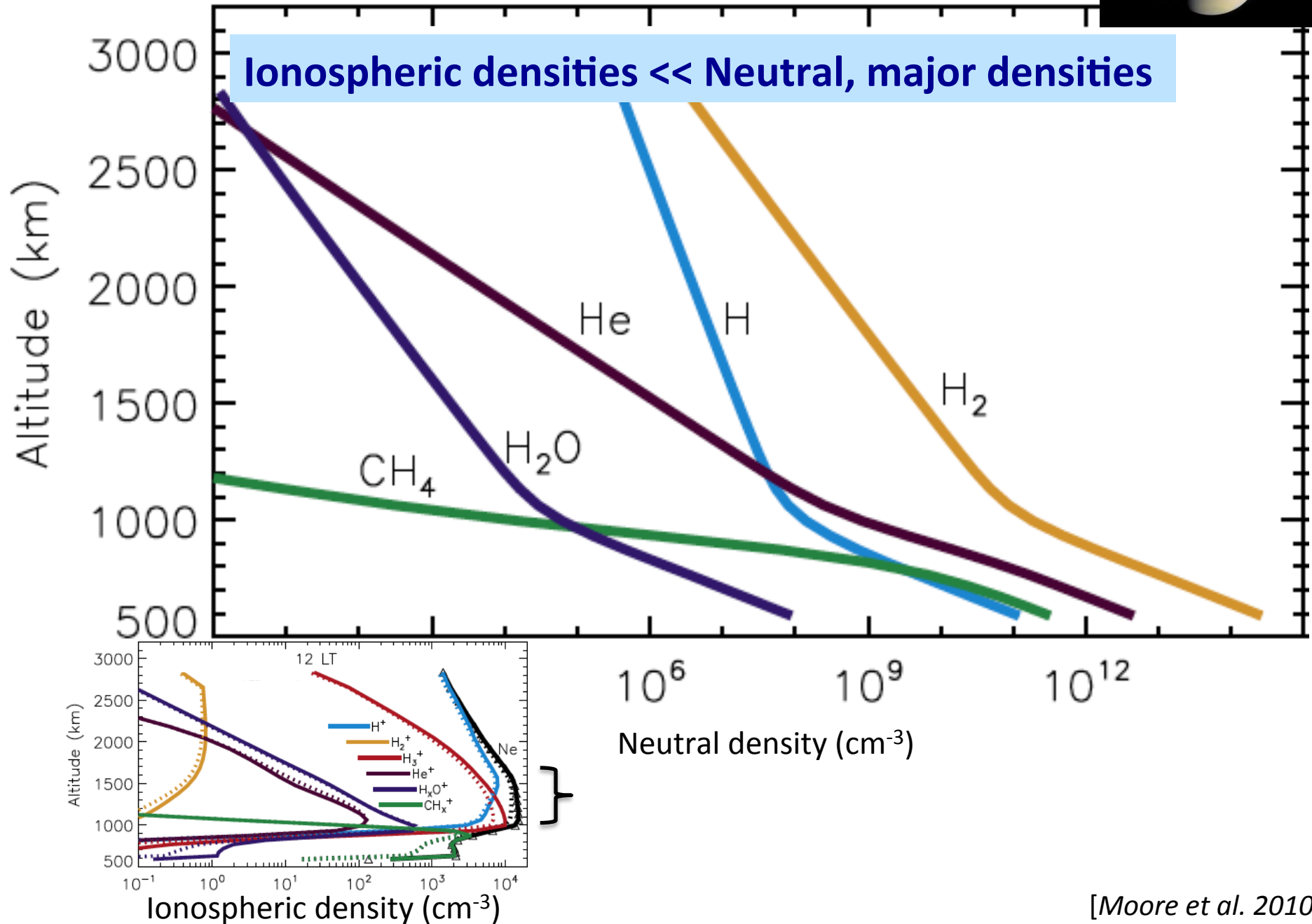




# Ionosphere

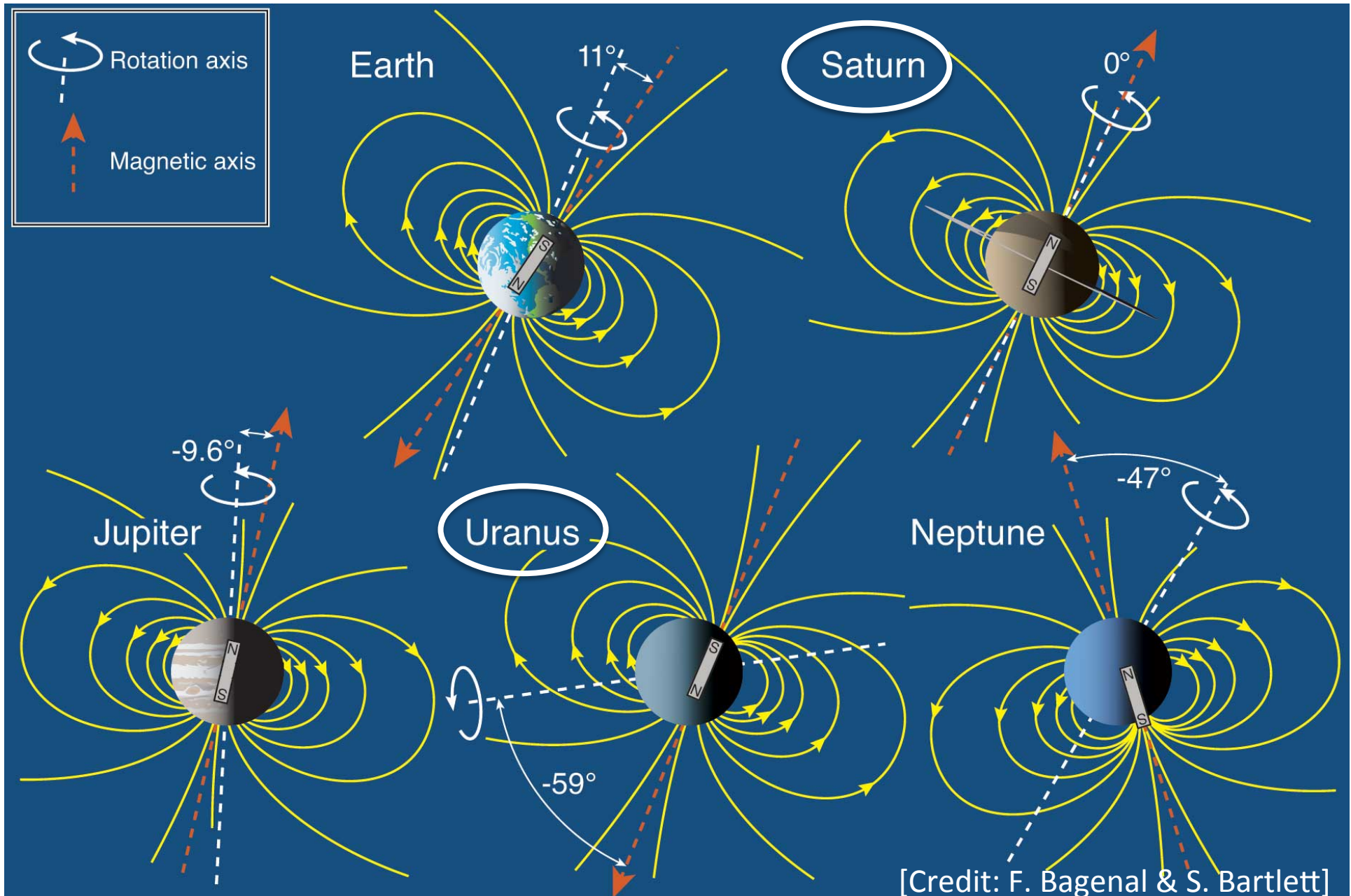
- ***Ionized part*** of the upper atmosphere
- ***Present at any planet*** (or moon) which has an atmosphere
- ***Conducting layer***:
  - Key source of *heating* of the high latitude upper atmosphere
- ***Key layer for coupling*** between the upper atmosphere and the magnetosphere:
  - *Closure of magnetospheric current system*

# Saturn's ionosphere

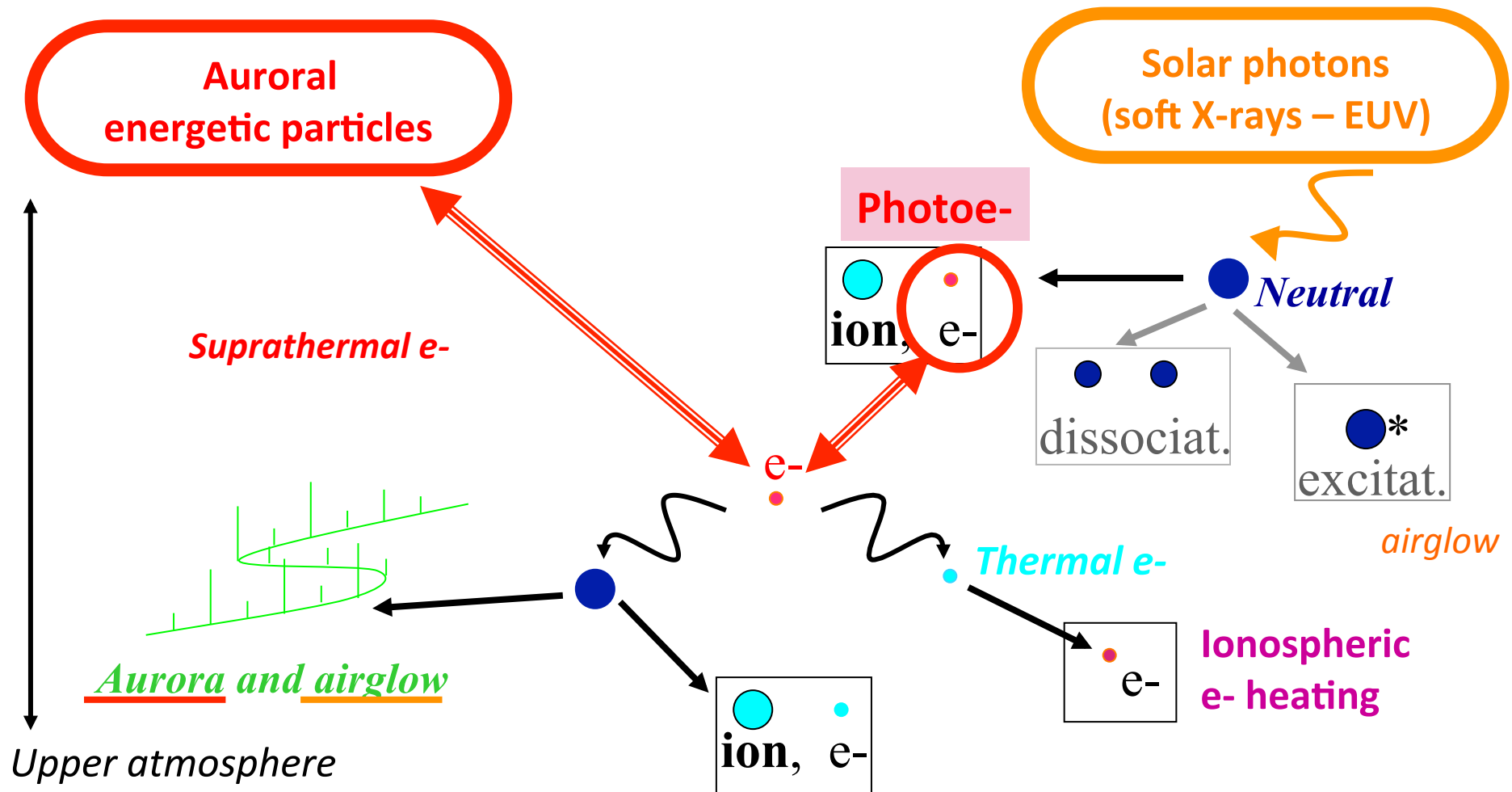


[Moore et al. 2010]

# Tilts of planetary magnetic fields with respect to their rotation axis



# Introduction: Solar and magnetospheric particle deposition



- *Creation of an ionosphere*
- *Major contribution on chemistry, dynamics, & energetics*

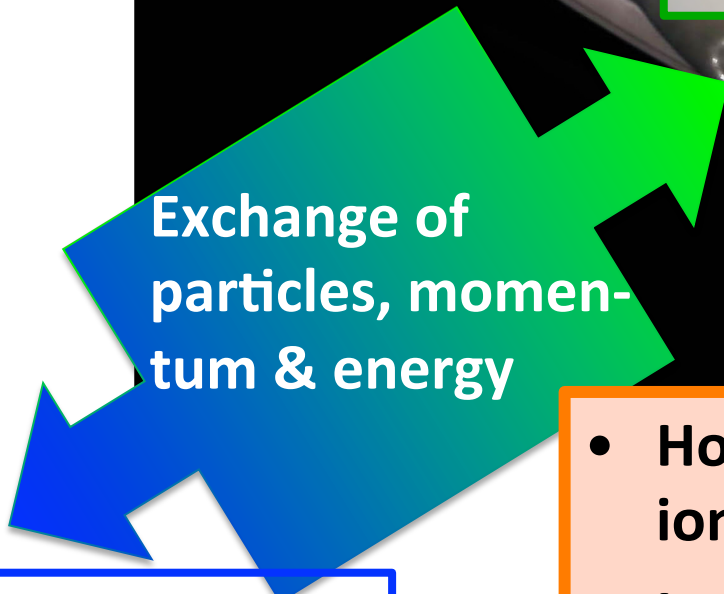
# MAGNETOSPHERE-IONOSPHERE-THERMOSPHERE COUPLING

AURORAL  
THERMOSPHERE  
IONOSPHERE

Exchange of  
particles, momen-  
tum & energy

MAGNETOSPHERE

- How well is the vertical and latitudinal ionospheric structure captured?
- Is there any atmosphere-magnetosphere link at low- & mid-latitudes?



# Ionosphere at the Giant Planets

## Ionization sources

- **Ionisation potential:**

- H<sub>2</sub>: 15.43 eV  $\leftrightarrow$  80 nm

- H: 13.60 eV  $\leftrightarrow$  91 nm

- CH<sub>4</sub>: 12.55 eV  $\leftrightarrow$  99 nm

13 eV  $\leftrightarrow$  ~100 nm

- **Solar EUV radiation (1-10, 10-100 nm):**

- Solar flux / (Sun-planet distance)<sup>2</sup>

- **Energetic particles** from the space environment

- A few keV to a few 100s keV

# Energy sources

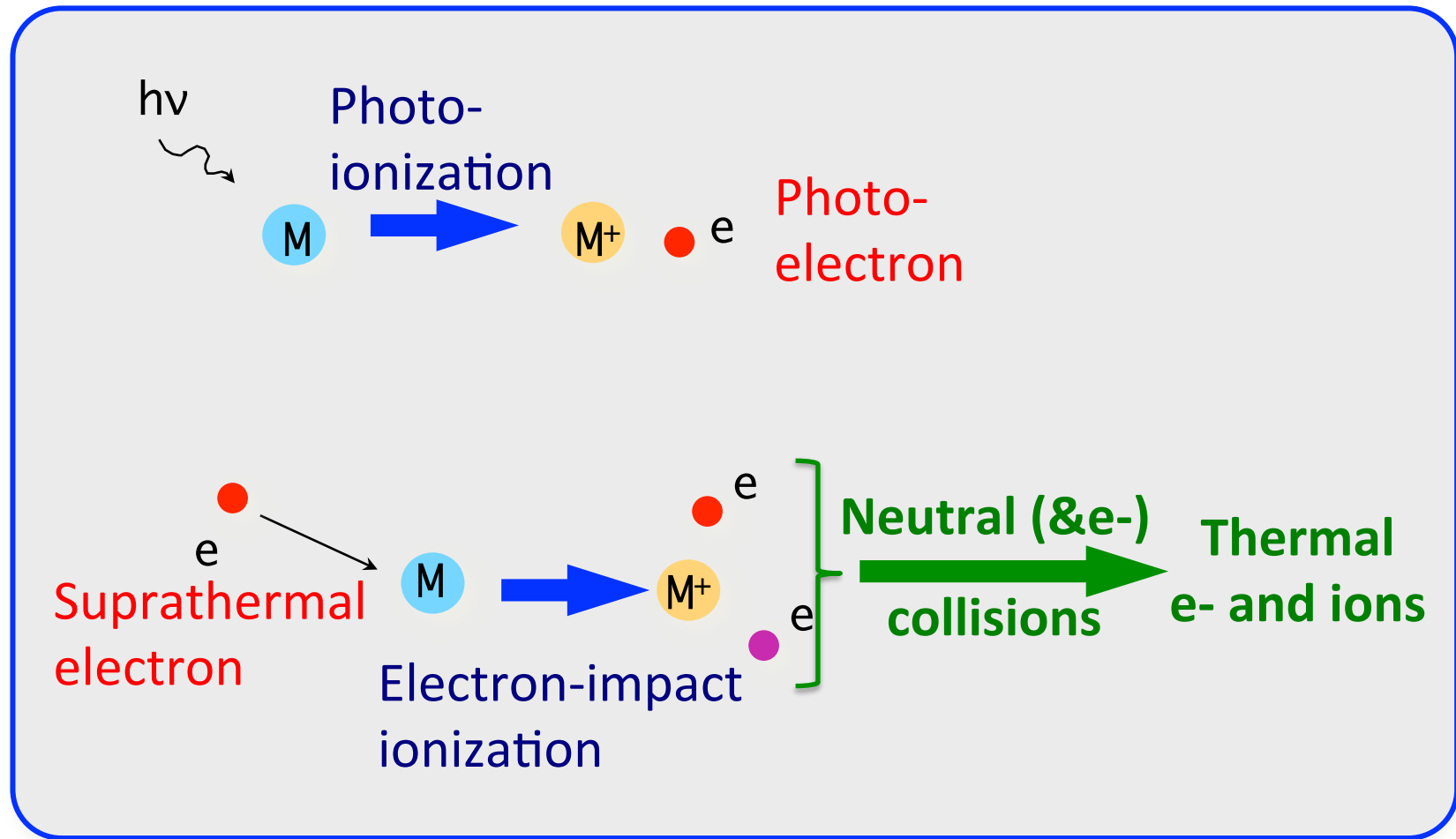
	Solar EUV input*	Auroral input*	Auroral particle input**
Earth (1 AU)	500 GW	80 GW	1-10 keV
Jupiter (5.2 AU)	800 GW	$10^5$ GW	30-200 keV 2-30 mW m <sup>-2</sup>
Saturn (9.5 AU)	200 GW	$(5-10) \times 10^3$ GW	10-20 keV ~ 1 mW m <sup>-2</sup>
Uranus (19 AU)	8 GW	100 GW	-
Neptune (30 AU)	3 GW	1 GW	-

\* Auroral input refers to “particle + Joule heating” (Strobel 2002)

\*\* Values valid for the main auroral oval, inferred from the analysis of auroral emissions (e.g., Fox et al. 2008, Gustin et al. 2004, 2009)



# Two electron populations



# Plasma population in the upper atmosphere

- **Ionospheric, thermal population** ( $e^-$ , ions):
  - Bulk of the plasma population
  - Thermalized through collisions
  - **Fluid treatment: macroscopic quantities**
    - Number density  $n$  (continuity equation), bulk velocity  $\underline{u}$  (momentum equation), temperature  $T$  (energy equation)
- **Suprathermal particles** ( $e^-$ , ion, neutrals):
  - Of solar or magnetospheric origin (also: meteor, ...)
  - Small in terms of density (high energy tail)
  - Source of energy to the upper atmosphere (e.g., ionisation, excitation/dissociation, thermal electron heating)
  - **Kinetic treatment: microscopic quantity**
    - Particle intensity  $I_e$  (Boltzmann equation) →  $P_e, Q_e, \dots$

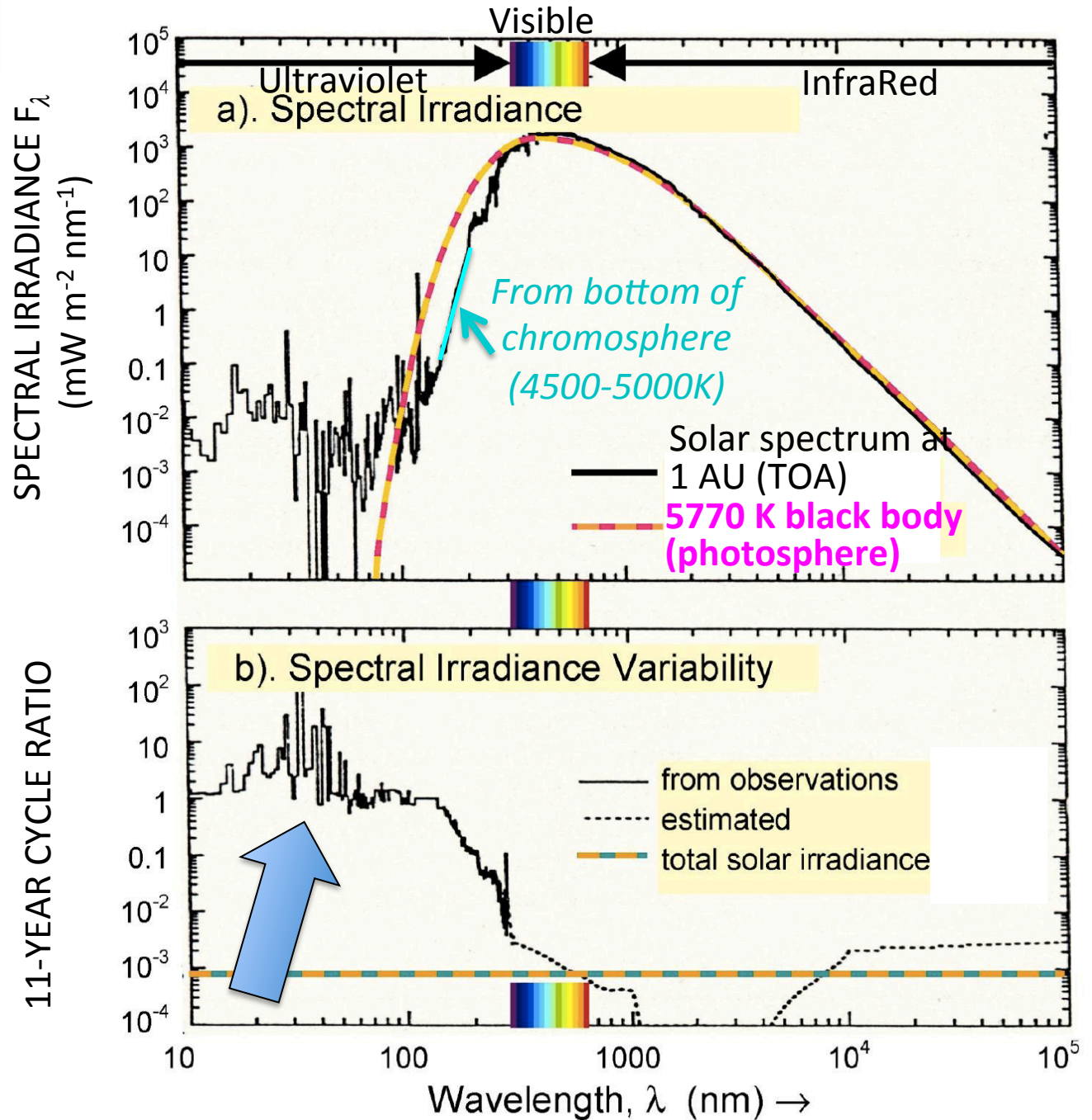
# Solar spectrum and its variability

SOLAR SPECTRUM:  
 9% in UV,  
 38% in visible,  
 53% in near IR (0.7-4  $\mu\text{m}$ )

