

Which topic is (probably, at this point in time) your primary interest?

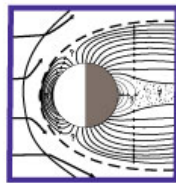
1. Solar physics
2. Heliosphere
3. Earth ionosphere/magnetosphere
4. Planetary space physics

Planetary Dynamos & Magnetospheres

See vol. III ch. 7 & vol. I ch. 13

MERCURY:

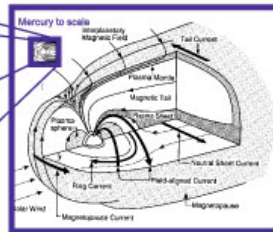
- Small
- Minute timescales
- Solar wind dominated



Mariner,
MESSENGER

EARTH:

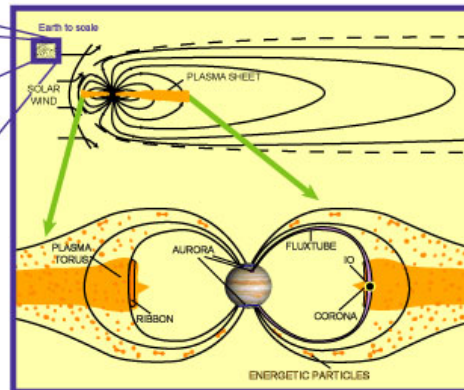
- Intermediate
- Hour timescales
- Solar wind driven



~100 missions since 1957
e.g. Polar, Geotail, FAST,
SAMPLEX, Cluster

JUPITER:

- Giant
- Timescales - minutes to months?
- Rotationally driven - solar wind triggered?



Pioneer, Voyager, Ulysses,
Galileo, Cassini

Testing our understanding of Sun-Earth connections through application to other planetary systems

Table 3 Planetary Magnetic Fields

	Ganymede	Mercury	Earth	Jupiter	Saturn	Uranus	Neptune
B_{Dipole}^a (nT)	719	170-270	30,600	430,000	21,400	22,800	14,200
Maximum / Minimum ^b	2	~2	2.8	4.5	4.6	12	9
Dipole Tilt and Sense ^c	-4°	~+10°	+9.92°	-9.4°	-0.0°	-59°	-47°
Dipole offset ^d (RP)	-	-	-	0.119	0.038	0.352	0.485
Obliquity ^e	0°	0°	23.5°	3.1°	26.7°	97.9°	29.6°
Range in Solar Wind Angle ^f	90°	90°	67-114°	87-93°	64-117°	8-172°	60-120°

^a Surface field at dipole equator. Values derived from modeling the magnetic field as an offset, tilted dipole (OTD).

^b Ratio of maximum surface field to minimum (equal to 2 for a centered dipole field). This ratio increases with larger non-dipolar components and tends to increase with the planet's oblateness.

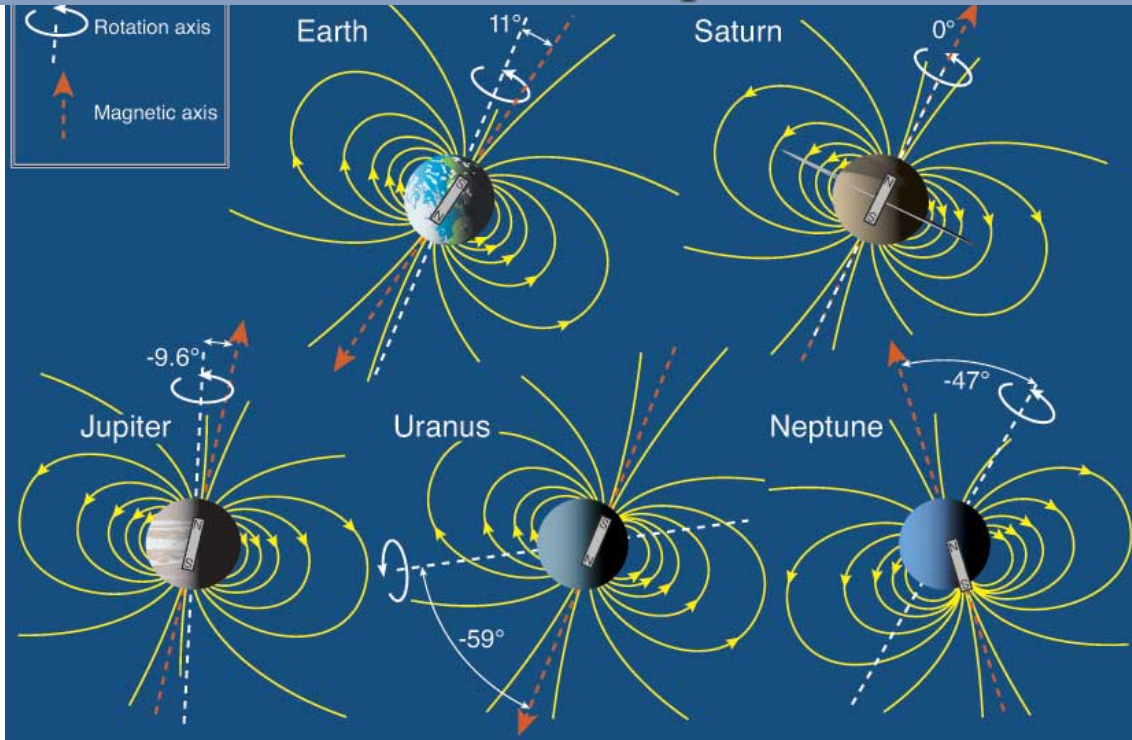
^c Angle between the magnetic and rotation axes. Positive values correspond to magnetic field directed north at the equator.

^d Values for the giant planets come from dipole (OTD) models of Connerney (1993). The Earth's dipole is from the International Geomagnetic Reference Field while the magnetic dip poles of the Earth's field are located (in 2010) at 85°N and 64°S latitudes and moving over 10 degrees per century (Finlay et al. 2010).

^e The inclination of a planet's spin equator to the ecliptic plane.