At 1 AU

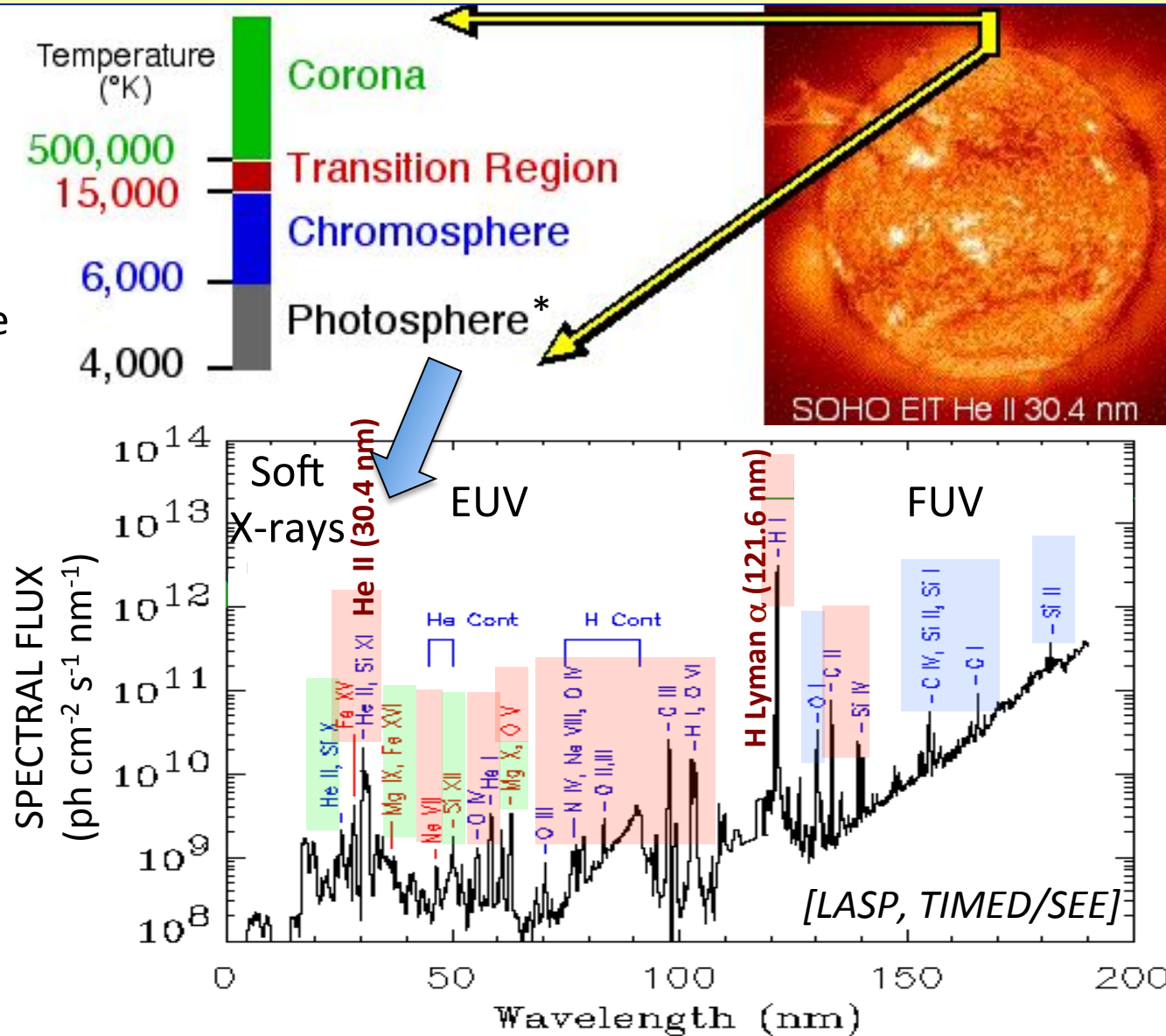
$$\frac{F_{\lambda}^{\max} - F_{\lambda}^{\min}}{F_{\lambda}^{\min}}$$

Lean (1991)  
 [adapted by Lockwood]



# Solar spectrum – origin of the soft X-ray, EUV and FUV

\* visible surface of the Sun

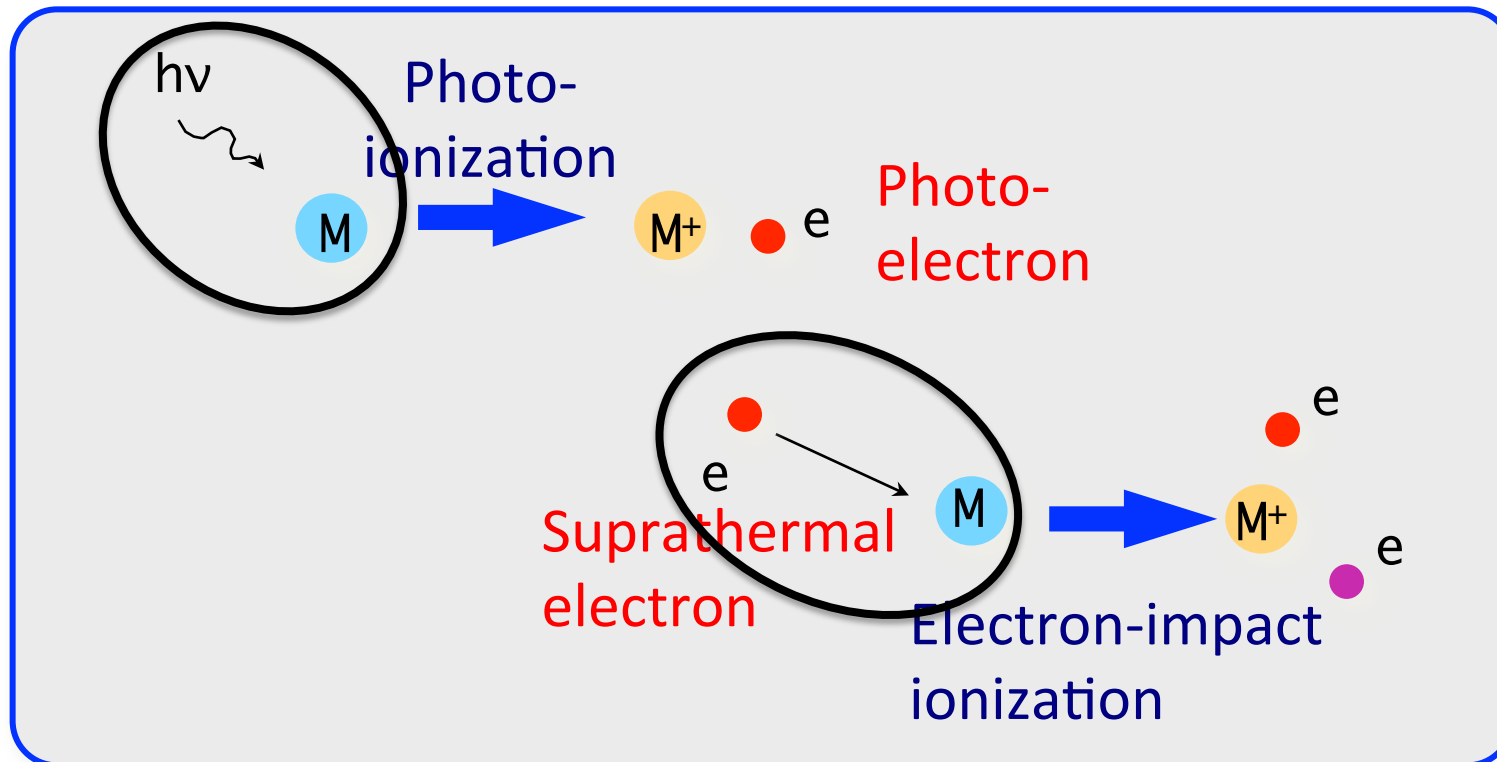


The solar soft X-ray and EUV radiation from the hot transition region and corona are responsible for the creation of an ionosphere.

# Ionization processes

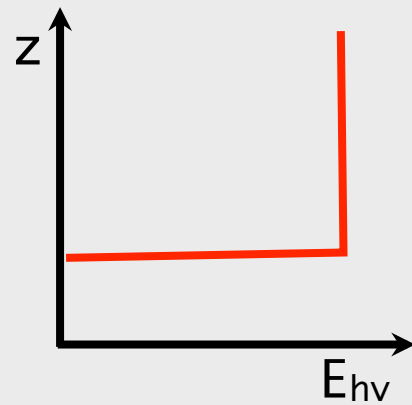
Range of ionization processes:

- Photoionization:  $h\nu + M \rightarrow M^+ + e^-$
- e-impact ionization:  $e^- + M \rightarrow e^- + M^+ + e^-$



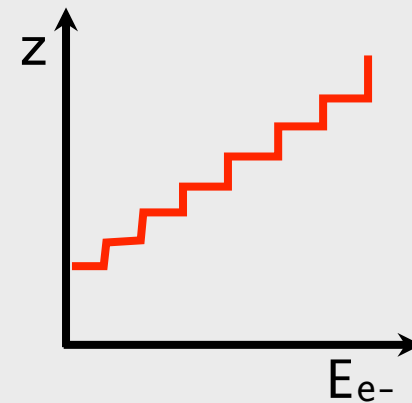
# Energy deposition

Solar photons  
(no scattering)



Absorption

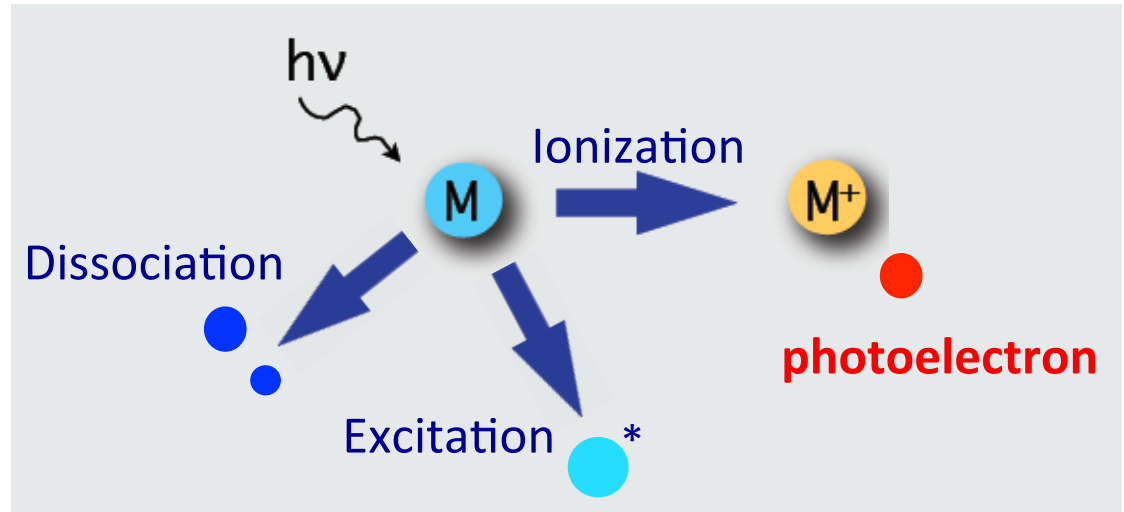
Suprathermal  $e^{-*}$



Collisions:  
discrete energy loss

\* *Photoelectrons, Auroral electrons, secondary electrons*

# Absorption of solar radiation in an atmosphere



- ✓ In the EUV, primarily extinction in the beam  
→ apply Beer-Lambert Law

- ✓ **Attenuated solar flux** at wavelength  $\lambda$  and at altitude  $z$ :

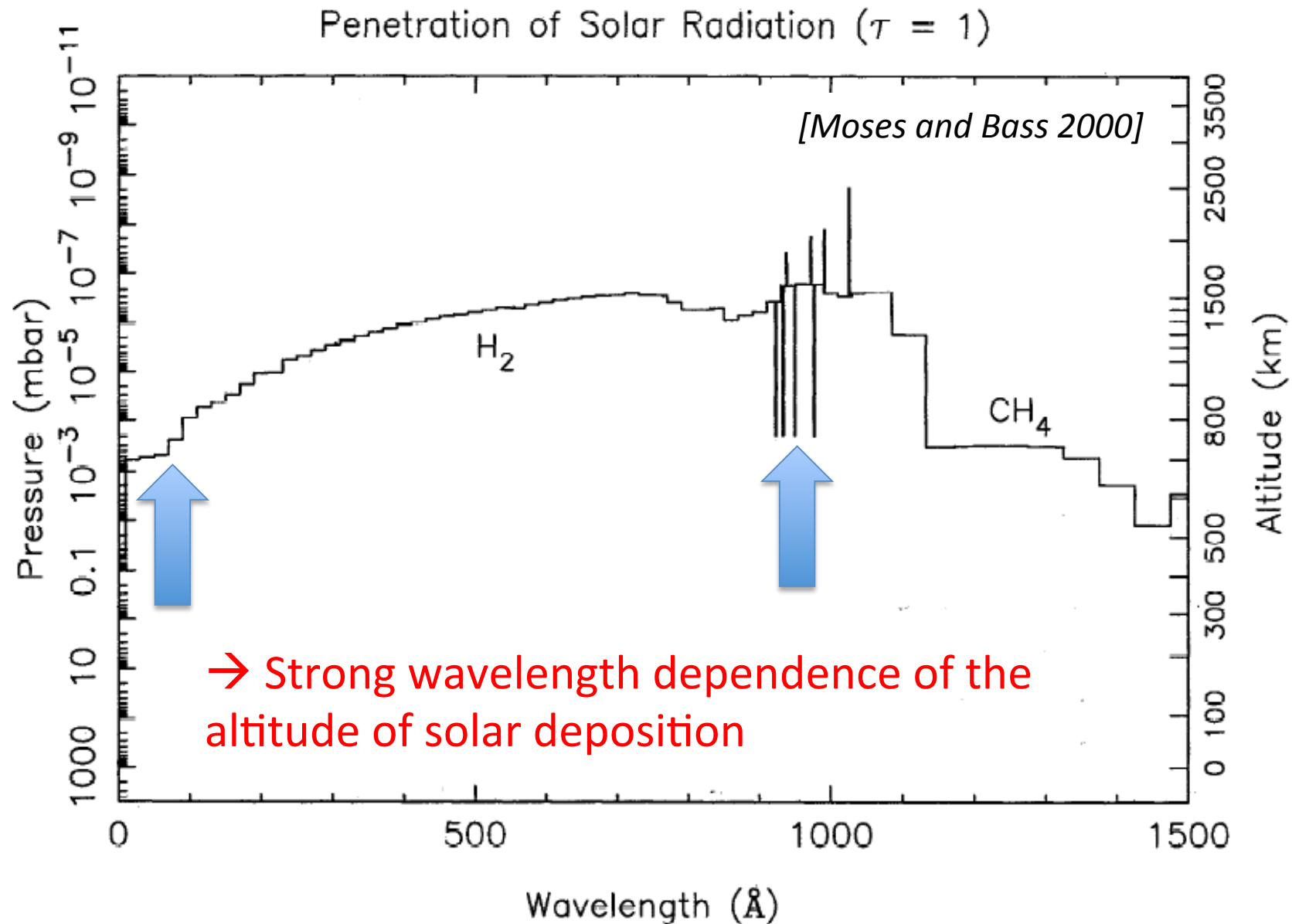
$$I_{\lambda}(z) = I_{\lambda}^{TOA} \exp\left(-\sum_n \sigma_n^{abs}(\lambda) \int_z^{\infty} n_n(z') \sec(\chi) dz'\right)$$

Solar flux at top of atmos  $\rightarrow$   $I_{\lambda}^{TOA}$   
 Photo-absorption cross section  $\rightarrow$   $\sigma_n^{abs}(\lambda)$   
 Number density of neutral  $n$   $\rightarrow$   $n_n(z')$   
 Solar zenith angle  $\rightarrow$   $\chi$   
**Optical depth  $\tau$**   $\rightarrow$   $\int_z^{\infty} n_n(z') \sec(\chi) dz'$

- ✓ **Photoelectron production rate** at  $\lambda$ :

$$P_{e,\lambda}(z) = \sum_n \sigma_n^{ioni}(\lambda) n_n(z) I_{\lambda}(z) \propto I_{\lambda}^{TOA} \propto \frac{1}{(d_{Sun-Planet})^2}$$

# Altitude of deposition

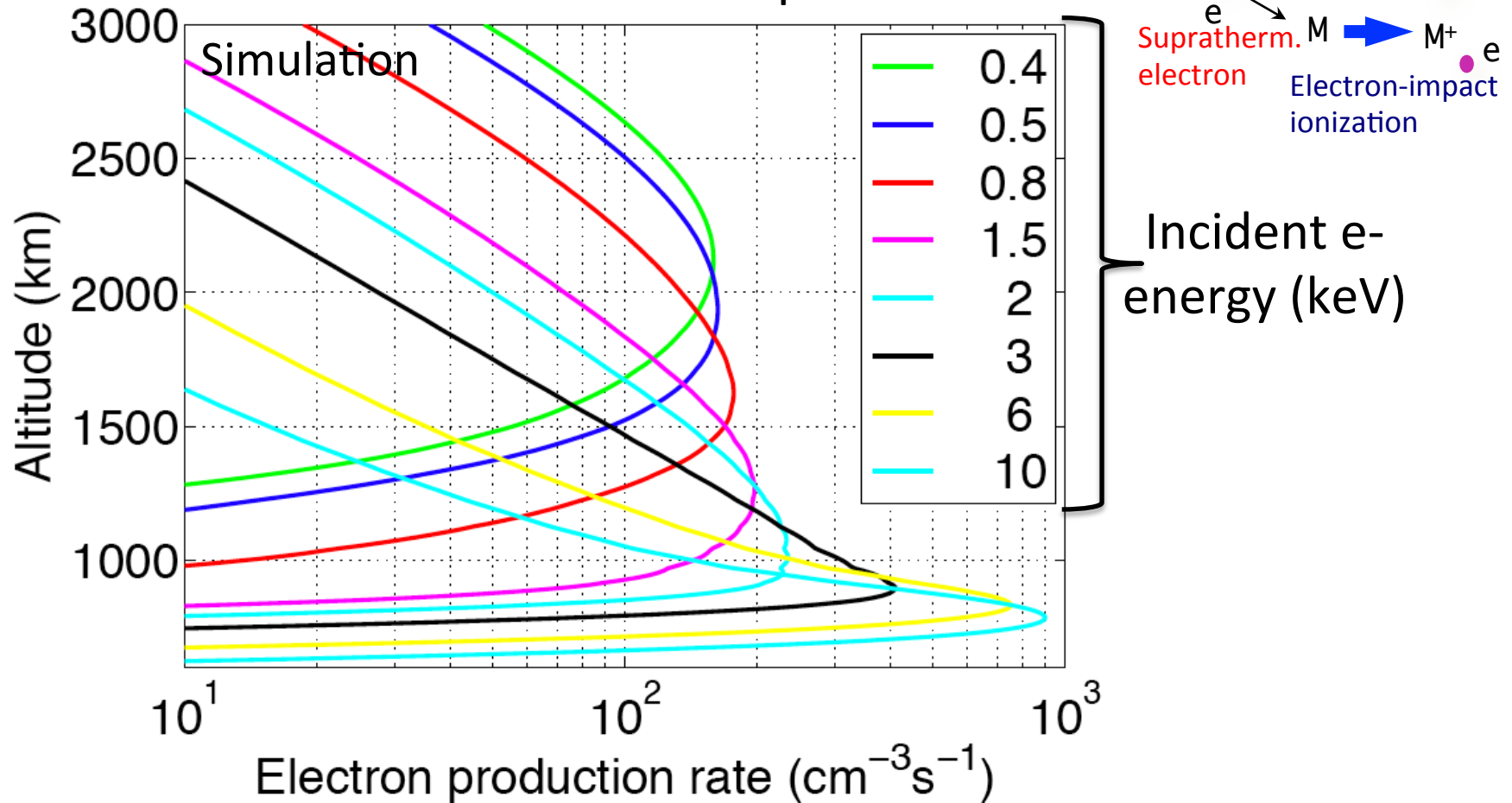




# Energy deposition of auroral electrons

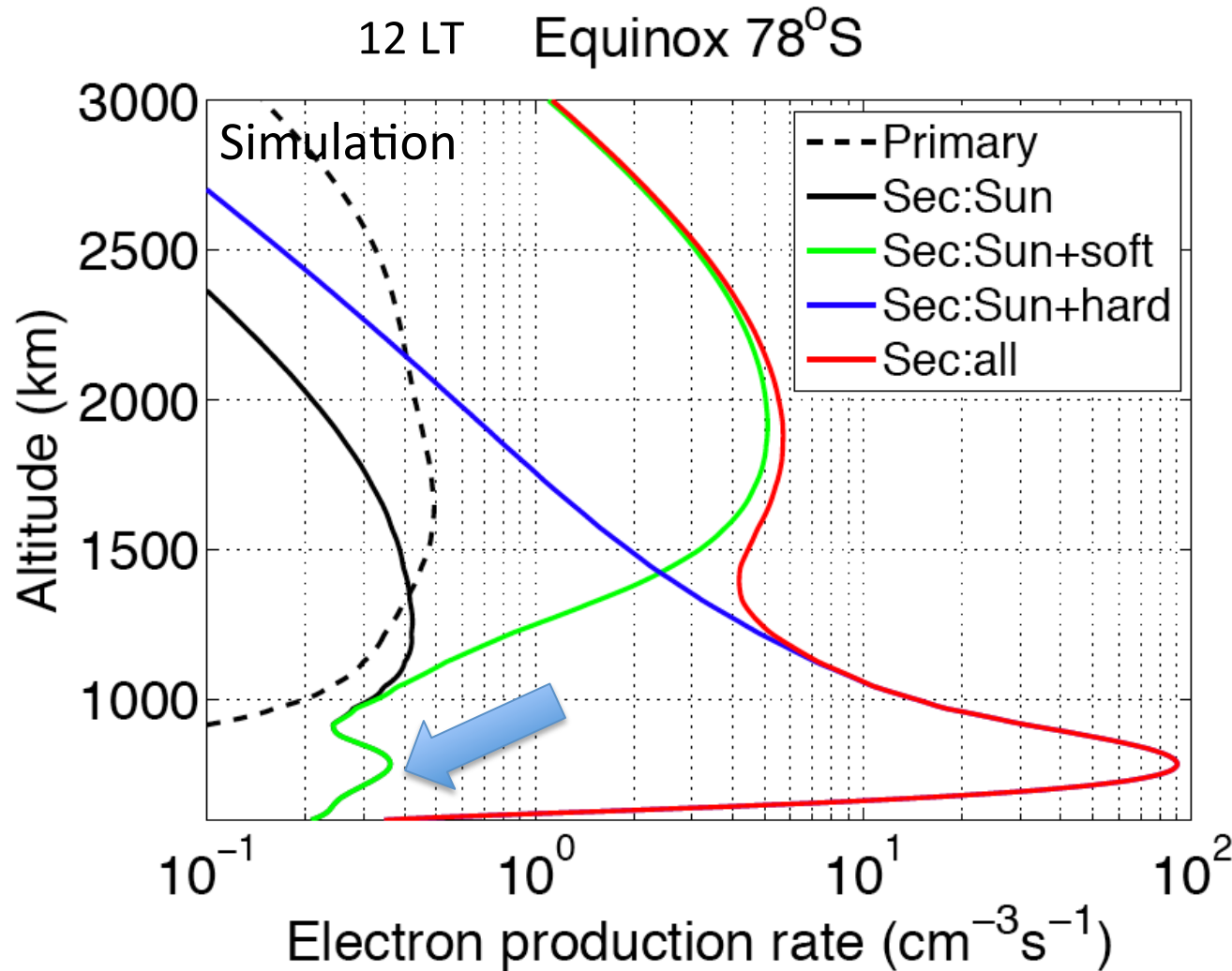


**Auroral electrons – Energy flux  $Q_{prec} = 1 \text{ mW m}^{-2}$**



$$P_e(z) = 4\pi \sum_n n_n(z) \int_{E_{\min}}^{E_{\max}} I_e(E, z) \sigma_n^{ioni}(E) dE \propto Q_{prec}$$