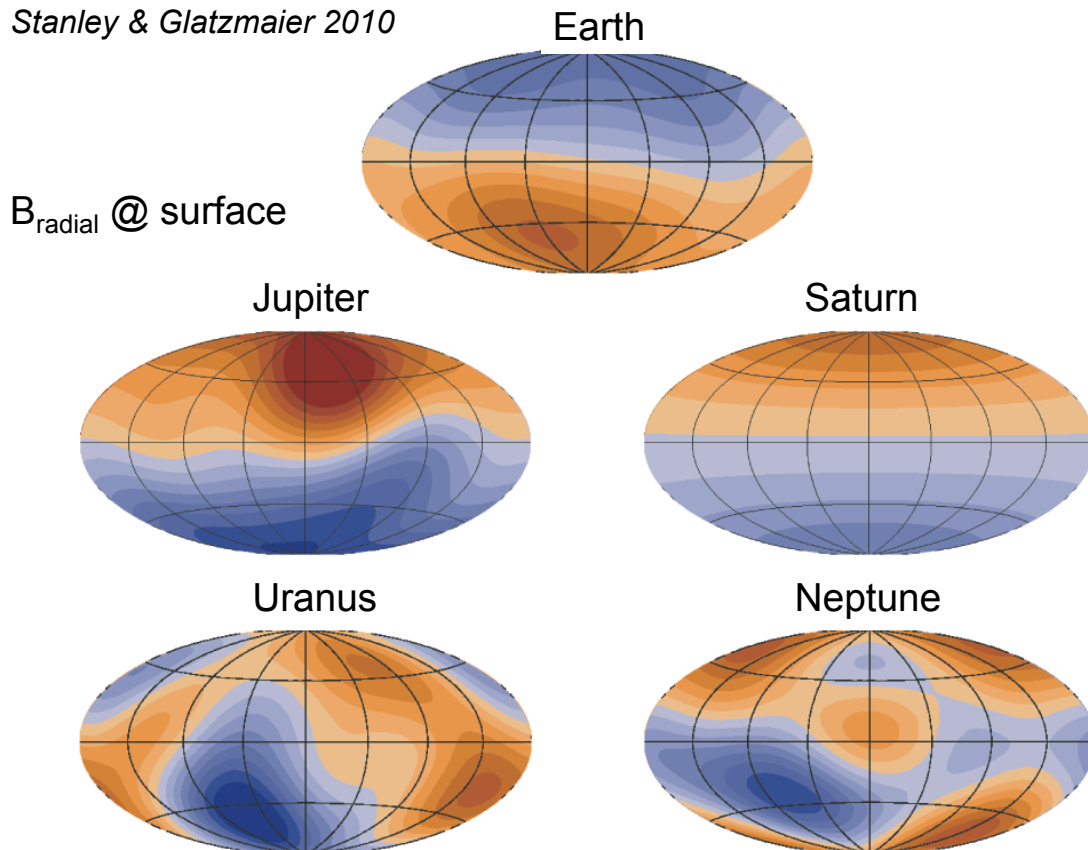
^f Range of angle between the radial direction from the sun and the planet's rotation axis over an orbital period. In Ganymede's case, the angle is between the corotational flow and the moon's spin axis.

Tilts and Obliquities



Offset Tilted Dipole (poor) Approximation

Stanley & Glatzmaier 2010



Magnetic Potential 3-D harmonics

$$\mathbf{B} = -\text{grad } V$$

$$V = R_p \sum_{n=1}^{\infty} \sum_{m=0}^n \left(\frac{R_p}{r}\right)^{n+1} P_n^m(\cos \theta) (g_n^m \cos m\lambda + h_n^m \sin m\lambda), \quad (7.1)$$

coefficients - constants

functions

$$P_0^0(\cos \theta) = 1$$

$$P_1^0(\cos \theta) = \cos \theta$$

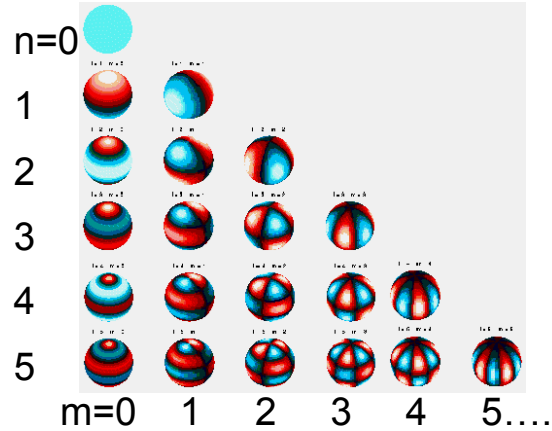
$$P_1^1(\cos \theta) = -\sin \theta$$

$$P_2^0(\cos \theta) = \frac{1}{2}(3 \cos^2 \theta - 1)$$

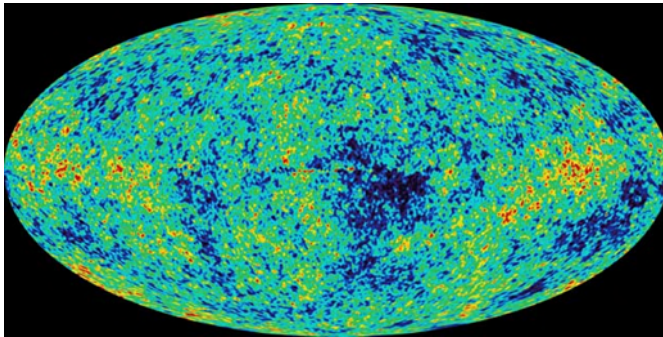
$$P_2^1(\cos \theta) = -3 \cos \theta \sin \theta$$

$$P_2^2(\cos \theta) = 3 \sin^2 \theta$$

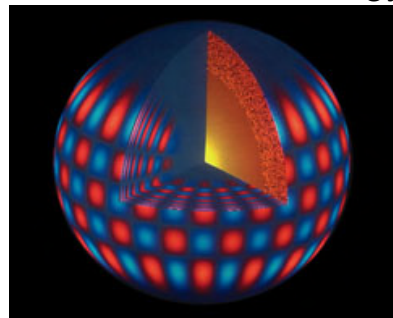
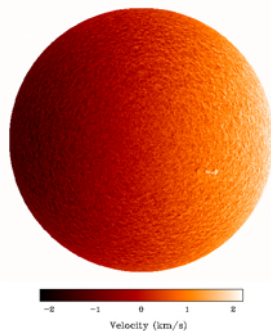
$$P_3^0(\cos \theta) = \frac{1}{2}(5 \cos^3 \theta - 3 \cos \theta)$$



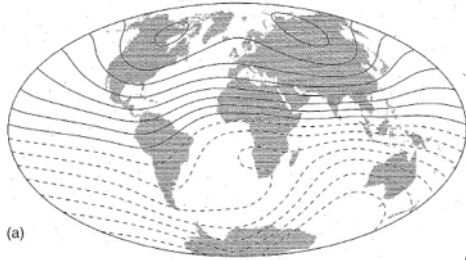
Same technique used to model cosmic microwave background



or interior of Sun with Helioseismology...

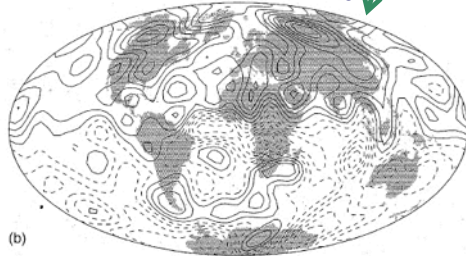


Br surface



(a)

Br core-mantle boundary



(b)

Fig. 7.2. Radial component of the geomagnetic field at the Earth's surface (a) and at the core-mantle boundary (b). Full lines for inward magnetic flux and dashed lines for outward flux. Contour intervals are arbitrary and different in the two panels.

From accurate measurement of surface field:

1. Extrapolate to core-mantle boundary = dynamo

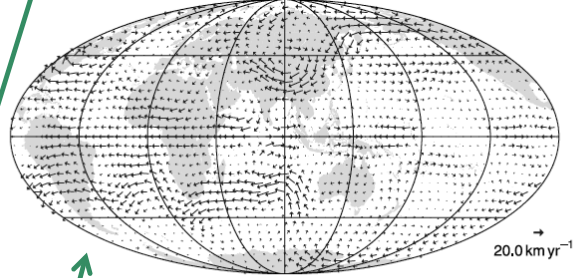


Fig. 20 Large-scale core surface flow constructed by Holme and Olsen (2006) under the frozen-flux and tangential geostrophy assumptions from a satellite observation-derived secular variation model

Hulot et al. 2010

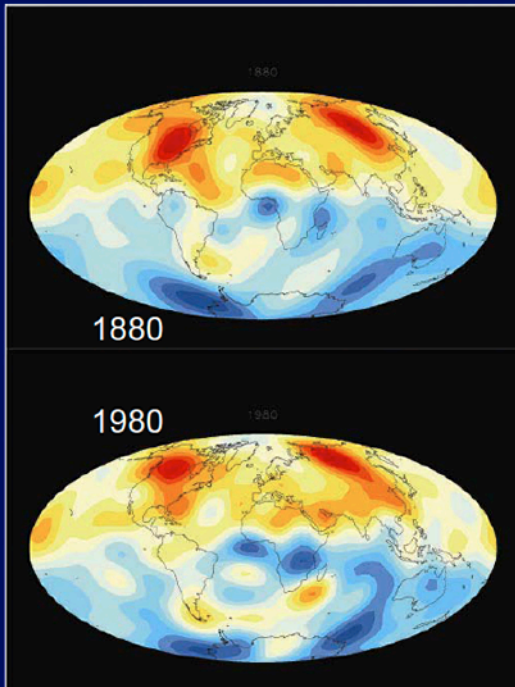
2. Derive core flows

$$\frac{\partial B_r}{\partial t} = -\nabla_h \cdot (\mathbf{u}_h B_r)$$

h=horizontal

3. Secular variation & reversals.....

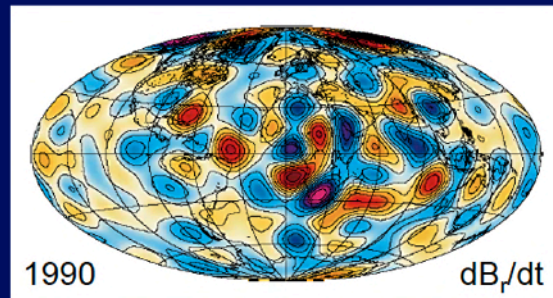
Secular variation



1880

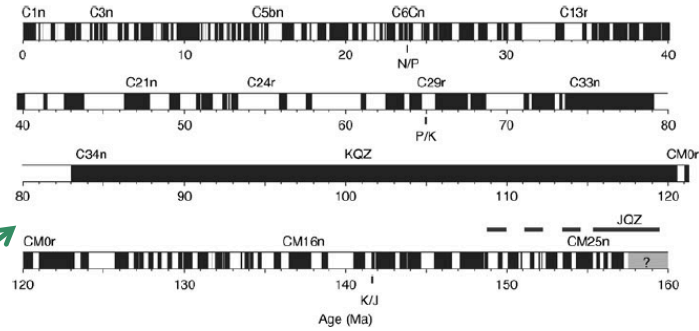
1980

Dipole dropped by 9% since 1840
 Reconstructions of core field morphology 1590 - now
 Fluctuations of non-dipole parts on time scales 50 – 400 yrs
 Stability of high-latitude flux lobes
 Westward drift in Atlantic / Africa



1990

dB_r/dt



Polarity reversals:

1. variable in duration and
2. rate

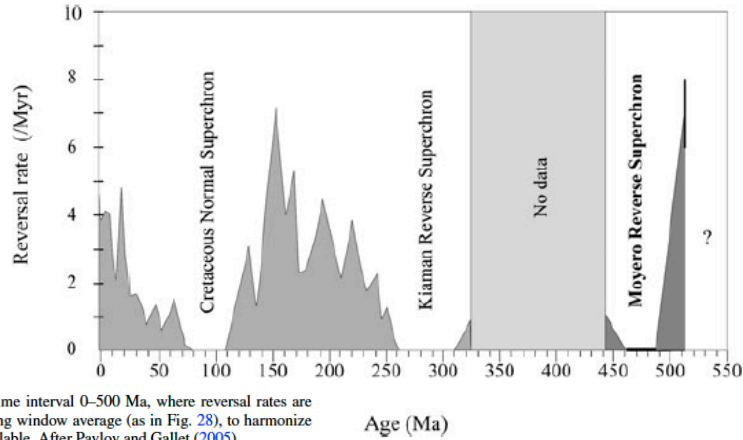
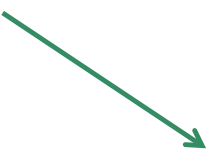
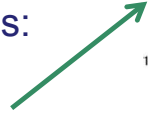
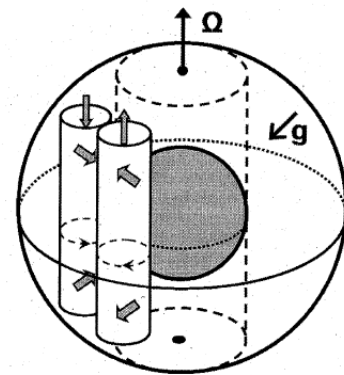


Fig. 29 Estimate of the reversal rate (in Myr^{-1}) for the time interval 0–500 Ma, where reversal rates are estimated by geological stage, rather than by using a moving window average (as in Fig. 28), to harmonize post- and pre- 150 Ma data for which much less data is available. After Pavlov and Gallet (2005)