# Solar versus auroral particle deposition



- **Soft component:**  
500 eV,  $0.03 \text{ mW}\cdot\text{m}^{-2}$

- **Hard component:**  
10 keV,  $0.1 \text{ mW}\cdot\text{m}^{-2}$

- **Soft + Hard  
component**

# Continuity equation applied to the ionospheric plasma

## Thermal Ion Continuity Equation

$$\frac{\partial n_i}{\partial t} = P_i - L_i - \nabla \cdot (n_i \underline{u})$$

Production

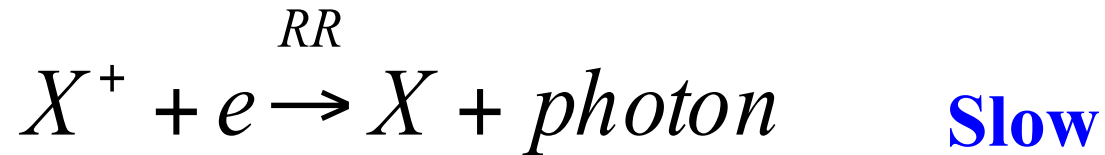
Loss

Transport

$\underline{u}$ , bulk velocity

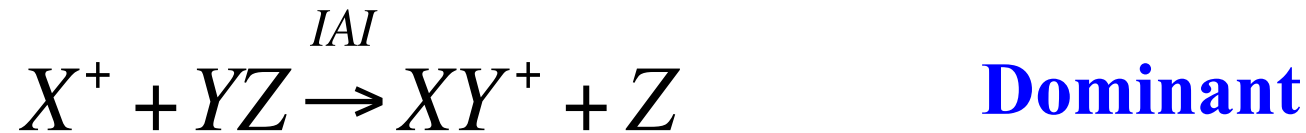
# Chemical loss of atomic ions

- **Radiative Recombination of Atomic Ions**



$$L_{X^+}^{RR} = \alpha_{X^+}^{RR} n_{X^+} n_e \quad \text{Ion loss Rate}$$

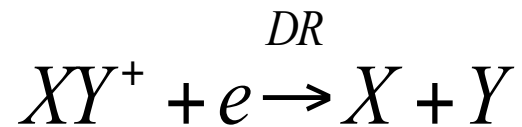
- **Ion-atom Interchange**



$$L_{X^+,YZ}^{IAI} = k_{X^+,YZ}^{IAI} n_{X^+} n_{YZ}$$

## Chemical loss of terminal, molecular ions

- **Dissociative Recombination of e- and molecular ions**



$$L_{XY^+}^{DR} = \alpha_{XY^+}^{DR} n_{XY^+} n_e$$

**Dominant  
for terminal  
ions**

- **If  $XY^+$  is the dominant ion,**

$$L_{XY^+}^{DR} \approx \alpha_{XY^+}^{DR} n_e^2$$

# Photochemical equilibrium

Thermal Ion Continuity Equation

$$\frac{\partial n_i}{\partial t} = P_i(z) - L_i(z) - \nabla \cdot (n_i \underline{u})$$

Production      Loss      Transport  
 $\underline{u}$ , bulk velocity

For molecular, terminal ion i:

$$P_i(z) = \alpha_i(Te) n_i(z) n_e(z) \approx \alpha_i(Te) n_e^2(z)$$

- Chemical loss timescales:**

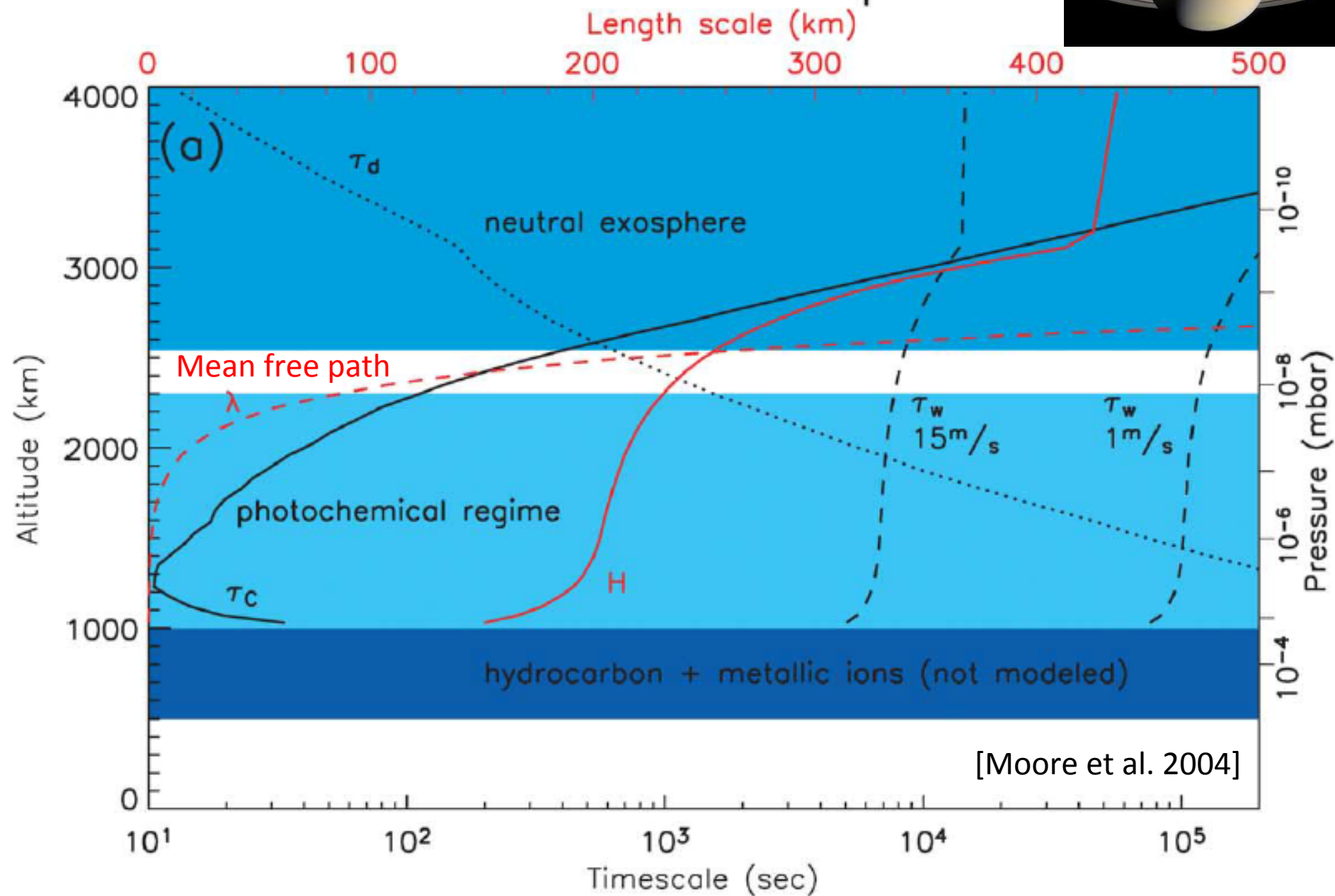
- Atomic ion or non-terminal ion:  $t_c = \frac{1}{k n_{YZ}}$        $X^+ + YZ \xrightarrow{IAI} X + YZ^+$
- Terminal molecular ion:  $t_c = \frac{1}{\alpha n_e}$        $XY^+ + e^- \xrightarrow{DR} X + Y$

- Transport loss timescales:**

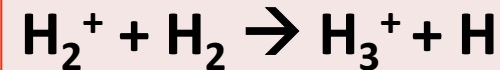
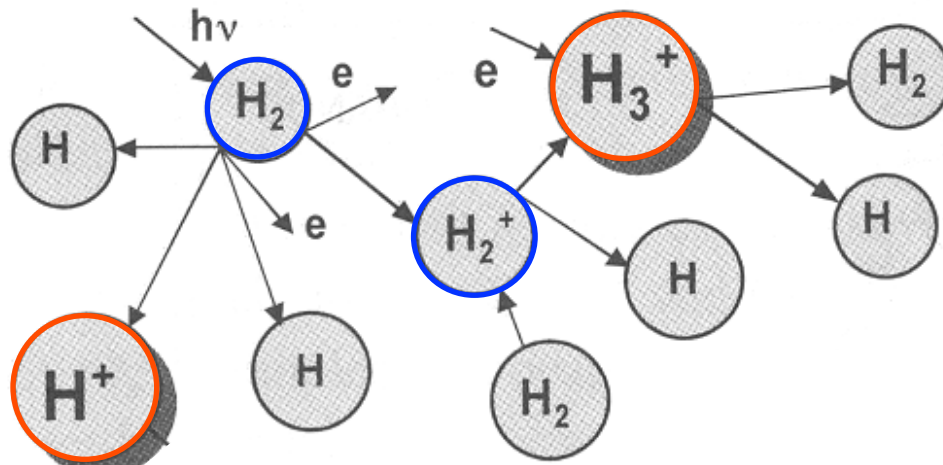
- Diffusion timescale:  $t_d = \frac{H_a^2}{D_a \sin^2 I}$
- Timescale associated with the neutral wind:  $t_w = \frac{H_p}{U \sin I \cos I}$

with  $H_a$  and  $H_p$  are the atmosphere and plasma scale heights (typical length scales),  $I$  is the magnetic field dip angle (cp w/ horizontal),  $U$  is the meridional (N/S) neutral wind speed and  $D_a$  is the ion-neutral diffusion coefficient.

# Photochemical equilibrium at Saturn



# Photo-chemistry in an H<sub>2</sub> atmosphere



$$k_0 = 2.0 \times 10^{-9} \text{ cm}^3 \text{ s}^{-1}$$



$$\alpha_0 = 1.73 \times 10^{-6} \times \text{Te}^{-0.5} \text{ cm}^3 \text{ s}^{-1}$$

with Te in K.

- Charge exchange reaction  $\text{H}^+ + \text{H}_2(\nu \geq 4) \rightarrow \text{H}_2^+ + \text{H}$  (1)

controls the abundance of H<sub>3</sub><sup>+</sup> as it is quickly followed by:



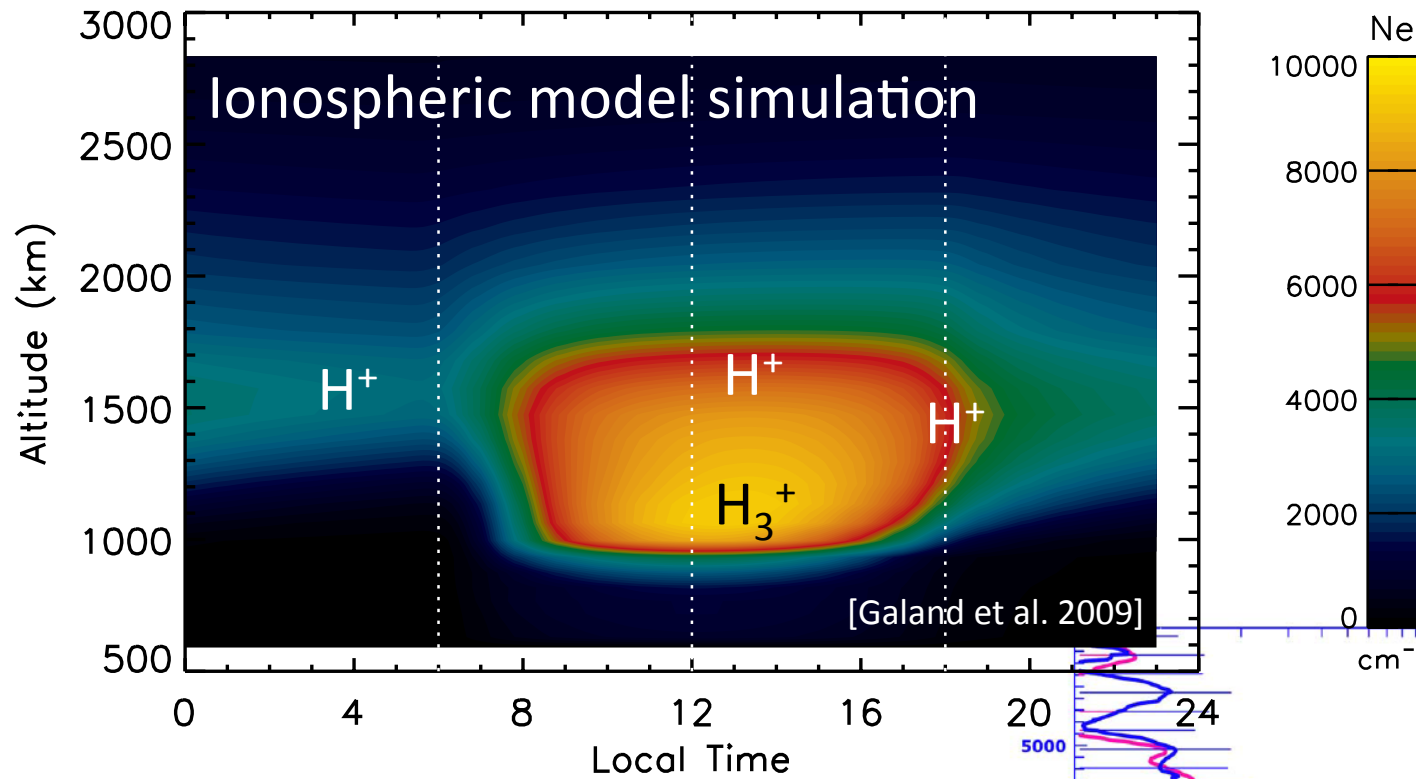
- Reaction rate  $k_1^* = k_1 [\text{H}_2(\nu \geq 4)] / [\text{H}_2]$

– Large [H<sub>2</sub>(ν ≥ 4)] → large k<sub>1</sub><sup>\*</sup> → more H<sup>+</sup> converted in H<sub>3</sub><sup>+</sup> → decrease in ionospheric densities

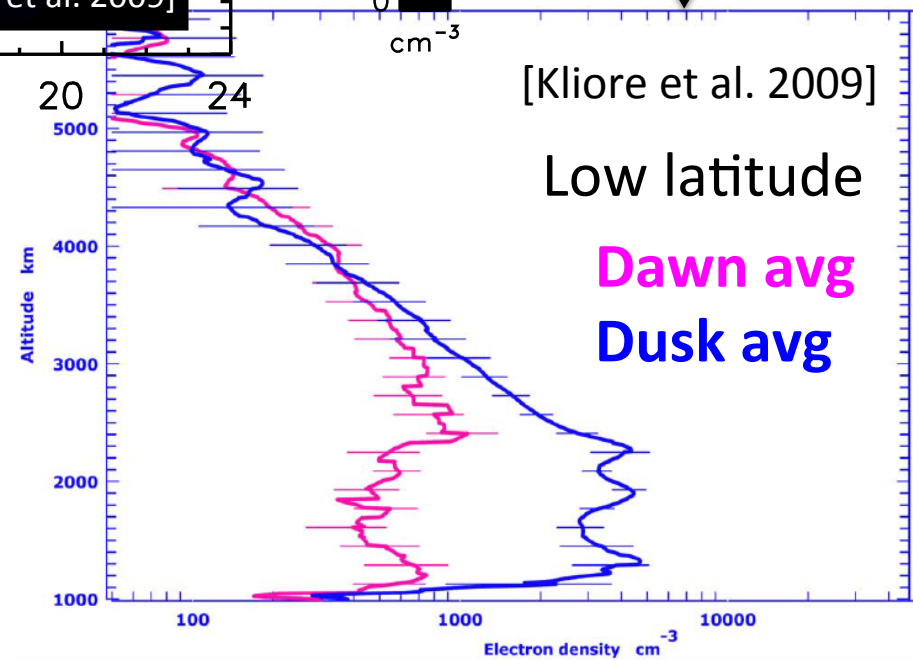
➤ k<sub>1</sub> = 10<sup>-9</sup> cm<sup>3</sup> s<sup>-1</sup> [Huestis, 2008]



# Dawn-dusk asymmetry



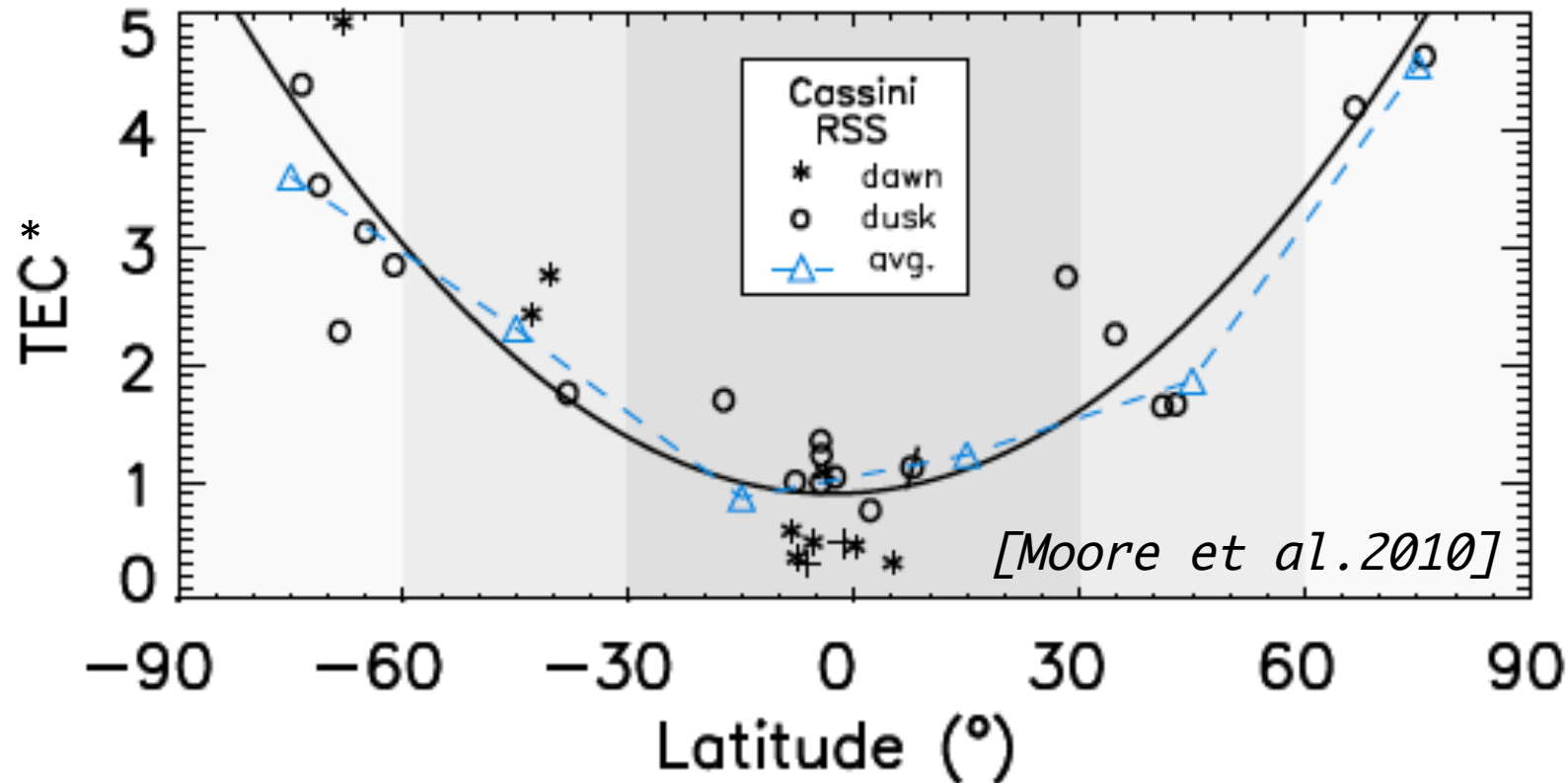
From Cassini/  
Radio Science  
Subsystem  
(RSS) data



# The ionosphere of Saturn



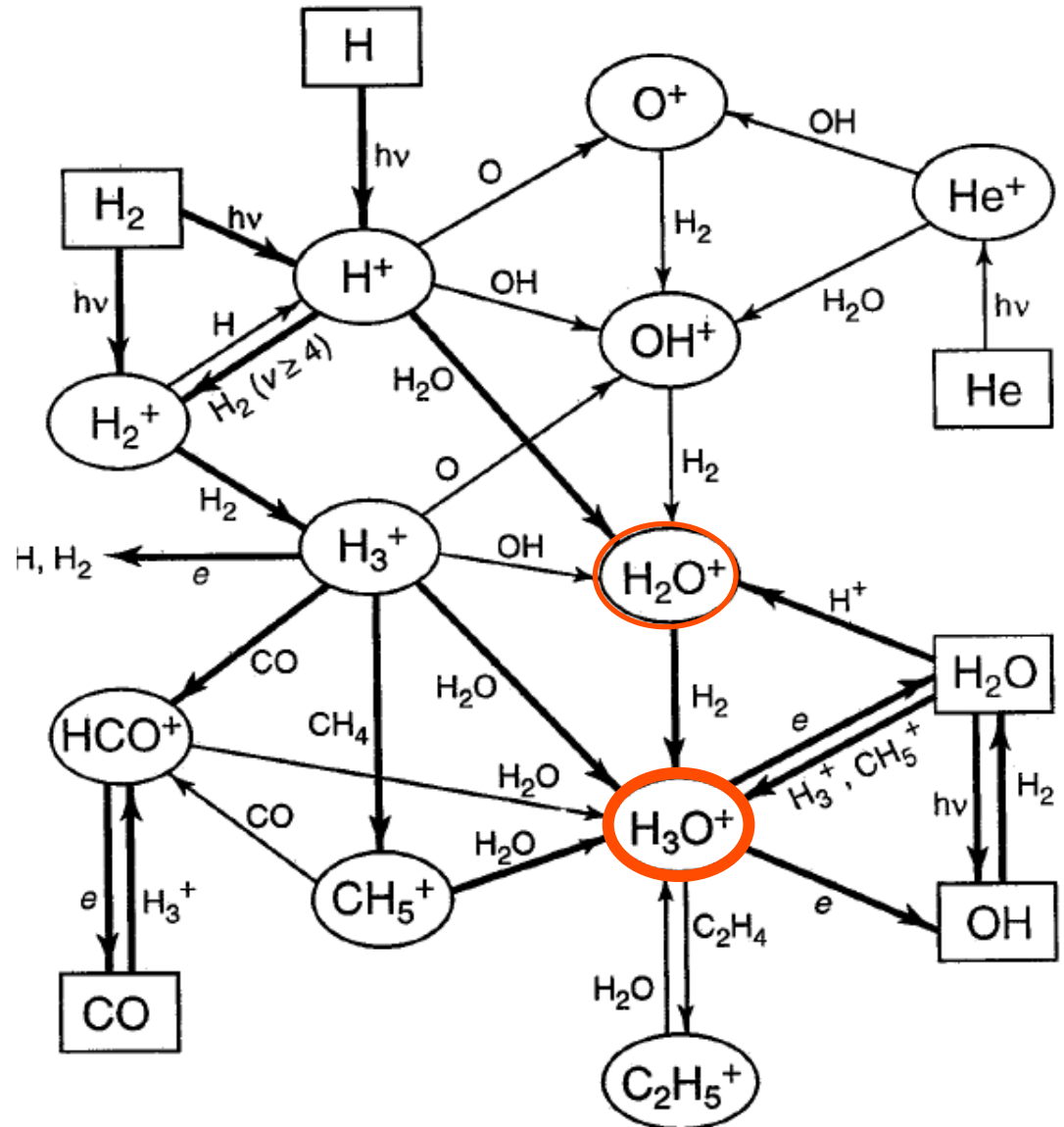
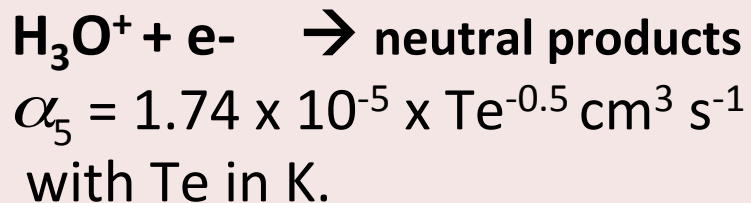
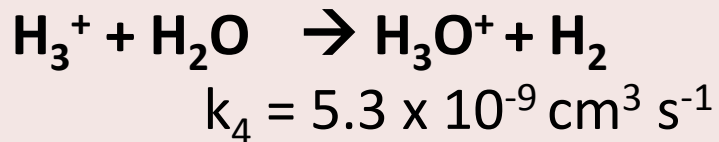
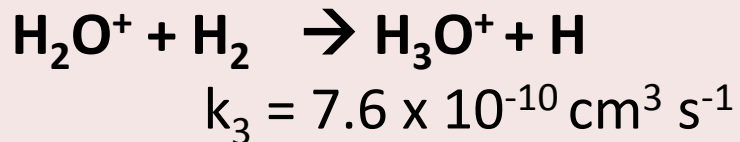
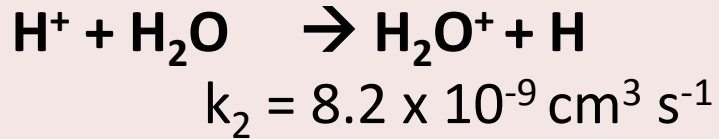
From Cassini/Radio Science Subsystem (RSS) data



\*Total Electron Content [1 TEC unit =  $10^{16}m^{-2}$ ]

Why is there a minimum in TEC at low latitudes?...

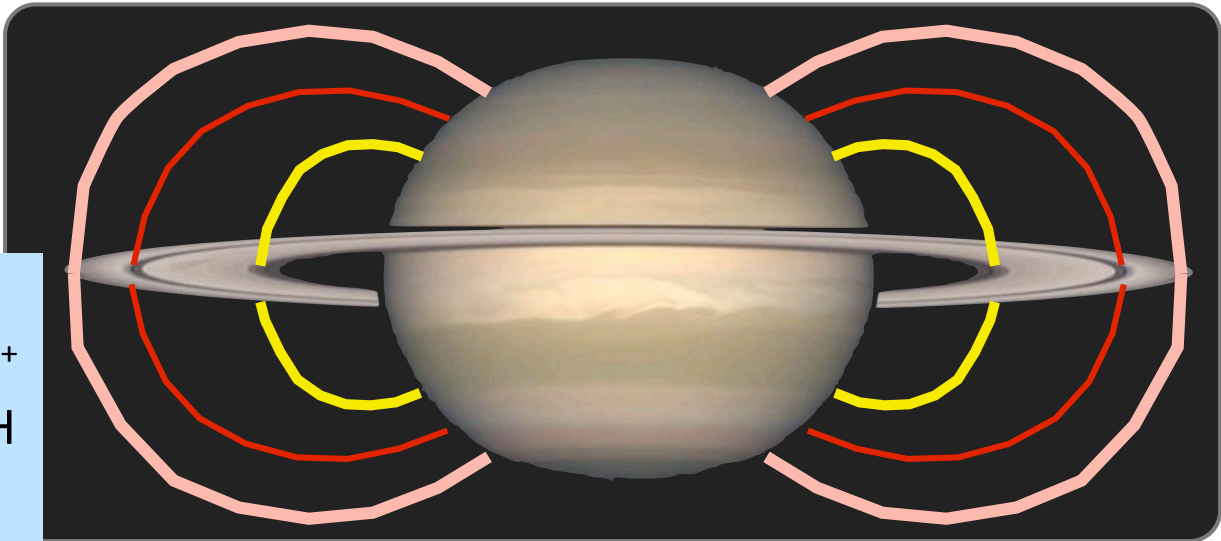
# Photochemistry in Gas Giant atmospheres



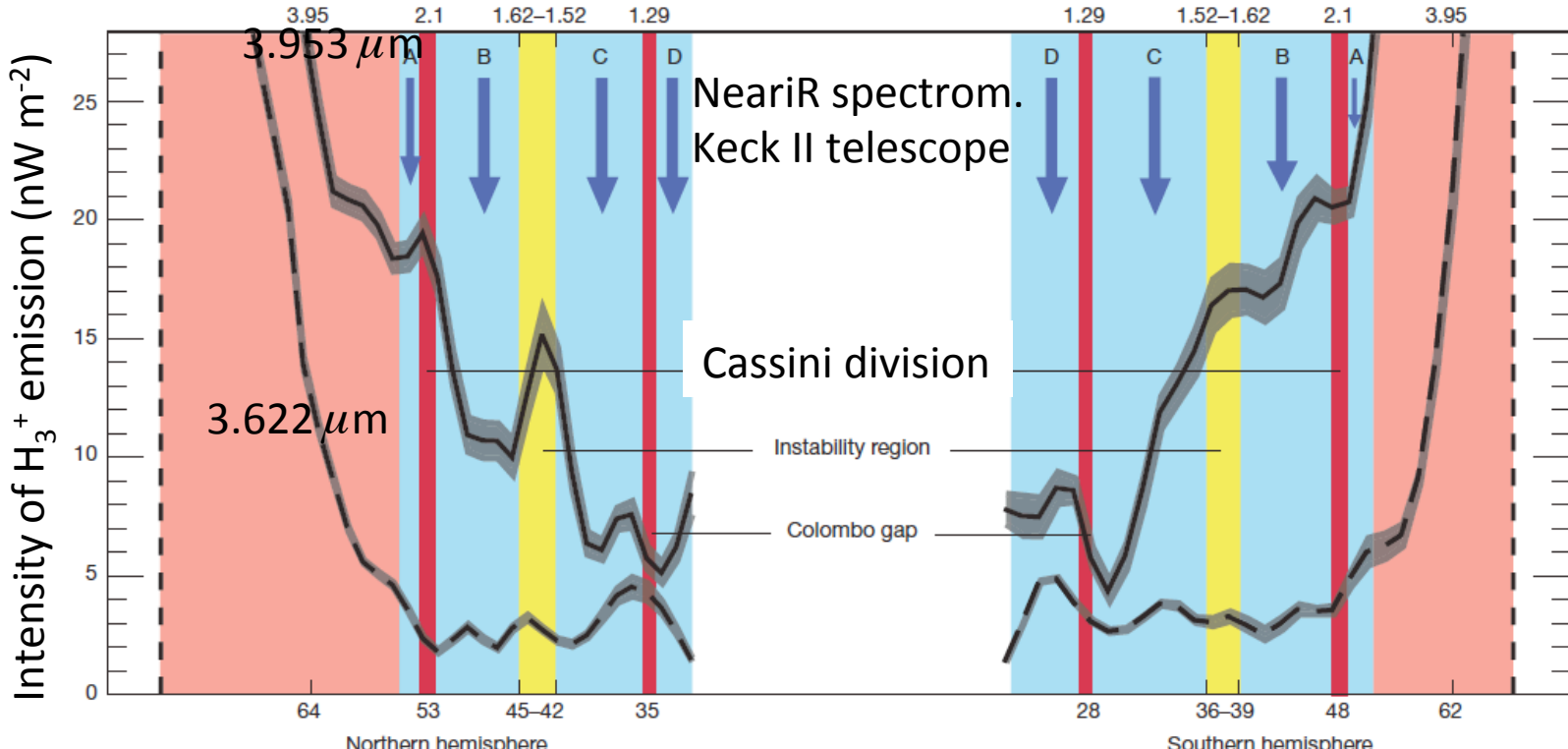
[Moses and Bass 2000]

# Ring rain at Saturn

From the rings:  $\text{H}_2\text{O}^+$   
 React with  $\text{H}_2 \dots \rightarrow \text{H}_3\text{O}^+$   
 $\text{H}_3\text{O}^+ + \text{e}^- \rightarrow \text{H}_2\text{O}$  or  $\text{OH}$   
 $\text{OH} + \text{H}_2 \rightarrow \text{H}_2\text{O} + \text{H}$   
 $\text{H}_3^+ + \text{H}_2\text{O} \rightarrow \text{H}_3\text{O}^+ + \text{H}_2$



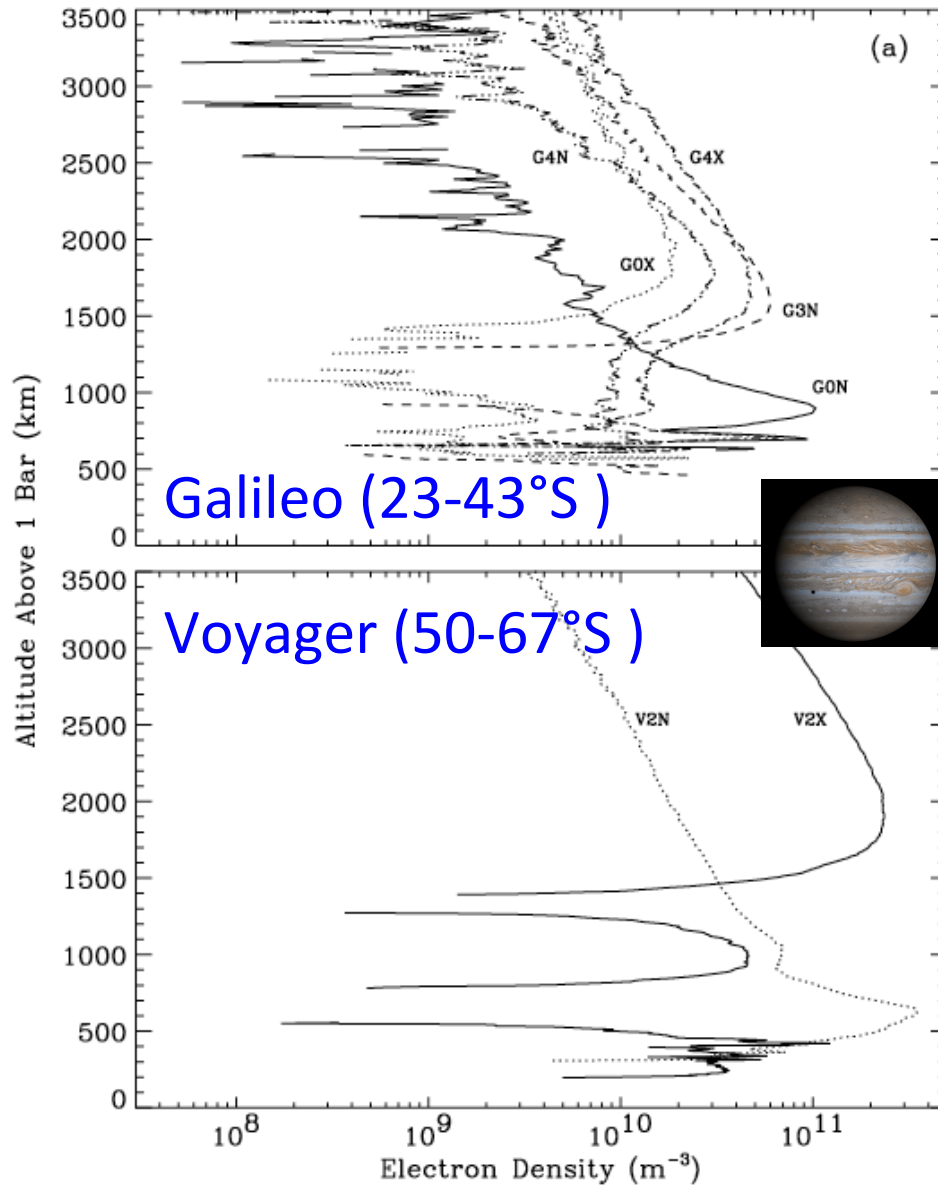
Equatorial magnetic mapping ( $R_s$ )



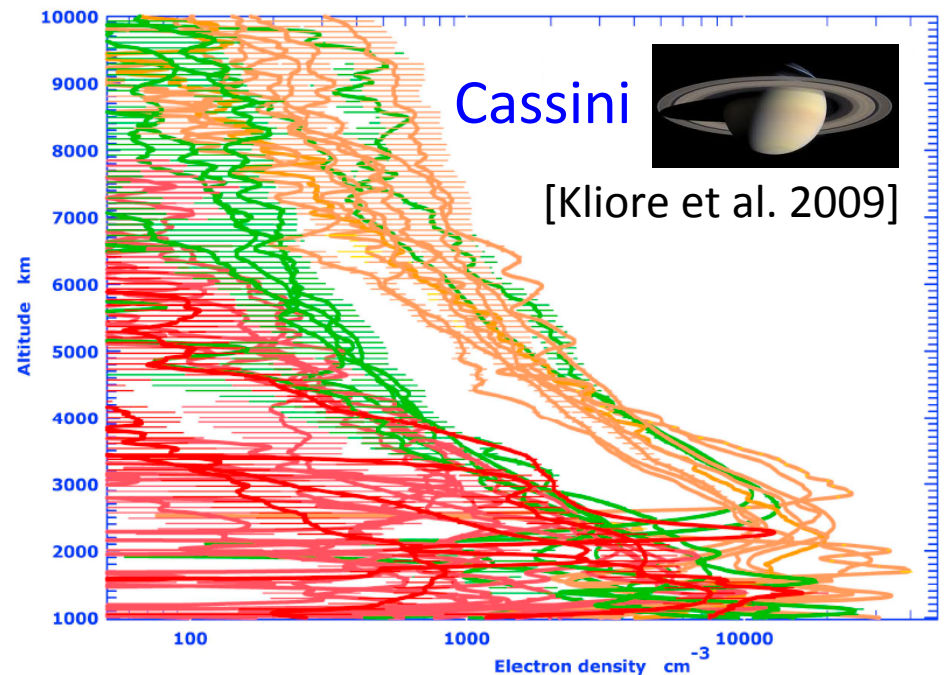
O'Donoghue et al. (Nature, 2013)

# Measured ionospheric profiles at Jupiter and Saturn

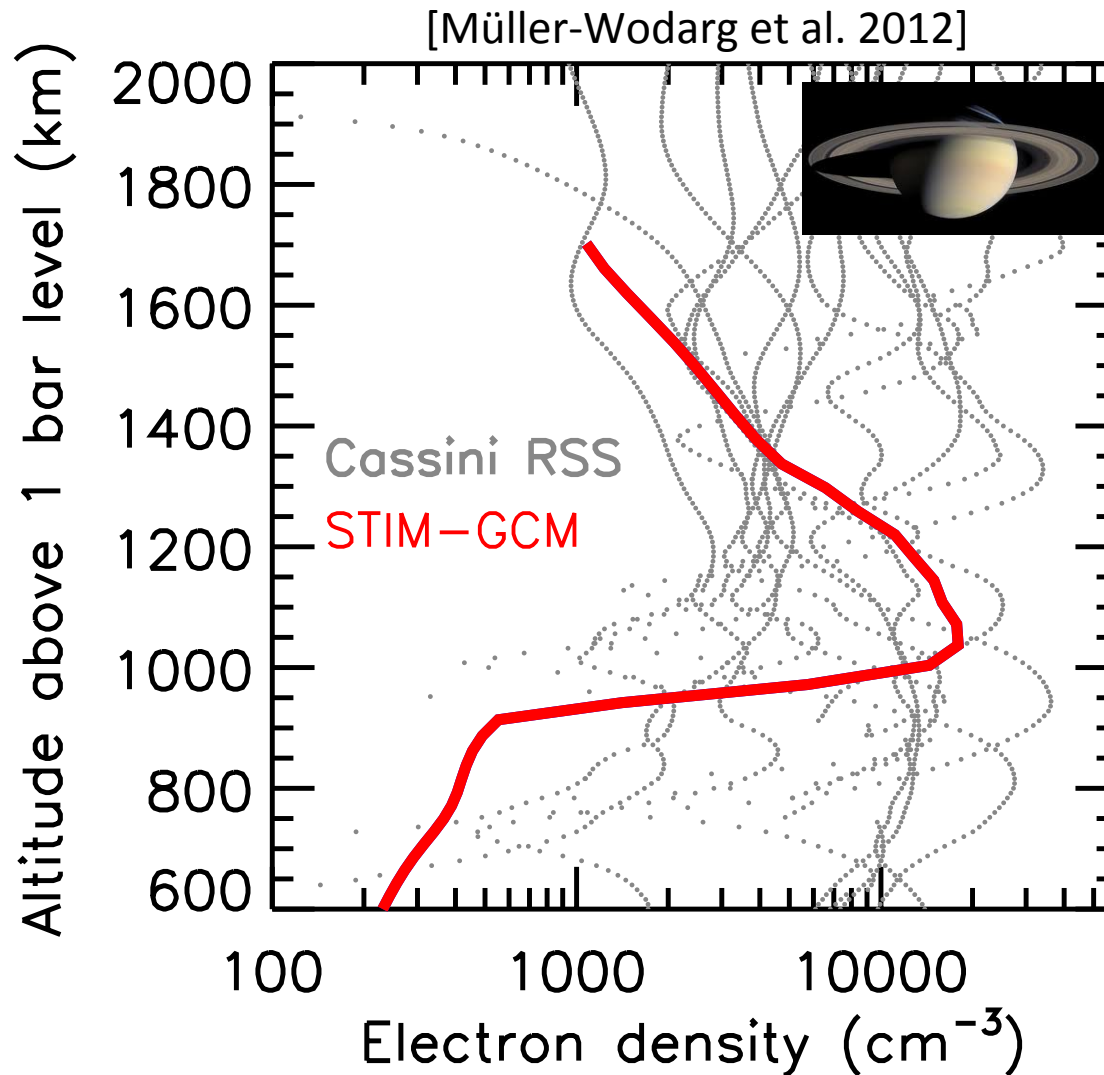
[Yelle and Miller 2004]



Low latitude  
Mid latitude  
High latitude



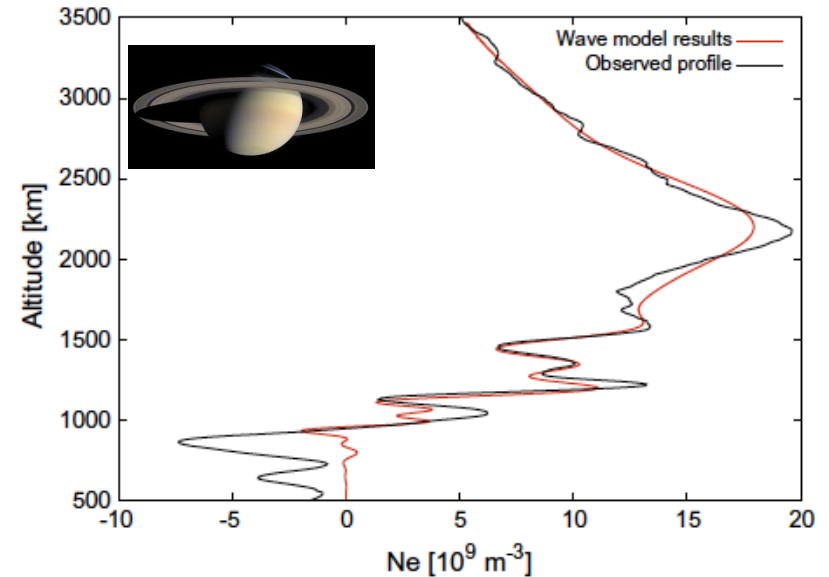
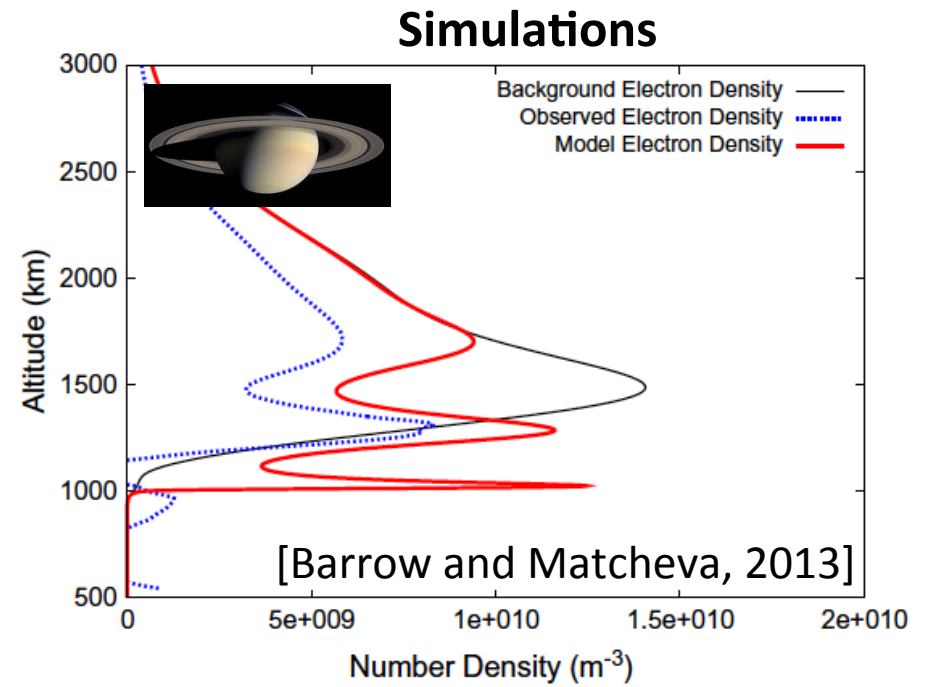
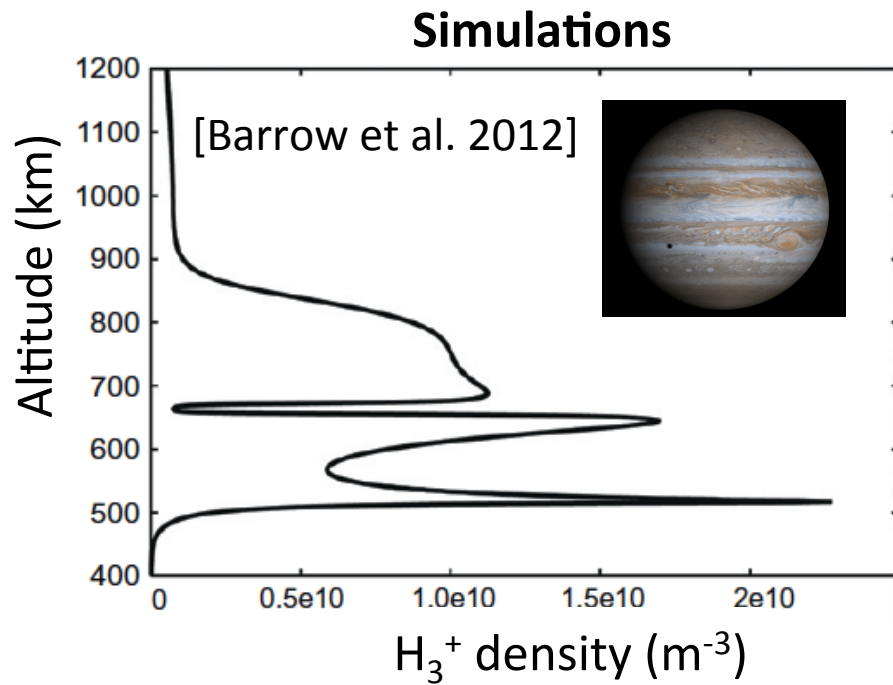
# Vertical structure of the ionosphere



\* Cassini/Radio Science Subsystem

\* Saturn Thermosphere-Ionosphere Model – General Circulation Model

# Effect of atmospheric gravity waves



# TAKE HOME MESSAGE: Ionosphere

- ***Ionization sources:***
  - EUV solar radiation and magnetospheric, energetic particles
- ***Photo-absorption versus particle deposition:***
  - Altitude of absorption of solar radiation driven by:
    - Photo-absorption cross section
    - Column density of absorbers
  - Altitude of deposition of auroral electrons driven by:
    - Column density of atmosphere (primarily)
- ***Ionospheric composition:***
  - Dominant ion species photo-produced → usually not most abundant: need to take into account chemistry
- ***Ionospheric density:***
  - Limitation in electron density estimates: H<sub>2</sub> vibrat., ion drift, water inflow
- ***Vertical/latitudinal structure:***
  - Ring contribution to latitudinal and vertical structure
  - Atmospheric gravity wave contribution to vertical structure



# SECOND LECTURE

# Outline

## **1. Setting the scene**

- Thermosphere/Ionosphere
- Energy sources

## **2. Ionosphere at the giant planets**

- What? How? Where? How much?