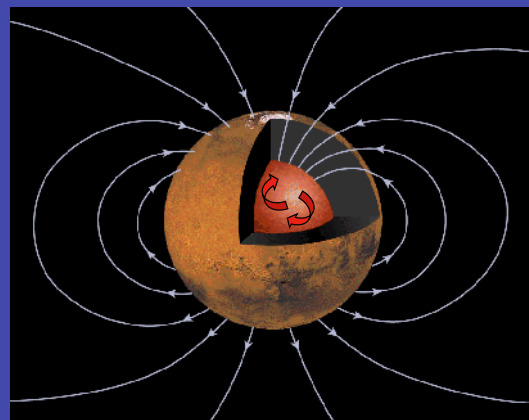
Planetary Dynamos

See vol. III ch. 7

**Volume of electrically
conducting fluid ①
which is convecting ②
and rotating**



All planetary objects probably have enough rotation - the presence (or not) of a global magnetic field tells us about ① and ②



Magnetic fields of solar system planets

Spacecraft detected magnetic fields at most (but not all) major planets

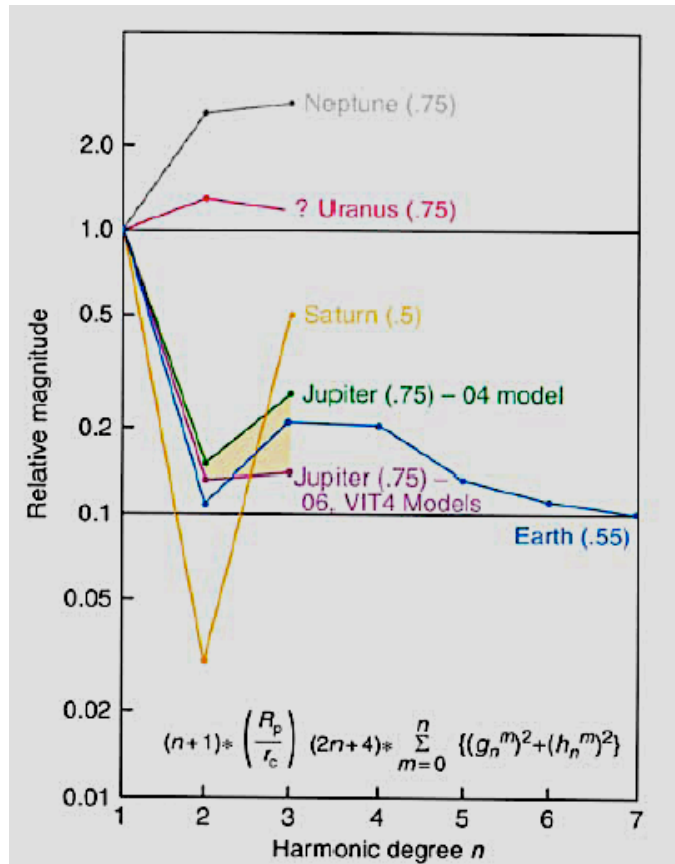
Planet	Dynamo	R_c/R_p	B_s [μ T]	Dip. tilt	Quadr / Dipole
Mercury	Yes (?)	0.75	0.35	<5°?	0.1-0.5
Earth	Yes	0.55	44	10.4°	0.04
Jupiter	Yes	0.84	640	9.4°	0.10
Saturn	Yes	0.6 ?	31	0°	0.02
Uranus	Yes	0.75	48	59°	1.3
Neptune	Yes	0.75	47	45°	2.7
Ganymede	Yes	0.3 ?	1.0	< 5° ?	?

V. dipolar

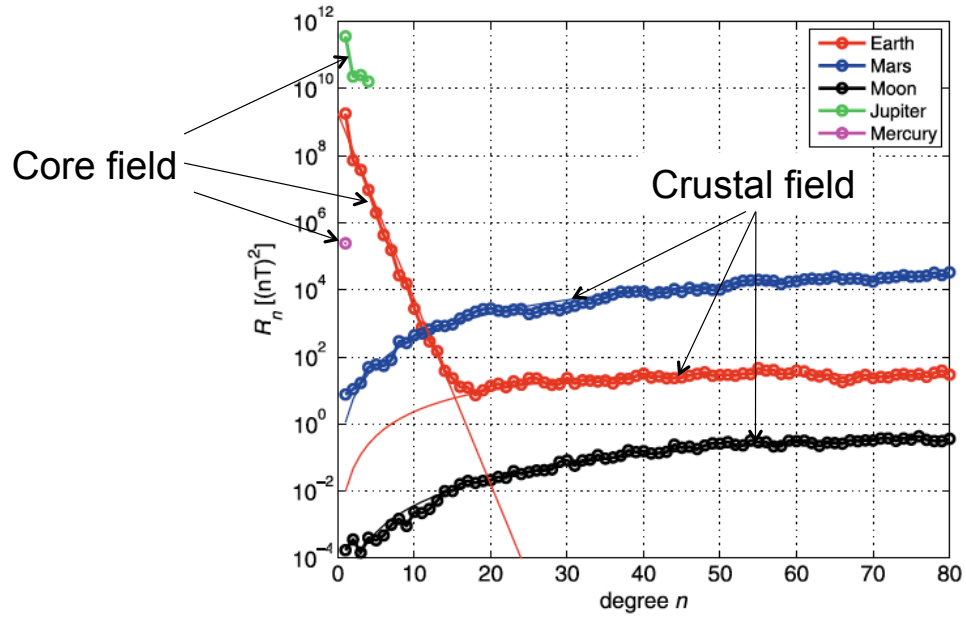
Irregular

R_c / R_p : core / planetary radius, B_s : Mean field at planet's surface, Quadr. / dipole power at R_c

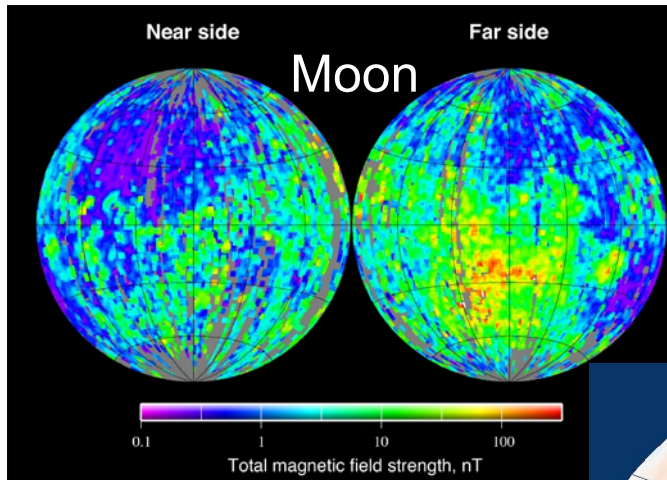
- Dipole – $n=1$
- Quadupole – $n=2$
- Quadrupole/Dipole ratio indicates irregular field



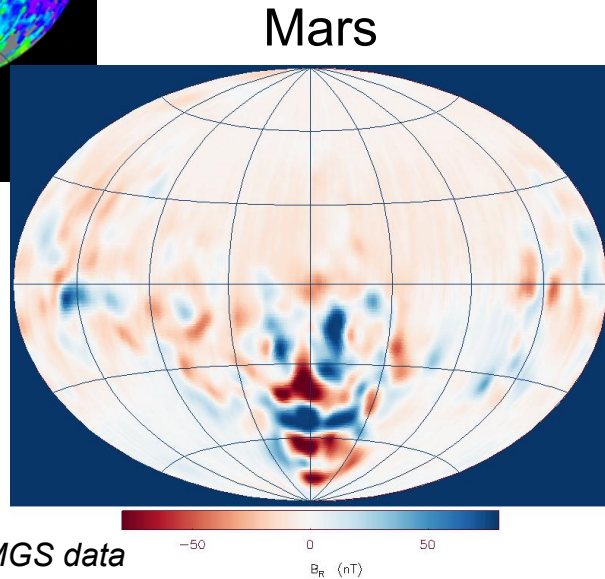
Moon & Mars: All Crustal Remanent Magnetization



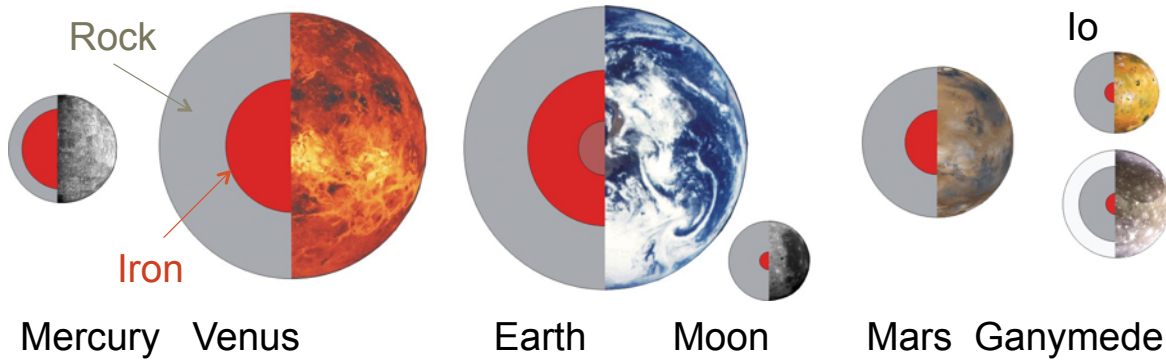
Power spectra of the field of internal origin for the Earth (after Olsen et al. 2009a and Maus et al. 2008), Mars (after Cain et al. 2003), Jupiter, Mercury (after Connerney 2008) and the Moon (after Purucker 2008) at their respective surface reference radius. Also shown are theoretical crustal spectra (thin curves, Voorhies et al. 2002) for the Earth, Mars and the Moon.



From Wikipedia – Lunar Prospector observations



From David Brain, MGS data



Mercury Venus Earth Moon Mars Ganymede

Planet	Density [$g\ cm^{-3}$]	R_{core}/R_{planet}
Mercury	5.43	0.75
Venus	5.24	0.55
Earth	5.515	0.55
Moon	3.36	0.2
Mars	3.94	0.5
Ganymede	1.94	0.3

*Why don't
Venus or Mars
have dynamos?*

What drives
the
geodynamo

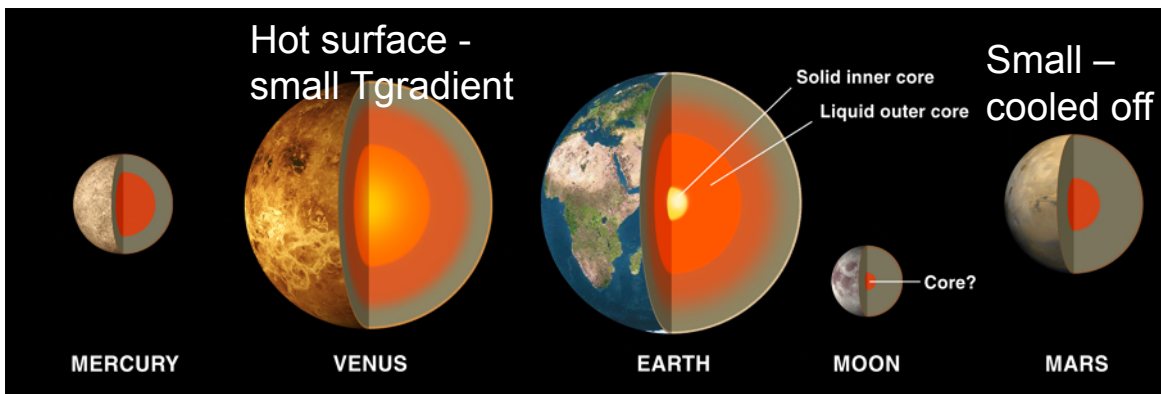
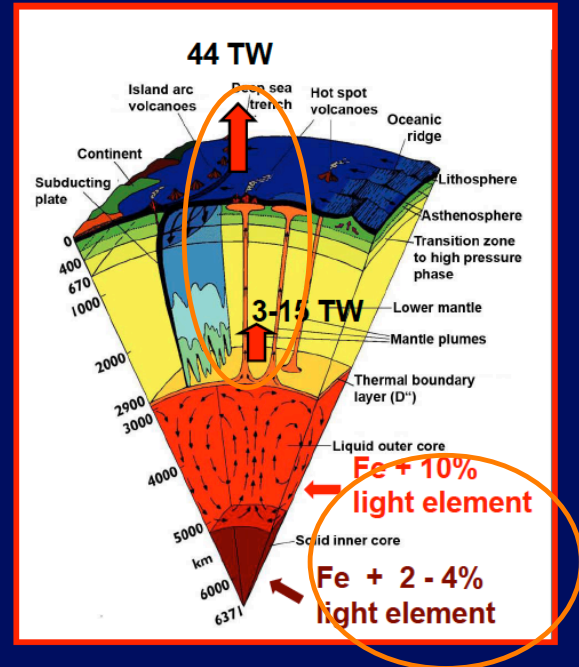
In words...
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Or by picture...