## **3. Auroral emissions**

## **4. Magnetosphere-ionosphere-thermosphere coupling**

- Energy crisis at the giant planets

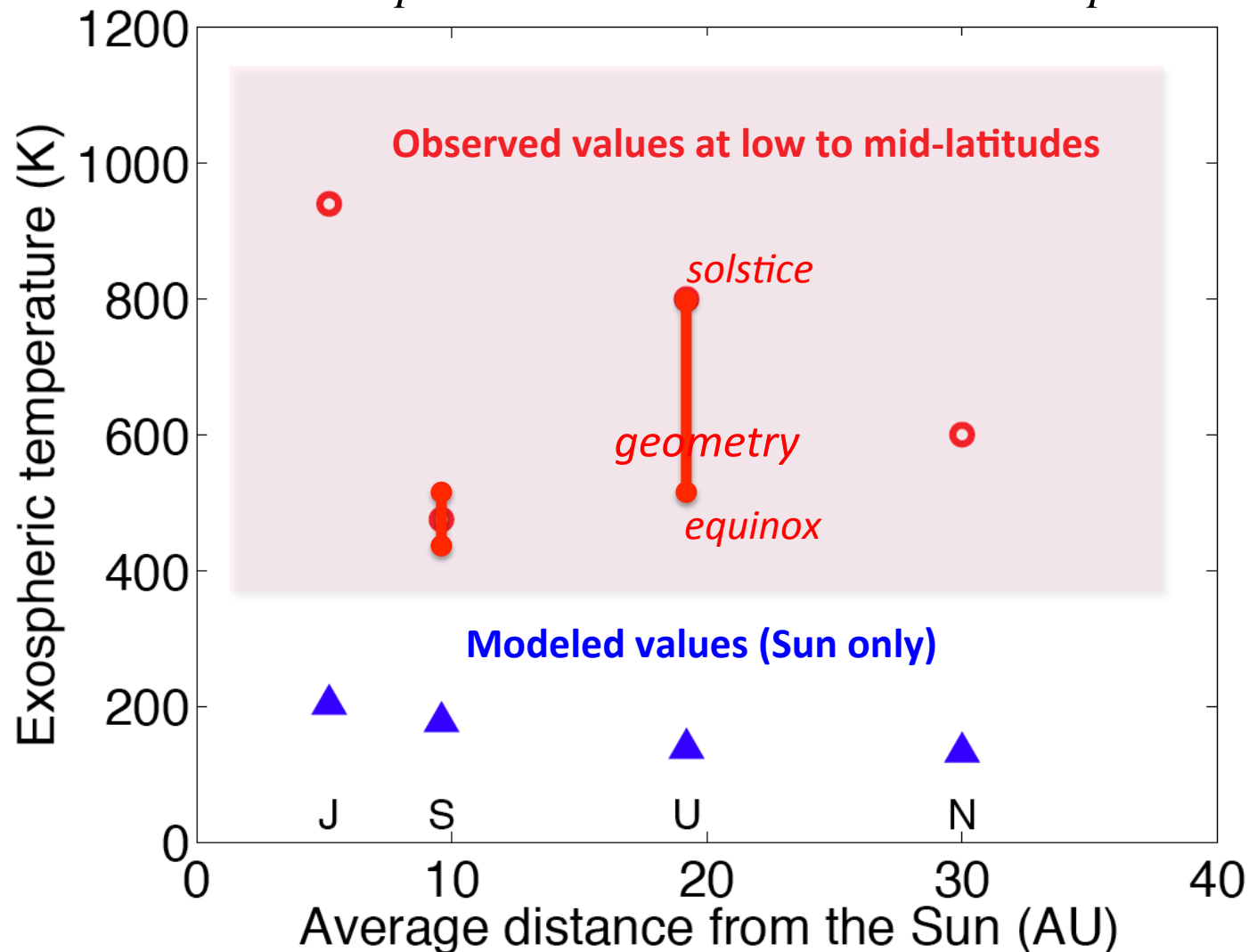
## **5. Future prospect**

## Open questions

- *What drive **ionospheric structure and variability** at the giant planets? (see 1<sup>st</sup> lecture)*
- *What causes the **energy crisis** at the giant planets?*
- *What drive the **hemispheric differences** observed at Saturn in the magnetosphere and auroral, ionospheric regions?*
- *Is the **variable rotation rate** observed in the Saturnian magnetosphere linked to atmospheric dynamics?*

# Energy Crisis at the Giant Planets

$$T_{\text{exosphere}}^{\text{computed with sun only}} \ll T_{\text{exosphere}}^{\text{observed}}$$



[After Yelle and Miller, 2004; Melin et al., 2011, 2013]

# Energy crisis at the giant planets

## ◆ HEATING SOURCES: forcing from above/below

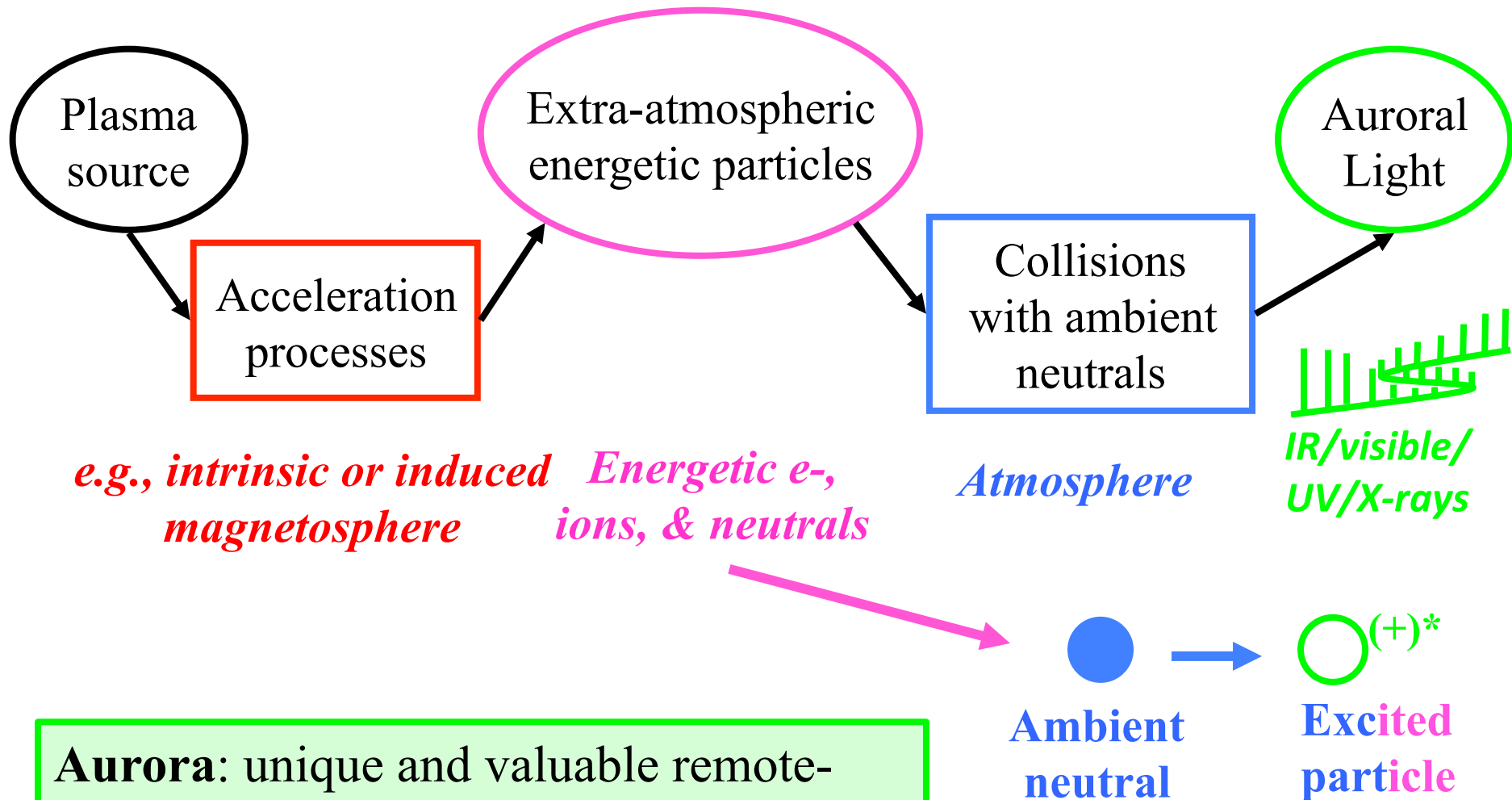
- **Solar heating** through excitation/dissociation/ionization + exothermic chemical reactions
- **Particle heating** via collisions + chemistry  
[e.g., Waite et al. 1983, 1987; Grodent et al., 2001]
- **“Ionospheric Joule heating”** via auroral electrical currents and ion-drag heating [Vasyliūnas and Song, 2005] **at high latitudes**
- **Dissipation of upward, propagating waves** (such as gravity waves, ...) [Matcheva and Strobel 1999, Hickey et al. 2000; Barrow et al. 2012]
- ...

[\*: Strobel, 2002]

- **Solar EUV/FUV heating\***: 0.5 TW (Earth), 0.8 TW (Jupiter), 0.2 TW (Saturn)
- **Auroral part./Joule heating\***: 0.08 TW (Earth), 100 TW (Jupiter), 5-10 TW (Saturn)

# Auroral emissions

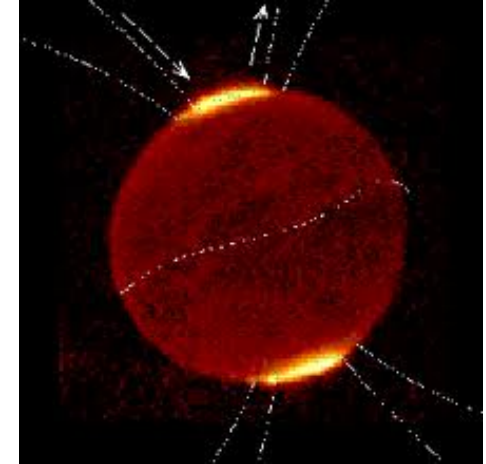
**Aurora** is the photo-manifestation of the interaction of energetic extra-atmospheric electrons, ions, and neutrals with an atmosphere.



**Aurora:** unique and valuable remote-sensing probe of the Solar System.

# Protonated molecular hydrogen: $\text{H}_3^+$

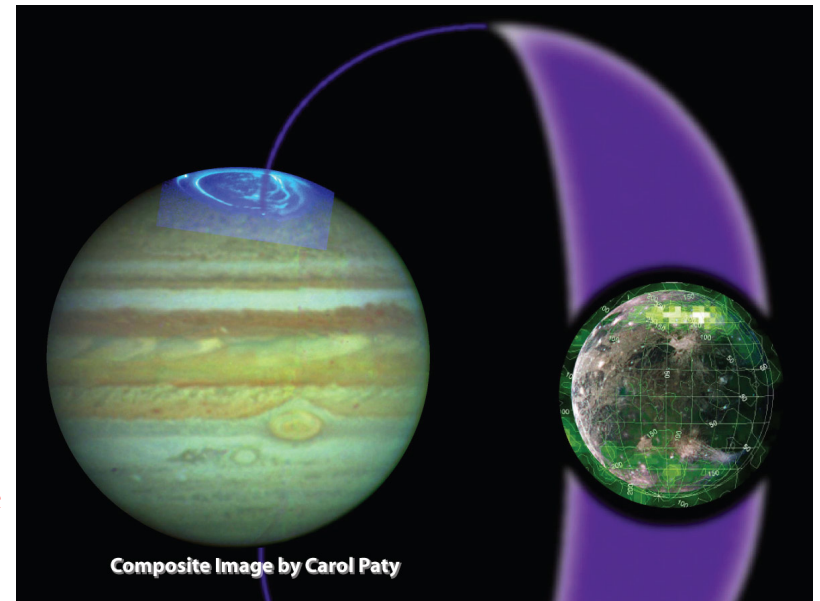
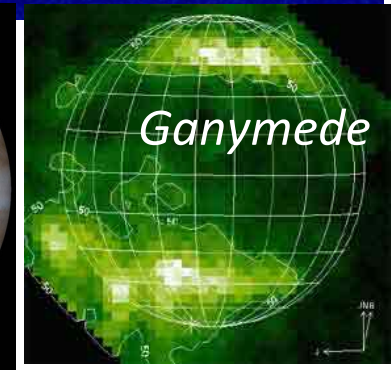
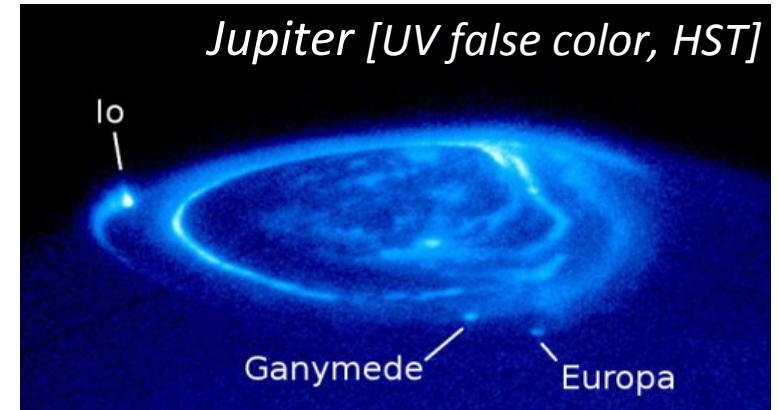
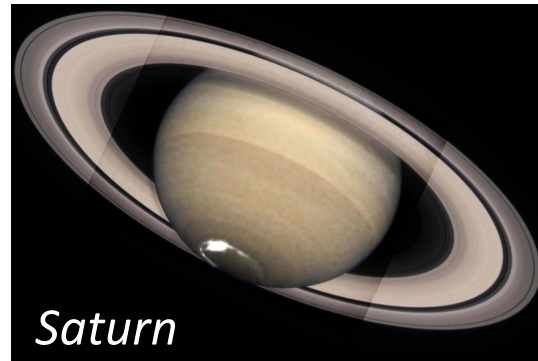
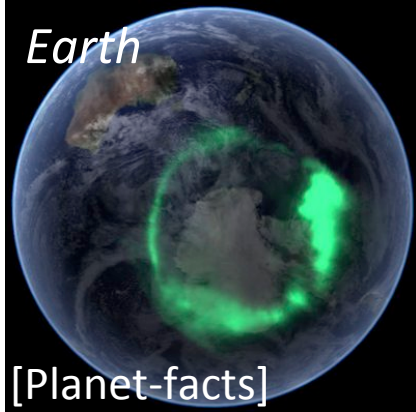
- First time detected in the universe through its IR, thermal emissions
  - **Jupiter aurora** with Voyager/UVS (Drossart et al. 1989)
- Thermal emissions: tracer of **ionosphere-thermosphere response** to magnetospheric processes
  - ~10 min lifetime
  - Destroyed by hydrocarbons:  $\text{H}_3^+ + \text{X} \rightarrow \text{H}_2 + \text{XH}^+$
- Importance of  $\text{H}_3^+$  as a **coolant** (as  $\text{CO}_2$  at Earth)
  - Efficient thermostat at Jupiter
  - Exoplanet too close, fail to cool, as  $\text{H}_2$  dissociated, no  $\text{H}_3^+$  produced.



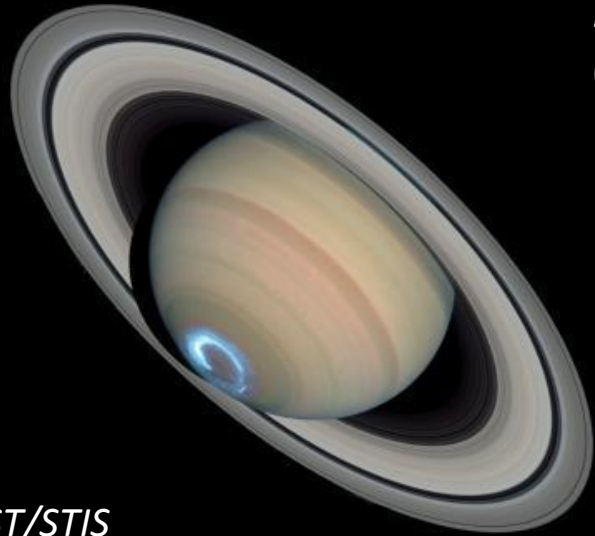
[Image: J. Connerney, T.Satoh, NASA Infrared Telescope Facility]



# Auroral emissions in the Solar System



**Aurora:**  
**“The TV-screen of the I-T-M coupling”**



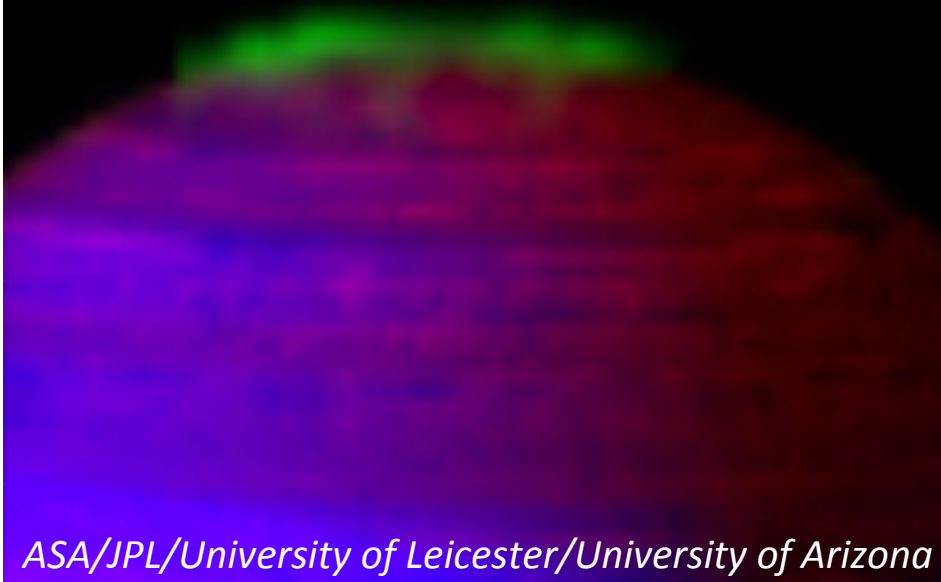
*HST/STIS [NASA/ESA/Jonathan Nichols  
(University of Leicester)]*



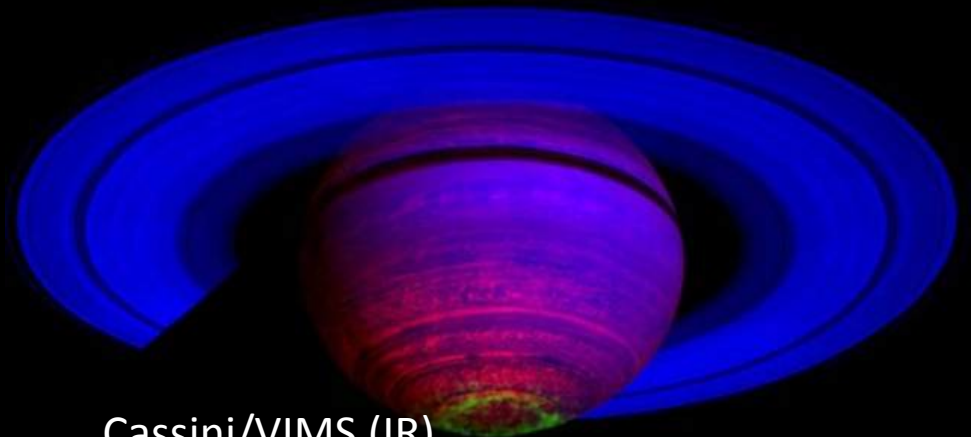
*Cassini/UVIS (UV)  
[UVIS team]*

*HST/STIS  
NASA/ESA/John Clarke (Boston University)*

UV spectroscopic analysis → Particle characteristics  
IR spectroscopic analysis →  $H_3^+$  density & temperature



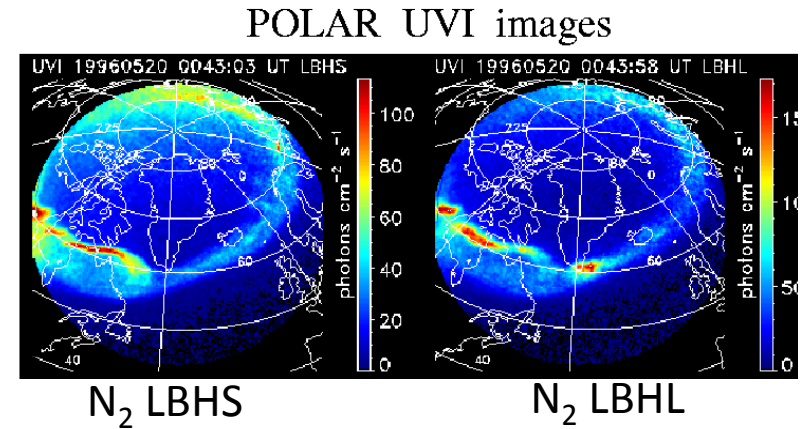
*ASA/JPL/University of Leicester/University of Arizona*