gravitational acceleration). In terrestrial planets, the adiabatic heat flow can be a large fraction of the actual heat flow, or it may exceed the actual heat flow, in which case at least the top layers of the core would be thermally stable. Near the top of Earth's core approximately 3–4 TW can be conducted along the adiabat (Lay *et al.*, 2008), i.e. close to the minimum estimates for the entire core heat flow. But even if all heat flux near the core–mantle boundary were carried by conduction, a convective dynamo can exist thanks to the inner core. At the inner core boundary, the adiabatic temperature profile of the convecting outer core crosses the melting point of iron. The latter increases with pressure more steeply than the adiabatic gradient, which is the reason why the Earth's core freezes from the center rather than from above. As the core cools, the inner core grows with time by freezing iron onto its outer boundary. This has two important implications for driving the dynamo. The latent heat that is released upon solidification is an effective heat source, which contributes to the heat budget approximately the same amount as the bulk cooling of the core. The heat flux that originates at the inner core decreases with radius as r^{-2} in the spherical geometry of the fluid core. The adiabatic temperature gradient is roughly proportional to r , because gravity decreases towards the center. Therefore, even if the actual heat flux were slightly less than the adiabatic heat flux near the core–mantle boundary, it must be superadiabatic deeper down. A second, perhaps more important effect is that the light elements in the outer core are preferentially rejected when iron freezes onto the inner core. Hence, they become concentrated in the residual fluid near the inner core boundary. This layering is gravitationally unstable because of the reduced density, which leads to compositional convection that homogenizes the light elements in the bulk of the fluid core. Compositional convection contributes as much as, or more than, thermal convection to the driving of the geodynamo in recent geological times.

Earth: Internal structure & energetics



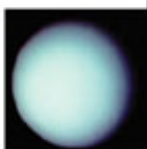

- Seismology: Dense core with $R_c/R_p=0.55$
- Fe only cosmochemically abundant element matching density
- No shear waves in outer core, hence it is liquid
- Solid inner core with $0.35R_c$
- ~10% light element (Si, S, O, ...) in outer core, less in inner core
- Earth heat flow 44 TW. Core fraction estimated 3-15 TW
- Core heat flow mostly due to secular cooling

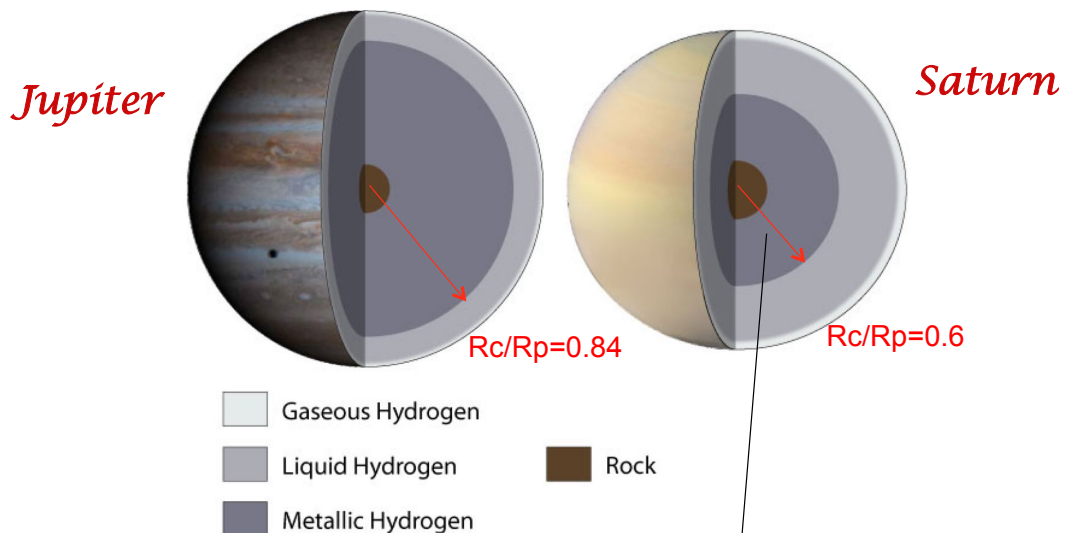


Why Don't Venus or Mars have Dynamos?

- Enough rotation – even for Venus
- Conducting fluid core – probably
- Lack of convection in core?
 - Mantle convection controls heat flow from core
 - Lack of plate tectonics suggests less efficient cooling of interior and lower heat flux from core
 - No inner core means no latent heat of solidification and no enhancement of lighter material in the outer core

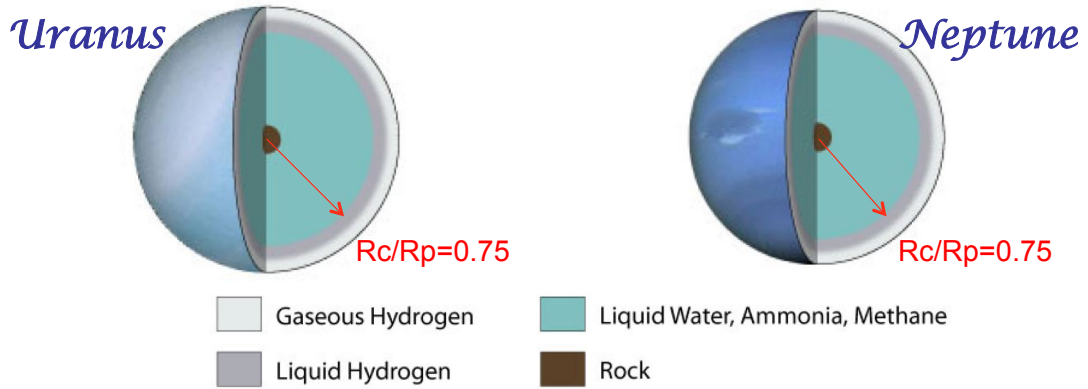
Giant Planets

		Distance AU	Mass Earth Mass	Radius Earth Radii	Density 1=water	Composition % by mass
	<i>Jupiter</i>	5.20	318	11.2	1.33	90% H, He
	<i>Saturn</i>	9.54	95	9.46	0.71	75% H, He
	<i>Uranus</i>	19.2	14	3.98	1.24	10% H, He Water Ammonia Methane
	<i>Neptune</i>	30.1	17	3.81	1.67	10% H, He Water Ammonia Methane