*Cassini/VIMS (IR)  
NASA/JPL/Tom Stallard (Univ. Leicester)/*

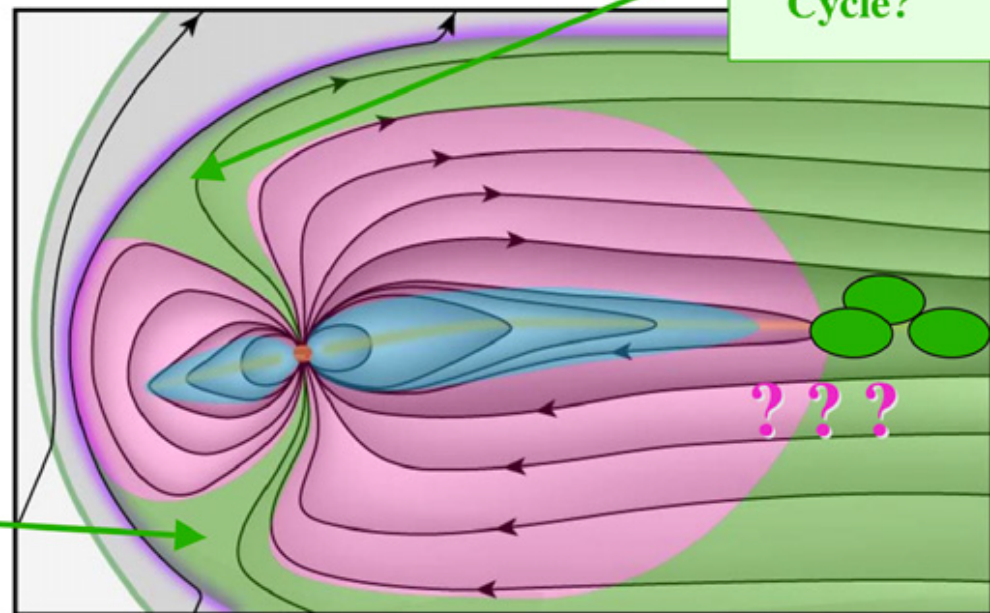
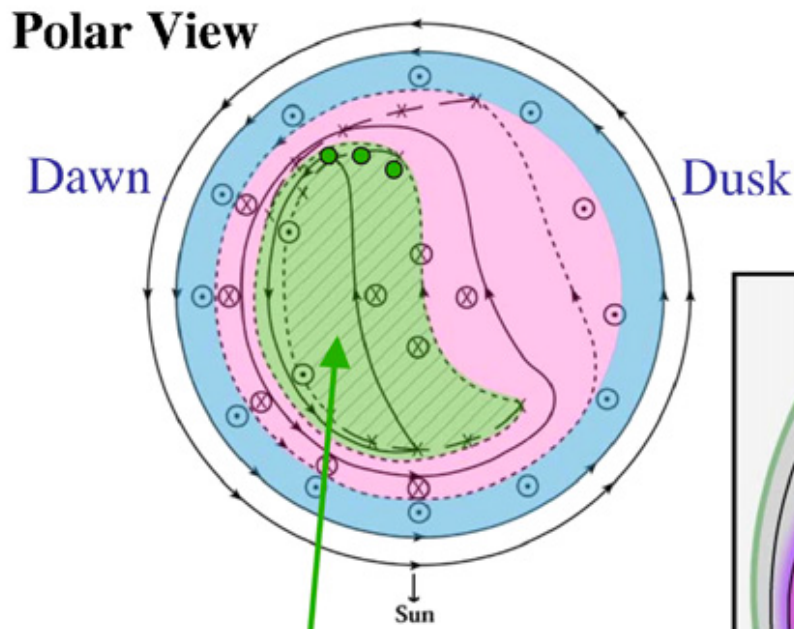
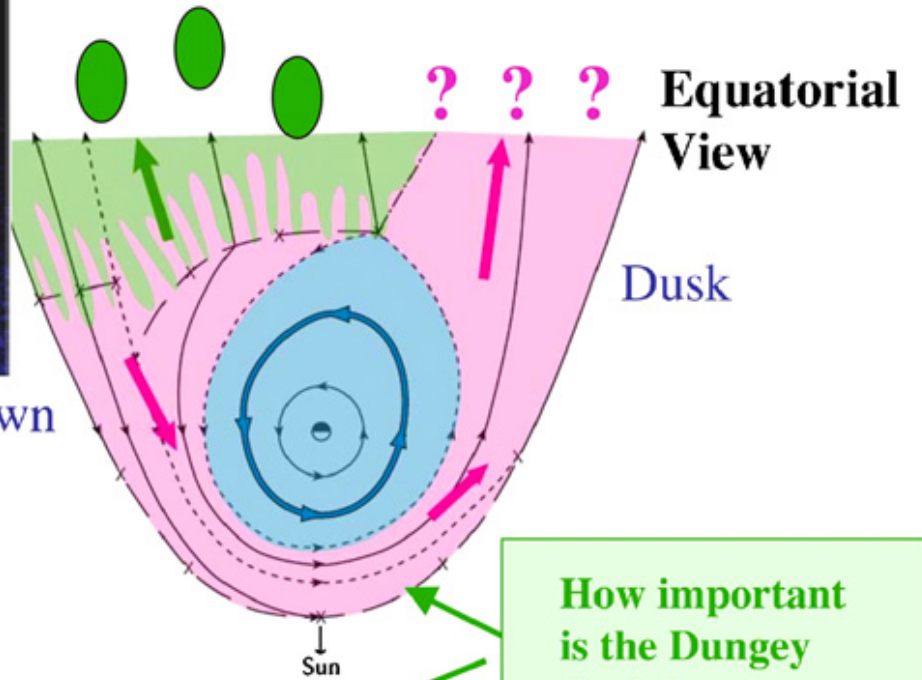
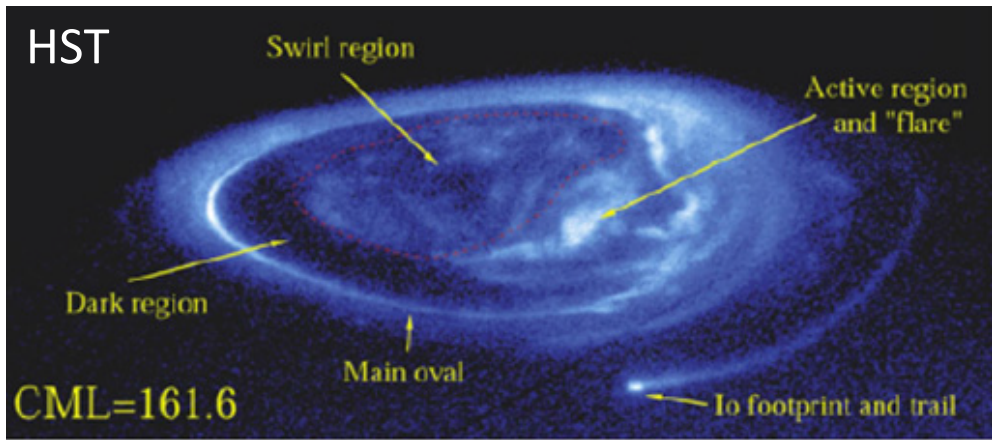
## AURORAL SPECTROSCOPIC ANALYSIS

- Identification of energetic particle type
- Assessment of  $(E_m, Q_{prec})$  of energetic particles
- ✓ *Supported by comprehensive modeling*



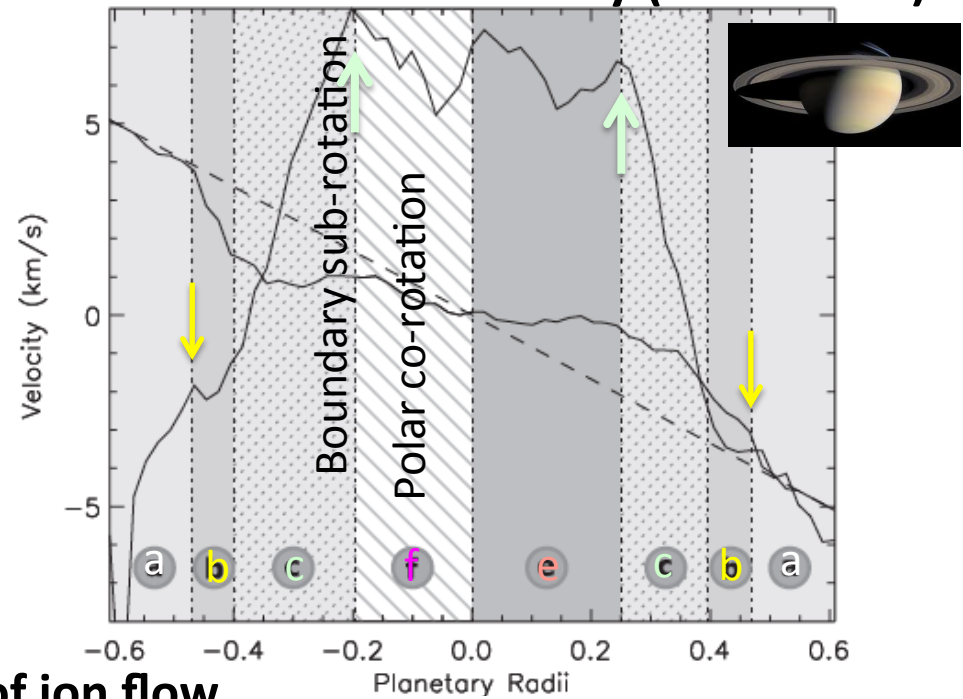
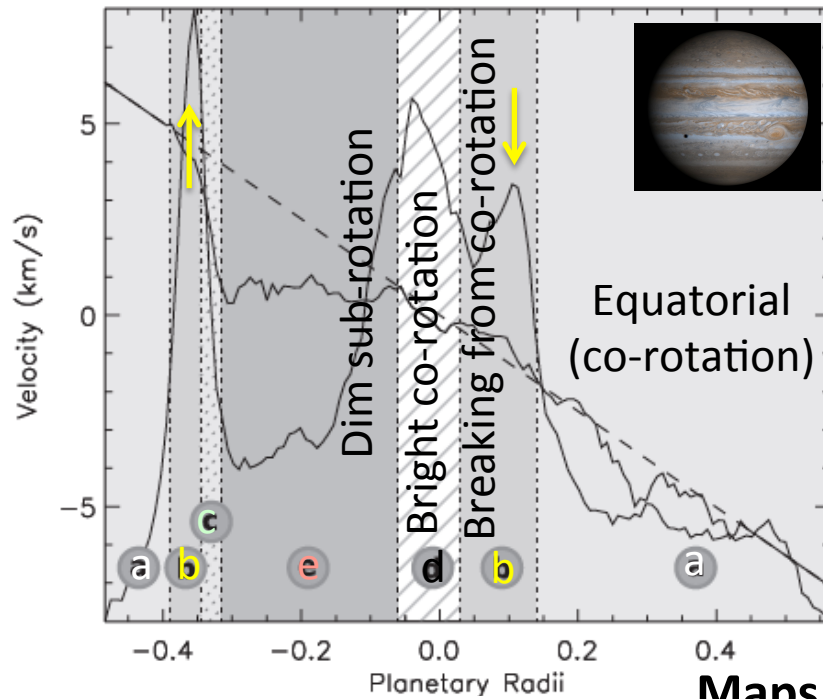
COLOR RATIO	Earth	Jupiter, Saturn
Two spectral bands chosen in:	N <sub>2</sub> LBH	H <sub>2</sub> Lyman and Werner
One band strongly absorbed by:	O <sub>2</sub> ( $< 160$ nm)	CH <sub>4</sub> ( $< 140$ nm)
Electron energy range covered	0.2 – 20 keV	~10 to 200 keV
Type of aurora identified:	Electron aurora (discrete only)	Electron aurora (diffuse + discrete)

✓ *Similar techniques can be applied at various planets  
BUT different limitations on the product*

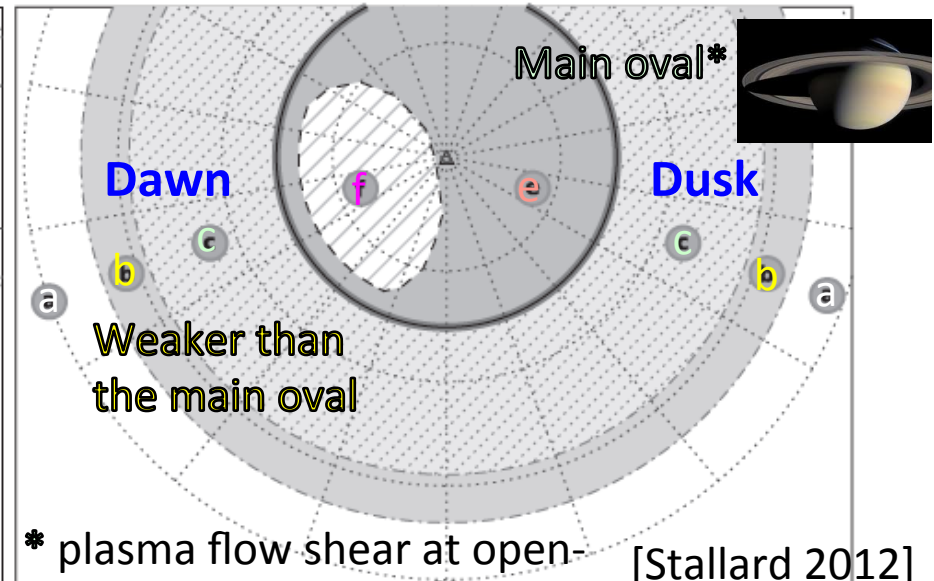
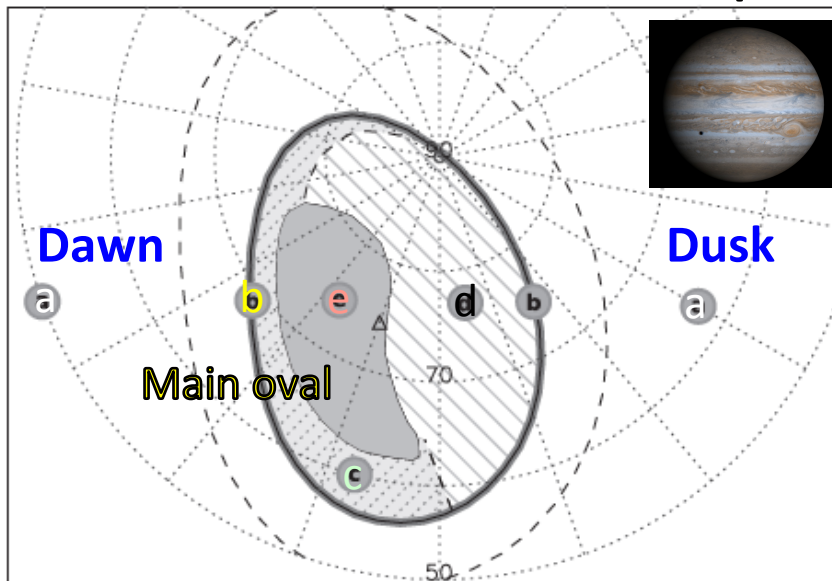


# IR AURORAL SPECTROSCOPIC ANALYSIS

$H_3^+$  line-of signed velocity and normalize intensity (NASA-IRTF)



## Maps of ion flow



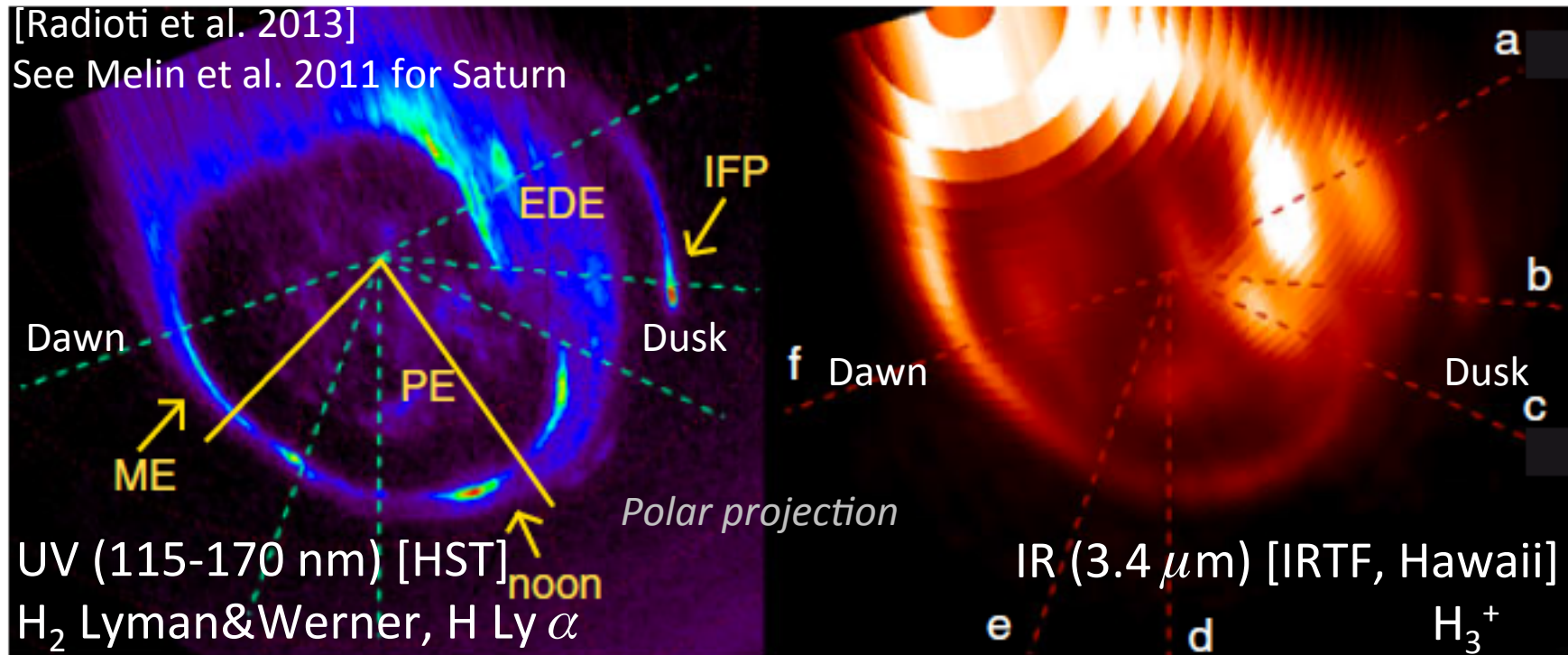
\* plasma flow shear at open-closed field line boundary [Stallard 2012]

# Simultaneous UV/IR auroral emissions



[Radioti et al. 2013]

See Melin et al. 2011 for Saturn



UV (115-170 nm) [HST]

H<sub>2</sub> Lyman & Werner, H Ly  $\alpha$

IR (3.4  $\mu\text{m}$ ) [IRTF, Hawaii]

H<sub>3</sub><sup>+</sup>

- Direct impact of auroral e<sup>-</sup> on H<sub>2</sub>
- Instantaneous
- Extend down to 250 km (visible)

- Thermal emissions ([H<sub>3</sub><sup>+</sup>], T<sub>H<sub>3</sub><sup>+</sup></sub> ~ T<sub>neutral</sub> (lifetime ~10 min))
- > 1000 km (destruction of H<sub>3</sub><sup>+</sup> by hydrocarbons below homopause)

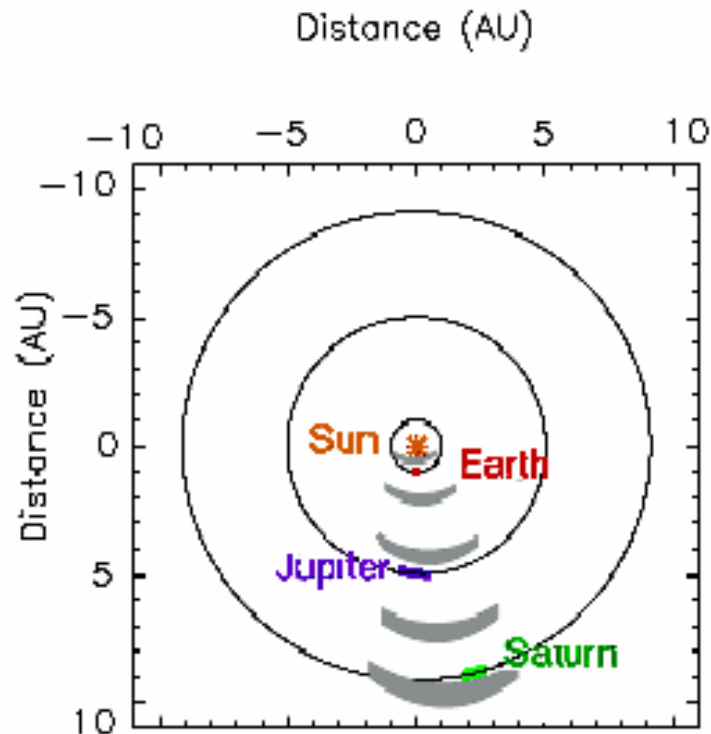
**Main oval:** Spatial distribution and temporal behavior of UV and IR emissions differ → IR emissions above homopause

**All** → IR emissions influenced by H<sub>3</sub><sup>+</sup> heating + time-delays.

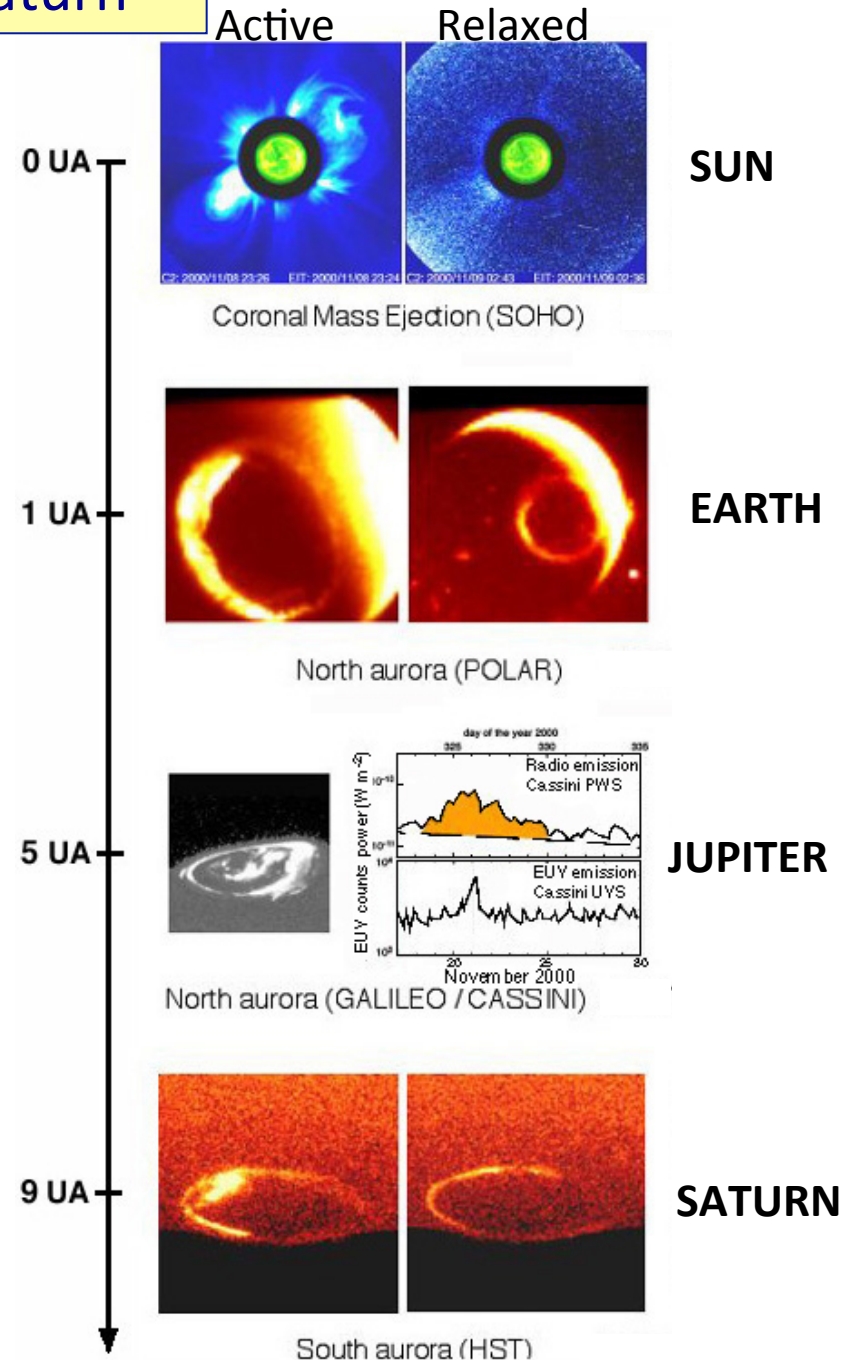
# Interplanetary shock to Earth, Jupiter, Saturn

First synoptic view of a CME-driven interplanetary shock hitting the Earth, Jupiter and Saturn, triggering major – but **different** – auroral responses at all three planets

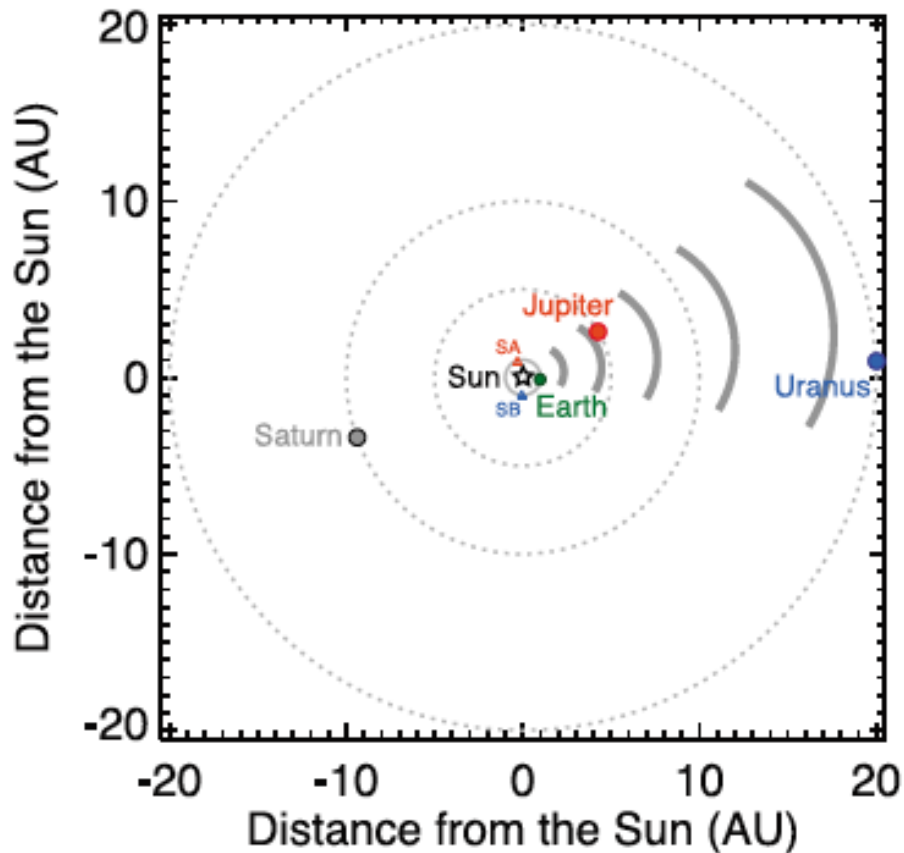
→ Highlights the difference in setting and allows us to learn more about sw-MI coupling by comparing different responses



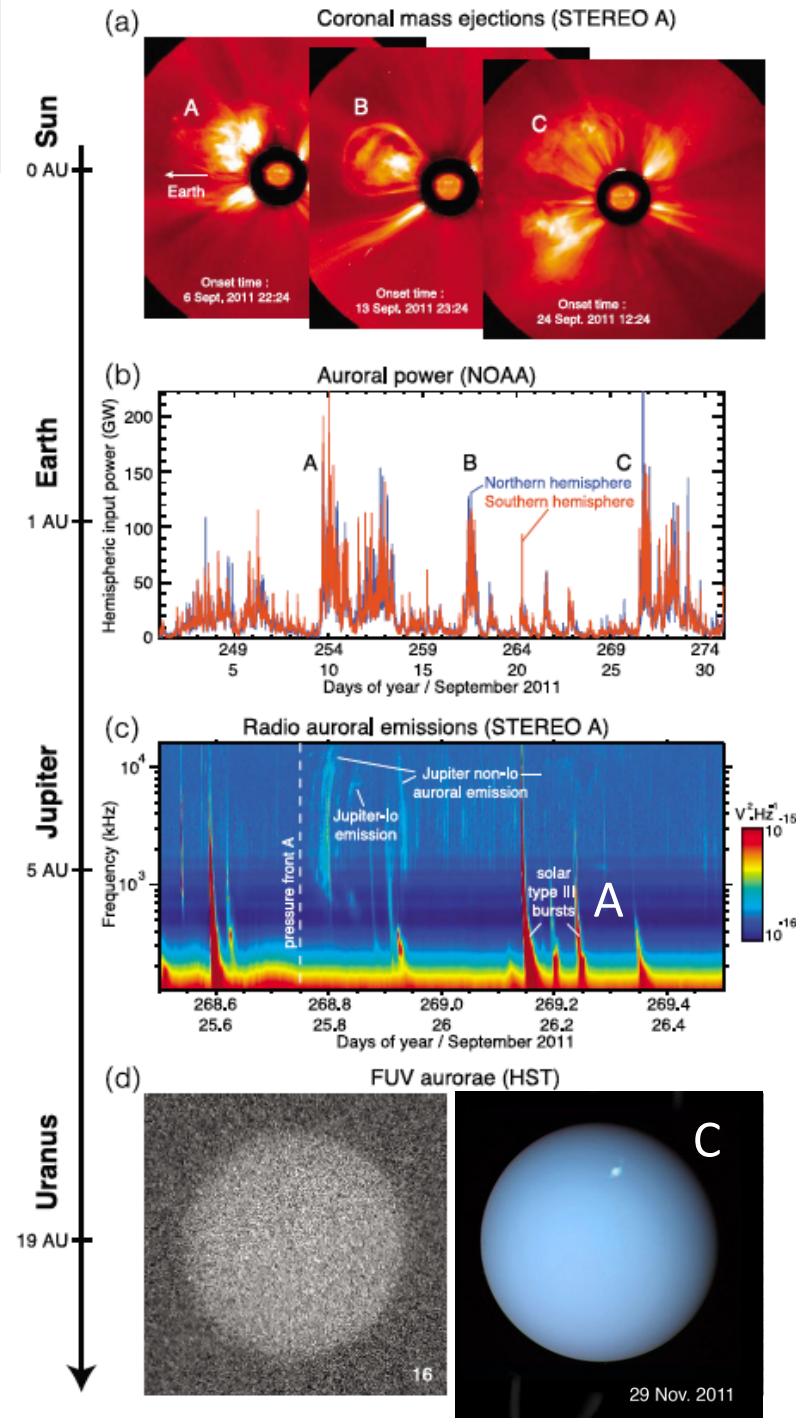
[Prangé et al., Nature, 2004]



# Interplanetary shock to Earth, Jupiter and Uranus



[Lamy et al., GRL, 2012]





# Thermosphere-ionosphere-magnetosphere coupling

# MAGNETOSPHERE-IONOSPHERE-THERMOSPHERE COUPLING

AURORAL  
THERMOSPHERE  
IONOSPHERE

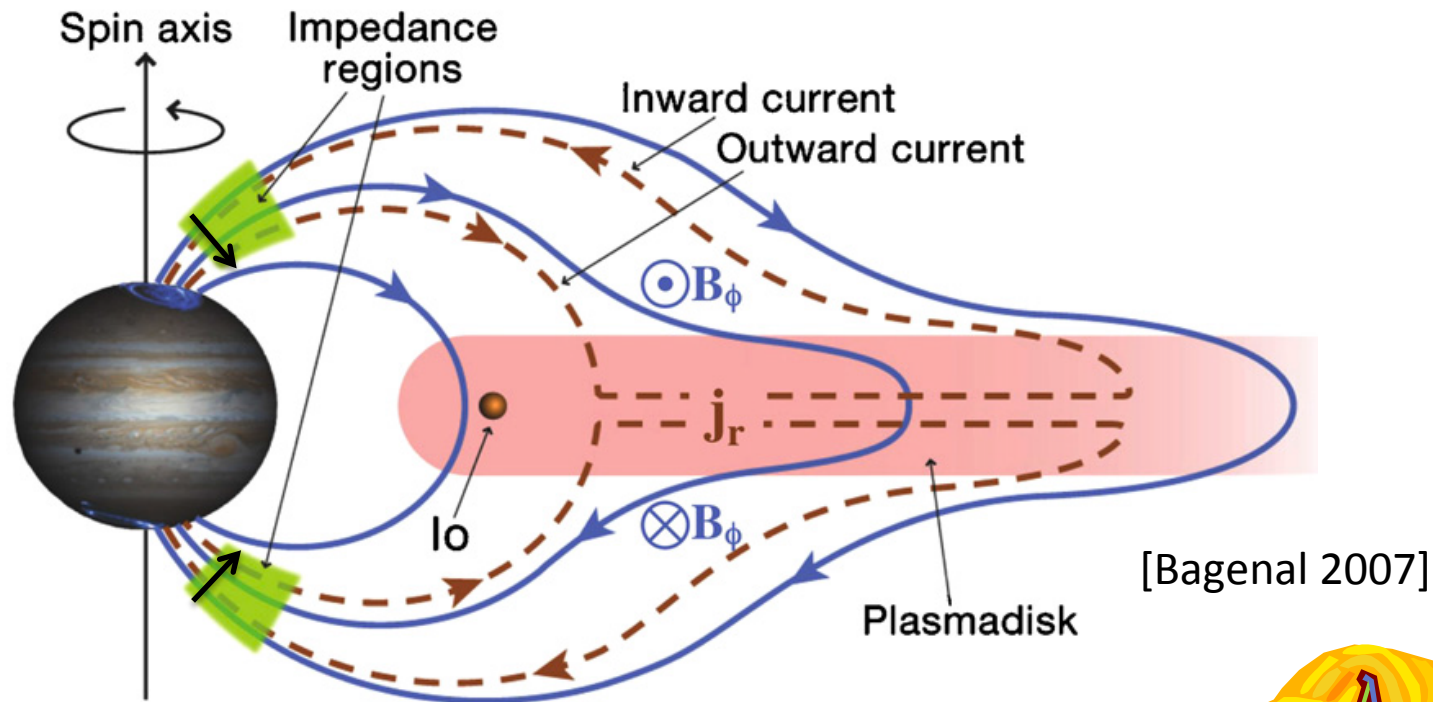
Exchange of  
particles, momen-  
tum & energy

MAGNETOSPHERE

Examples of ITM coupling:

- Angular momentum transfer
- Ion outflow, particle precipitation

# THERMOSPHERE-IONOSPHERE-MAGNETOSPHERE COUPLING



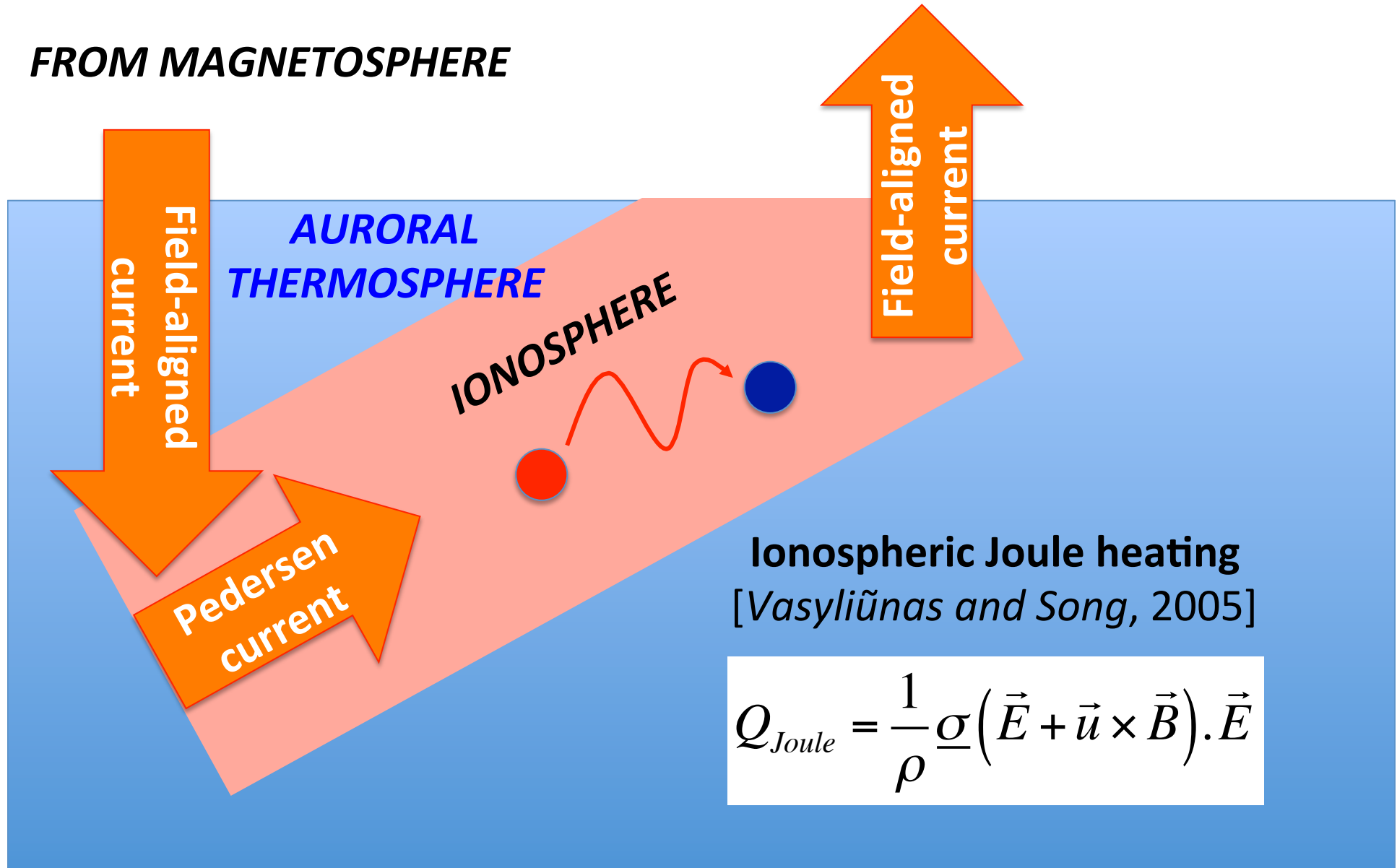
- Overall, at high latitudes:  
the magnetosphere extracts angular momentum from the upper atmosphere through the magnetic field-aligned currents



# THERMOSPHERE-IONOSPHERE-MAGNETOSPHERE COUPLING

**TO MAGNETOSPHERE**

**FROM MAGNETOSPHERE**



# Electrical, ionospheric conductivities

*Pedersen ( $\sigma_P$ )  
& Hall ( $\sigma_H$ )  
conductivities:*

with  $\omega = \frac{eB}{m}$

$$\sigma_P(z) = \sum_n \sum_i \frac{n_i e}{B} \left( \frac{\nu_{en\perp} \omega_e}{\nu_{en\perp}^2 + \omega_e^2} + \frac{\nu_{in} \omega_i}{\nu_{in}^2 + \omega_i^2} \right)$$

$$\sigma_H(z) = \sum_n \sum_i \frac{n_i e}{B} \left( \frac{\omega_e^2}{\nu_{en\perp}^2 + \omega_e^2} - \frac{\omega_i^2}{\nu_{in}^2 + \omega_i^2} \right)$$

*Pedersen ( $\Sigma_P$ ) & Hall ( $\Sigma_H$ )  
conductances:*

$$\Sigma_P = \int_{\text{ionosphere}} \sigma_P(z) dz$$



$$\Sigma_H = \int_{\text{ionosphere}} \sigma_H(z) dz$$

*Indices:* e (electrons), i (ions), n (neutrals)

*Variables:*

- n, number density
- $\omega$ , gyro-frequency (= e B / m)
- $\nu_{in}$ , ion-neutral collision frequency
- $\nu_{en\perp}$ , effective electron-neutral collision frequency perpendicular to  $\underline{B}$
- $\underline{B}$ , magnetic field

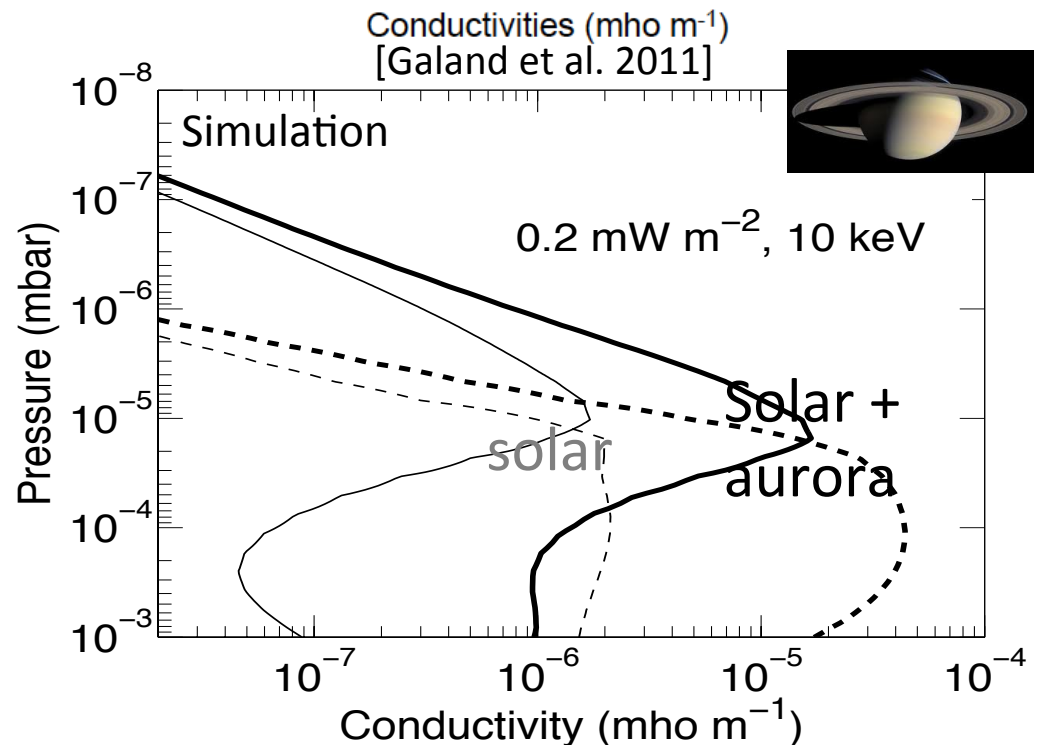
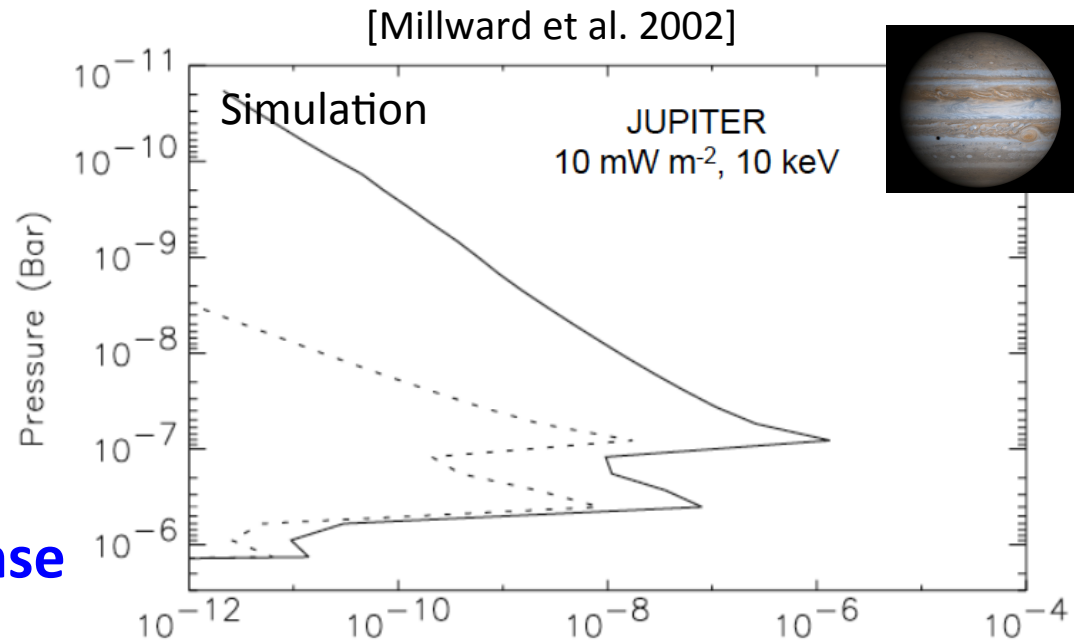
# Ionospheric conductivities in auroral regions

Pedersen   
 Hall 

At a given planet, when intense e- precipitation (within the reaction time of the ionosphere):  $\Sigma \propto Q_{prec}$

## Pedersen conductance

Jupiter	Saturn
60 keV, 10 mW m <sup>-2</sup>	10 keV, 1 mW m <sup>-2</sup>
2 mho	12 mho



# Ionospheric Conductances in auroral regions

Auroral electron mean energy & energy flux	<b>Earth</b> <i>[Fuller-Rowell and Evans, 1987]</i>	<b>Saturn</b> <i>[Galand et al., 2011]</i>	<b>Jupiter</b> <i>[Millward et al., 2002]</i>
10 keV 1 mW m <sup>-2</sup>	$\Sigma_P =$ <b>4-6 mho</b>	$\Sigma_P =$ <b>11-12 mho</b>	$\Sigma_P =$ <b>0.03 mho</b>

**Composition**  
**Altitude range**

$\Sigma_P = \max(\Sigma_P) / 10$  over 70 km (E)  
and 500 km (S)

**Strength of B field**

B field 20 times stronger  
at J cp w/ S

Slippage parameter<sup>(1)</sup> for Jupiter<sup>(2)</sup> & Saturn<sup>(3)</sup>:  $k = \frac{\Omega - \omega_n}{\Omega - \omega_i} = \sim 0.5$