Saturn has lower mass

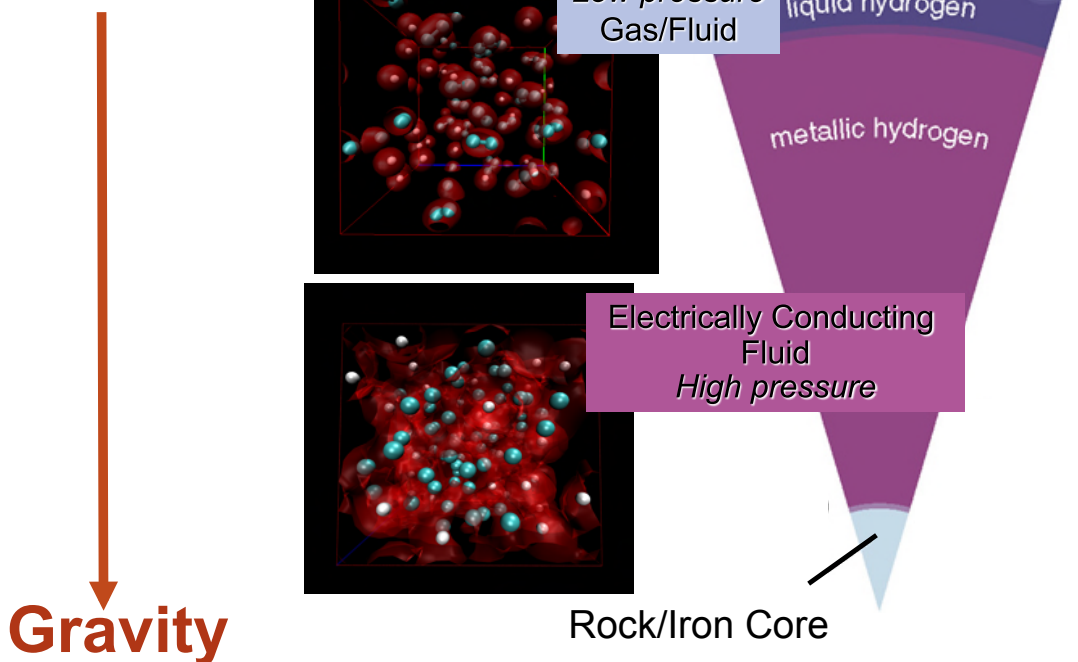
- lower pressures
- smaller region of metallic hydrogen
- weaker magnetic field



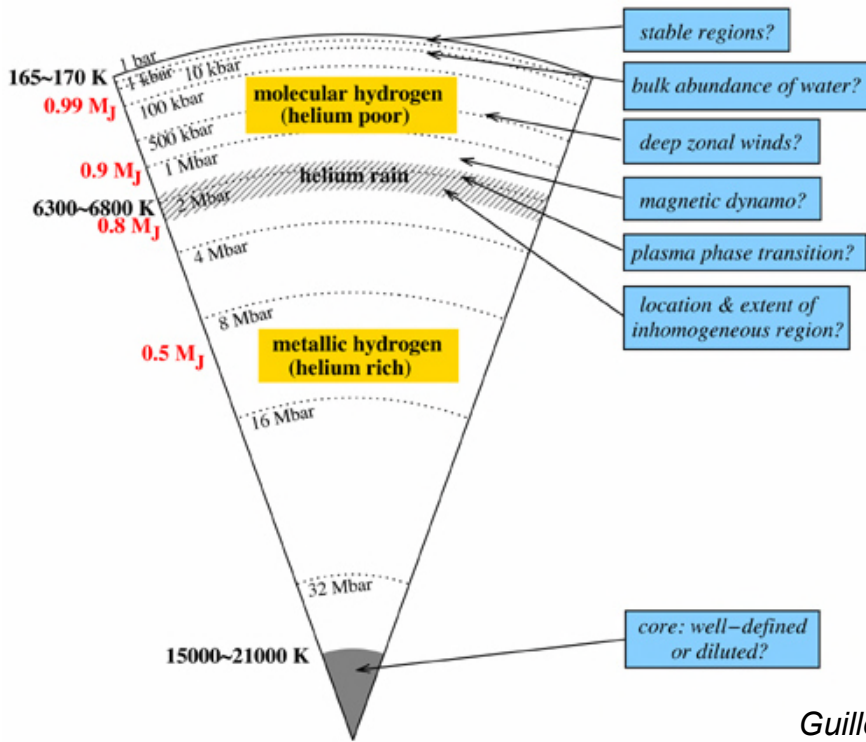
Uranus and Neptune have much less mass

- Lower pressures
- No metallic hydrogen
- Weak & irregular magnetic fields produced in water layer, deep below gas envelope

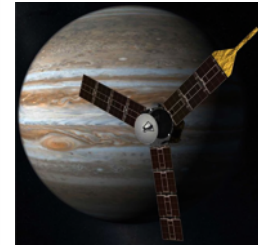
**Mass of Jupiter
~320 x Earth**



Interior of Jupiter Using Best Equation of State



Still lots of unknowns!!



Juno
Launch Aug. 5th

Guillot et al. 2004

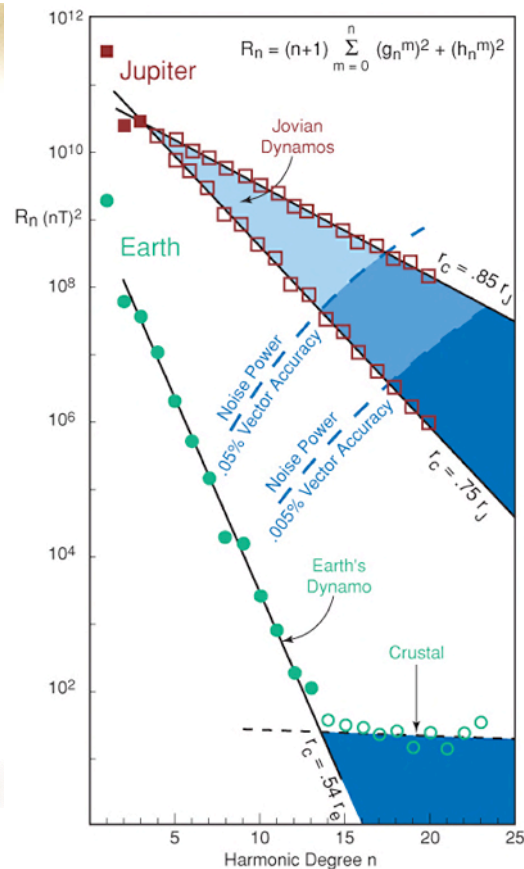
Juno Magnetic Spectra of Earth and Jupiter

Current knowledge of Jupiter is limited to $n < 4$

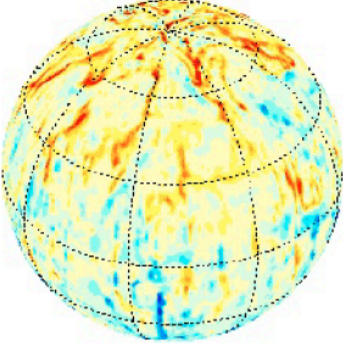
Earth dynamo at $n > 14$ is hidden by crustal field