Planetary angular velocity

$\Omega$

Angular velocity  
of the neutrals

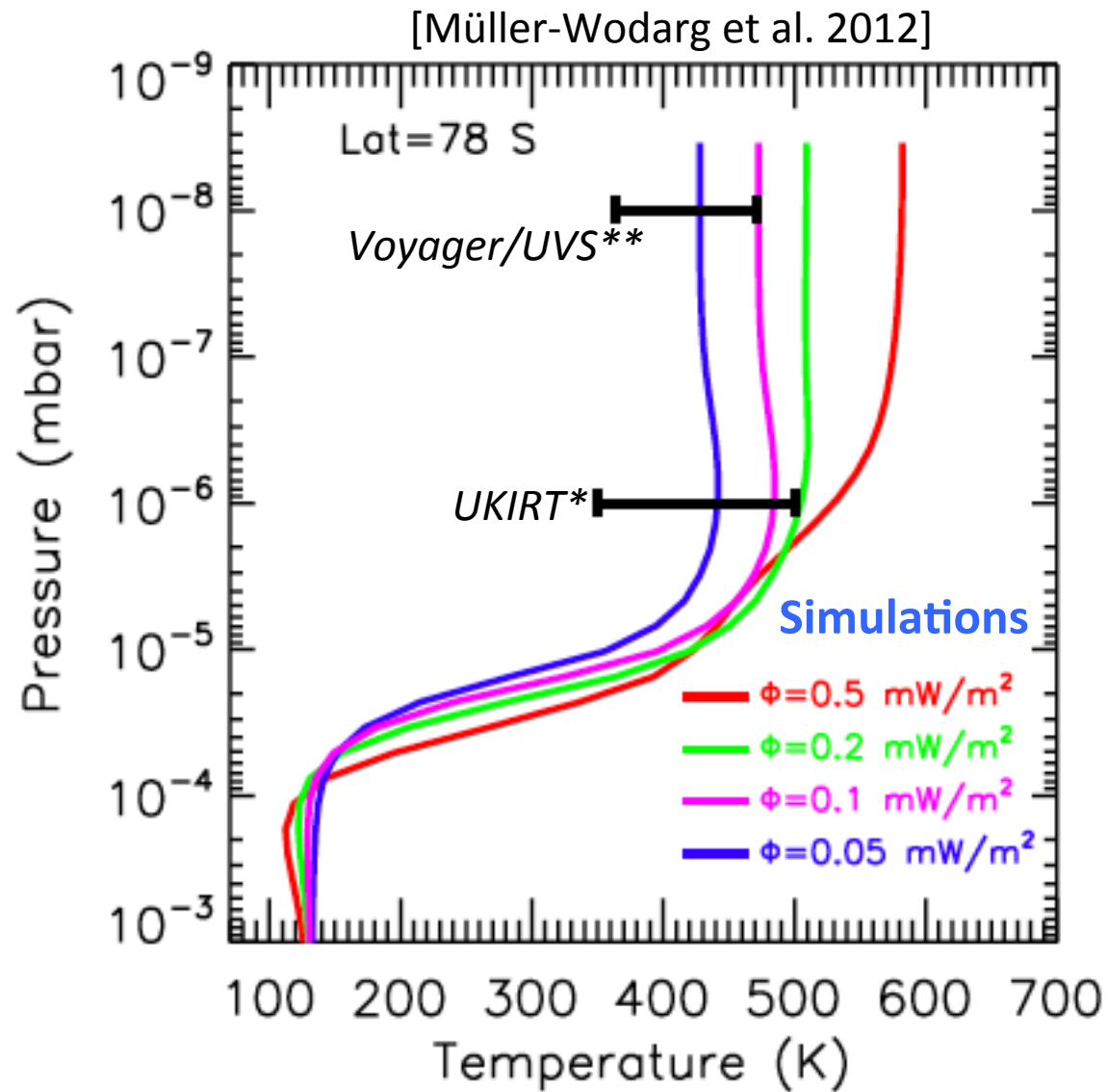
$\omega_n$

Angular velocity  
of the ions

$\omega_i$

(1) Huang and Hill [1989]; (2) Cowley et al. [2004]; (3) Galand et al. [2011]

# Ionospheric contribution to neutral temperature



\* Melin et al. (2007)

\*\* Vervack and Moses (2009)

**Auroral Joule heating sufficient to heat the high latitude ionosphere**



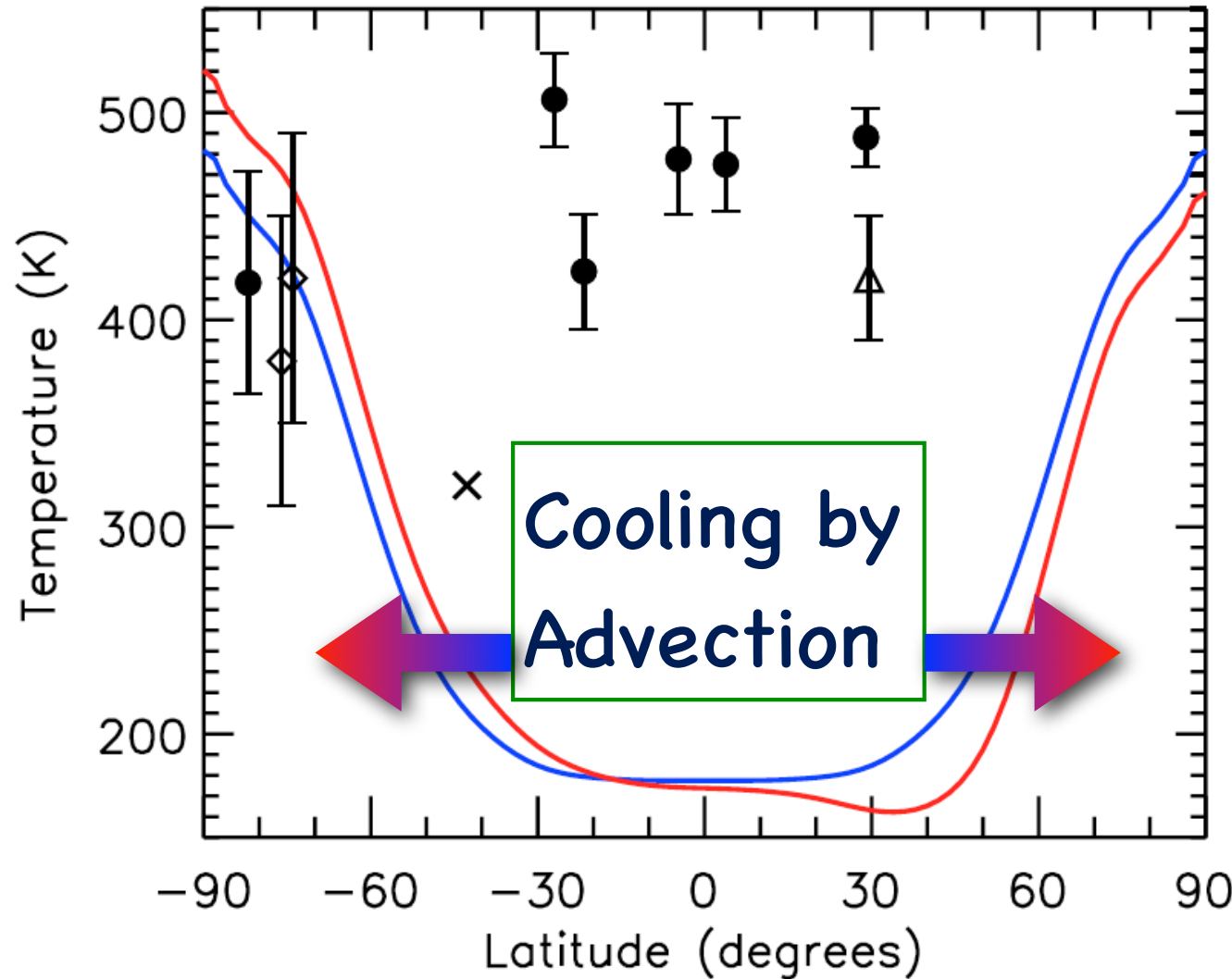
# Energy crisis at Saturn



Simulations:

STIM (Equinox)

STIM (Solstice)



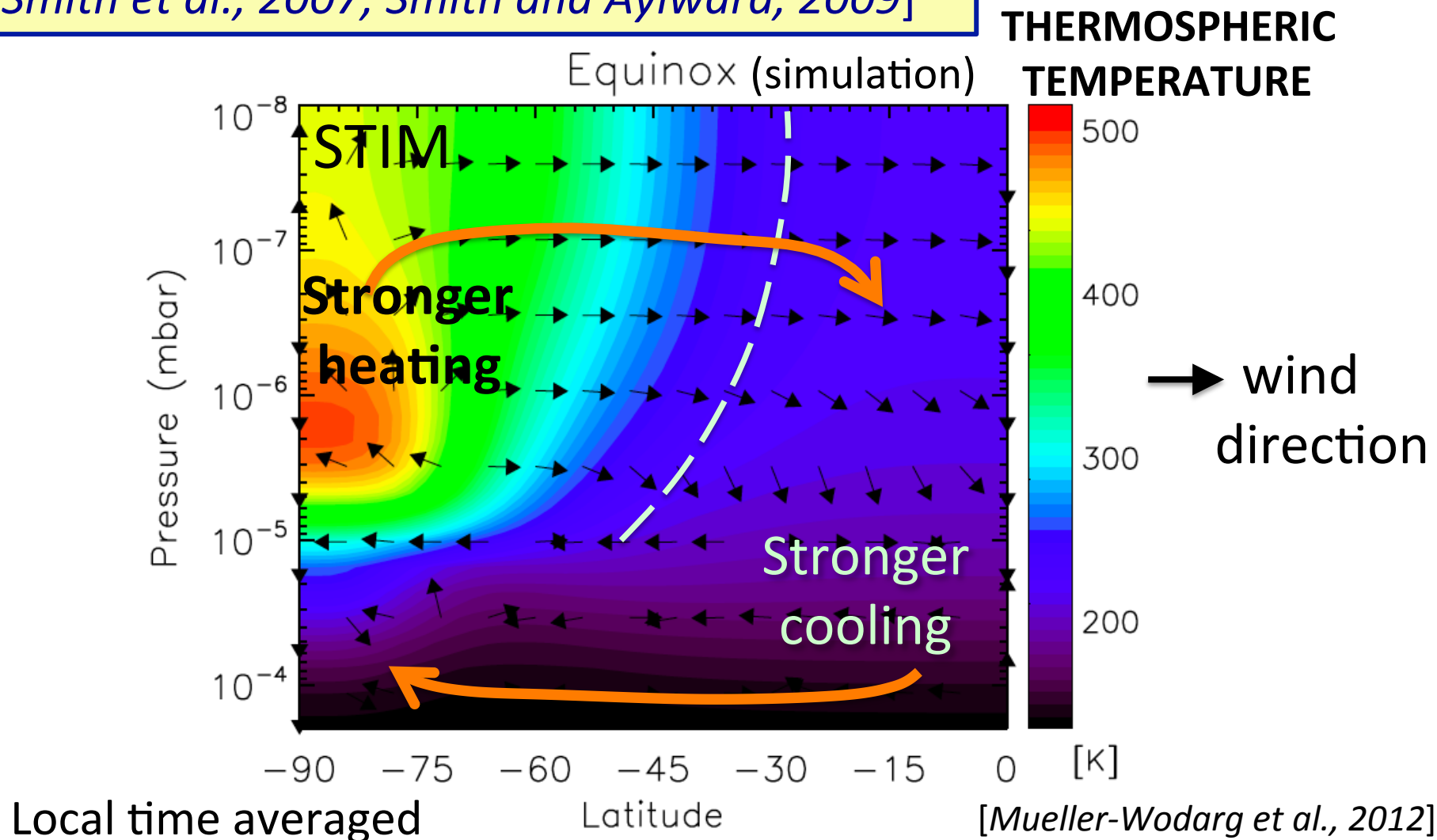
Points:

Measurements  
from Voyager  
UVS, Cassini  
UVIS and  
Ground Based  
IRTF



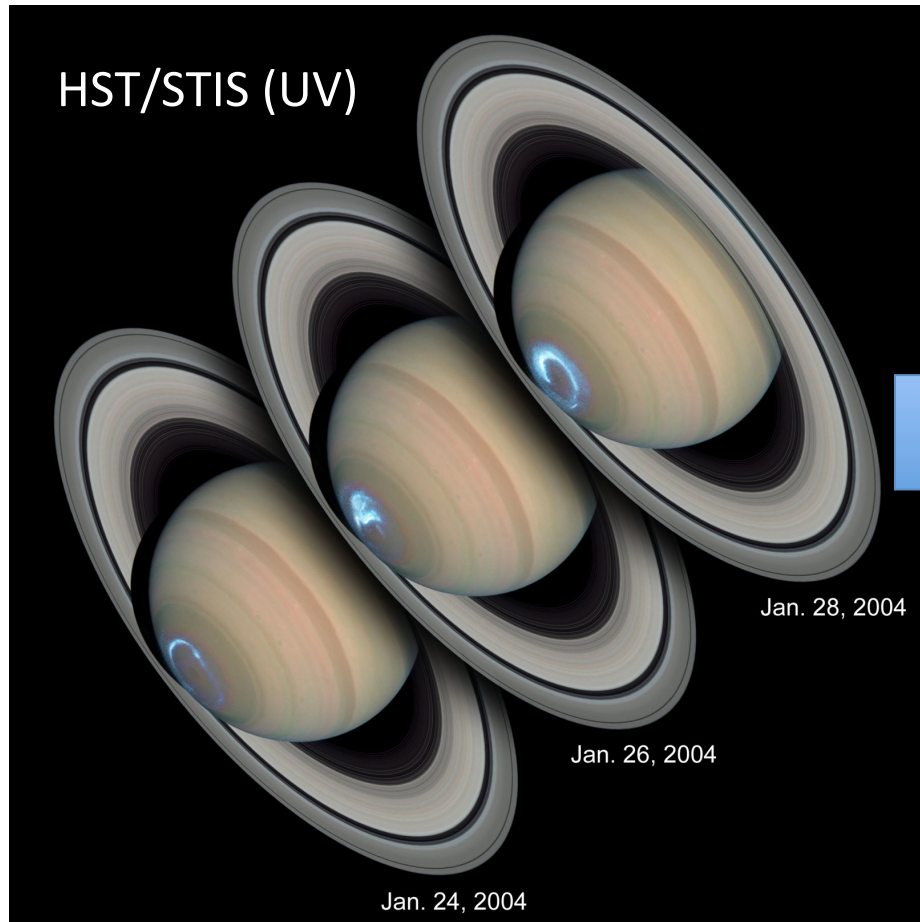
## Ion drag fridge mechanism

[*Smith et al., 2007; Smith and Aylward, 2009*]



**Polar sub-corotation due to auroral forcing (westward ion velocities due to ambient E fields) drives equator-to-pole circulation**

# Temporal change in auroral forcing



**“Magnetic storm”**  
**response** at Jupiter and Saturn remains largely **confined to auroral regions**

- no propagating gravity wave towards equator (as on Earth)

[Müller-Wodarg et al. 2013,  
Yates et al. 2013]

**Energy trapped in high latitude regions**

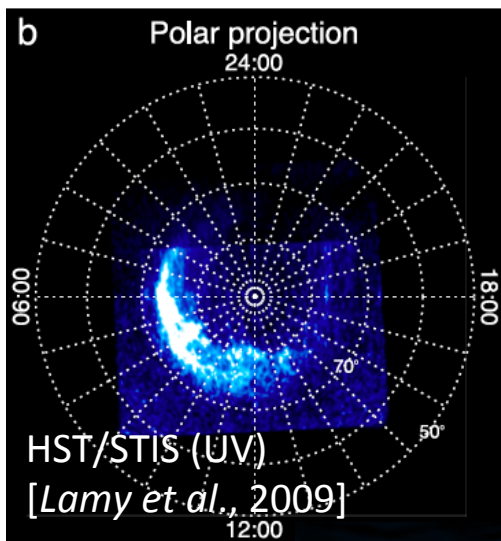
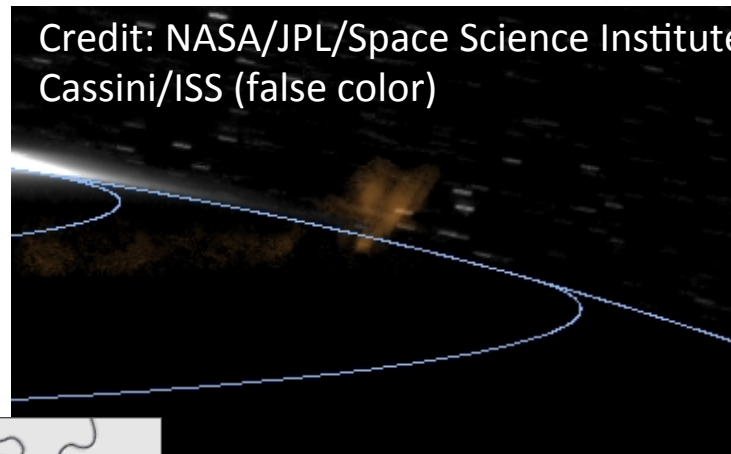
**→ Would shorter timescales be more efficient?**

Future prospect

**NASA/CASSINI  
(2004 → 2017)  
+ Earth-based support**

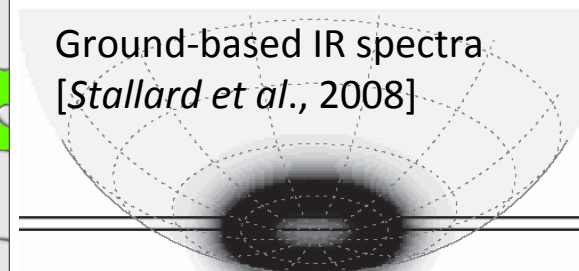
**THERMOSPHERE/  
IONOSPHERE  
COUPLED MODEL**

Credit: NASA/JPL/Space Science Institute  
Cassini/ISS (false color)



**Cassini/RSS radio  
occultations  
→ Ne**

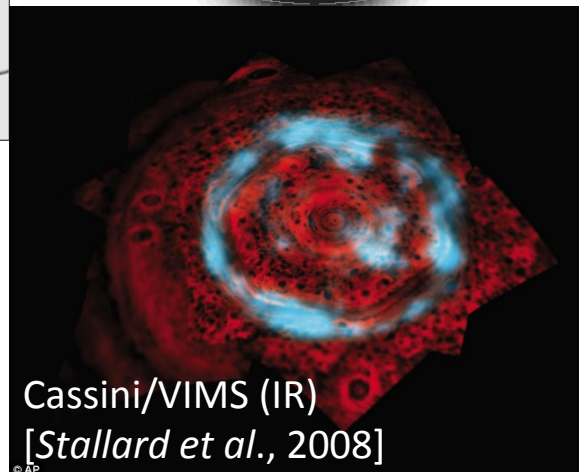
Ground-based IR spectra  
[Stallard et al., 2008]



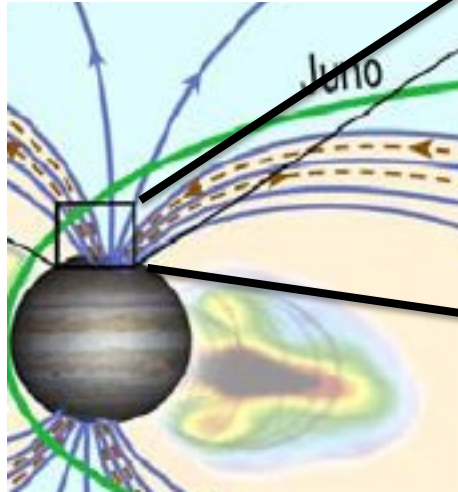
✓ **Combine as  
many as  
possible to  
better  
constrain the  
problem**



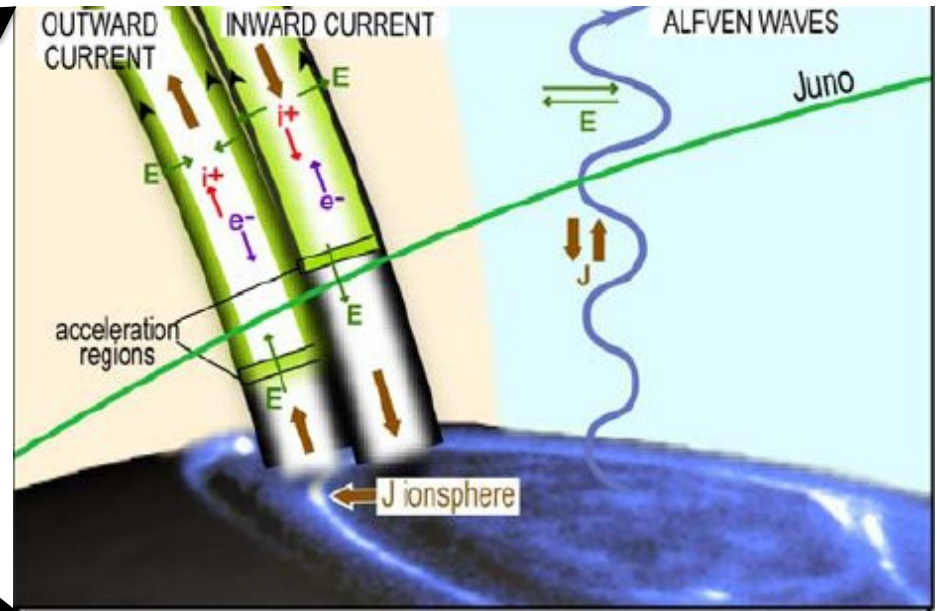
**Voyager/  
UVS  
Cassini/UVIS  
occultations  
→ Tn**



JUNO  
over the polar regions



Credit: Juno Team



Credit: Juno Team

July 2016!

**JUNO observations through the magnetic field lines connected to the auroral ionosphere, close/within the acceleration region (expected to be 2-3 RJ from center [e.g., Ray et al., 2009]):**

- Electric currents along magnetic field lines
- Plasma/radio waves revealing processing responsible for particle acceleration
- Energetic particles precipitating into atmosphere creating aurora
- Ultraviolet/IR auroral emissions regarding the morphology of the aurora

# TAKE HOME MESSAGE

- **ANALYSIS OF AURORAL EMISSIONS:**
  - **Valuable probe** of ionosphere (IR), auroral particle source, ITM coupling, and magnetic field line configuration
  - **Jupiter:** main oval driven by breaking down from co-rotation (Io)
  - **Saturn:** main oval mapped in the outer magnetosphere varying with solar wind conditions (Enceladus)
  - **Uranus:** solar wind dominated
- **T-I-M COUPLING:**
  - **Electrical, ionospheric conductances:**
    - Uncertainties in conductivities driven by limitation in electron density estimate
    - Differences in B field strength between Jupiter and Saturn yield differences in conductances. Larger energy fluxes at Jupiter do not seem to compensate for the strong B field. Implication on T-I-M coupling.
  - **Simulations:**
    - Critical to estimate the upper atmosphere response self-consistently
    - Play a key role in efforts to understand underlying physics
  - **Energy crisis** remains unsolved:
    - Investigate shorter timescale, E field variability unconstrained, role of waves, mid-lat e-?
  - Lessons learned from Saturn very useful for **upcoming exploration of Jupiter** (Juno / JUICE) and **exoplanets** (EChO)

# COMPARATIVE *CROSS BODY* APPROACH

✓ **Diversity** among solar system bodies in terms of:

▪ *physical, magnetic field, atmospheric, energy forcing settings*

→ makes *comparative aeronomy* an exciting and enriching field of research

✓ **Comparative aeronomy** challenges our understanding of atmospheric processes, coupling with neighboring regions, and planet evolution, as well as open new doors for extrapolating beyond our Solar System.