Juno will measure out to $n \sim 20$

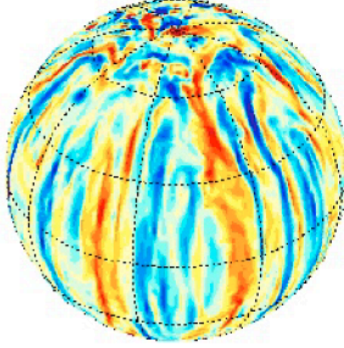
Determine spectral shape, dynamo radius, and secular variations



Dynamo Models – Lecture 1 Eqns. & discussion



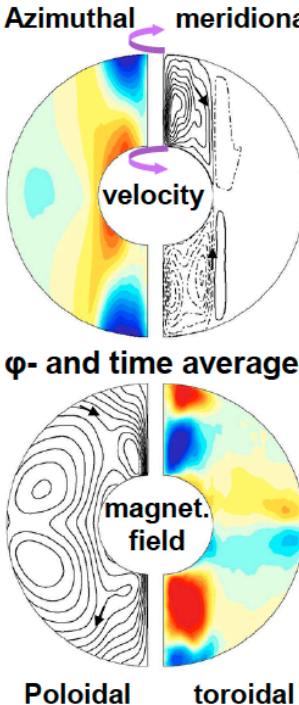
Radial magnetic field



Radial velocity

$E=10^{-5}$ $Ra/Ra_c=114$ $Pm=0.8$ $Rm=914$

- Flow columnar outside tangent cylinder
- Vigorous flow inside tangent cylinder; polar plumes
- Strong toroidal field inside tangent cylinder
- $\alpha^2\Omega$ – dynamo ? (Ω -effect inside tangent cylinder)



Azimuthal meridional

velocity

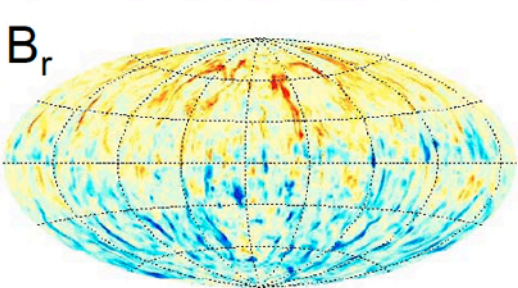
ϕ - and time average

magnet. field

Poloidal toroidal

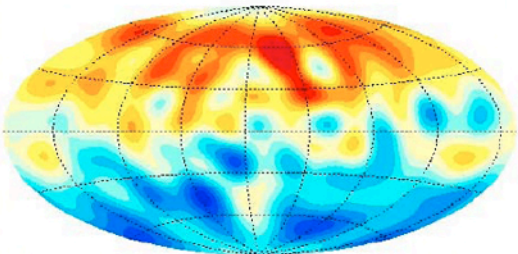
Heliophysics Summer School July 2009
Christensen: Planetary magnetic fields and dynamos
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Comparison with Earth: Field morphology



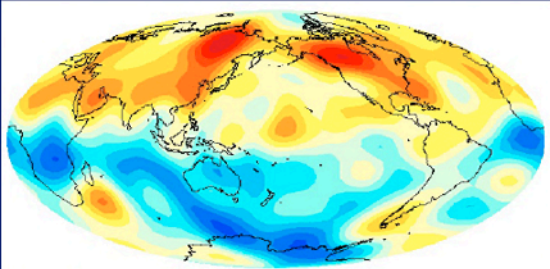
B_r

Dynamo model, full resolution



Dynamo model, filtered to $n < 13$

- Flux lobes at 60-70° latitude
- Weak flux at poles
- Flux spots of both polarities at low latitude. Expulsion of toroidal field bundles ?
- Westward vortex flow in polar cap



Earth's field at core mantle boundary

Heliophysics Summer School July 2009
Christensen: Planetary magnetic fields and dynamos
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Power-controlled field strength

Hypothesis: The magnetic energy density depends on thermodynamically available energy flux, that is the part of the energy flux that can be converted to magnetic energy and can balance ohmic dissipation
 The field strength is independent of rotation rate, conductivity, viscosity,...

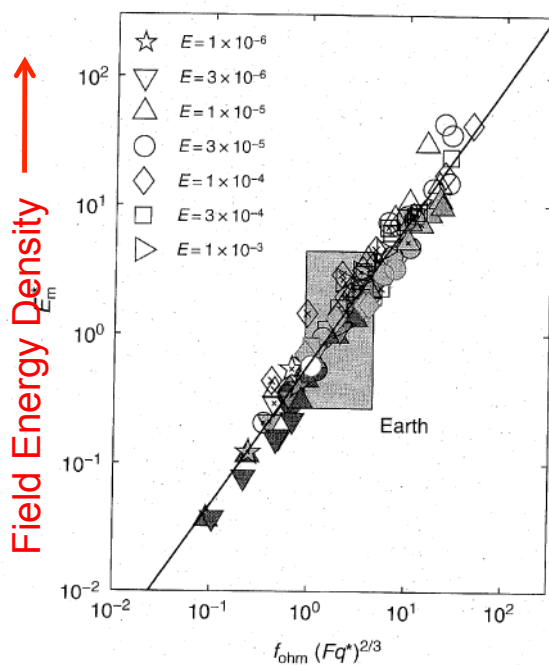
$$B^2/2\mu_0 \sim f_{\text{ohm}} \rho^{1/3} (L/H_T q_c)^{2/3}$$

q_c : convected heat flux, $H_T = c_p/(ag)$: temp. scale height, L : charact. radial length scale,
 ρ : density, f_{ohm} : ratio ohmic dissipation / total dissipation

Dynamo Scaling Laws

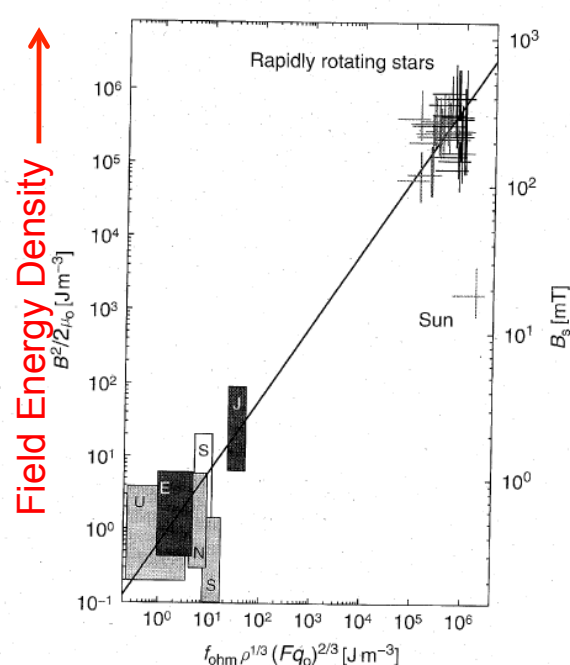
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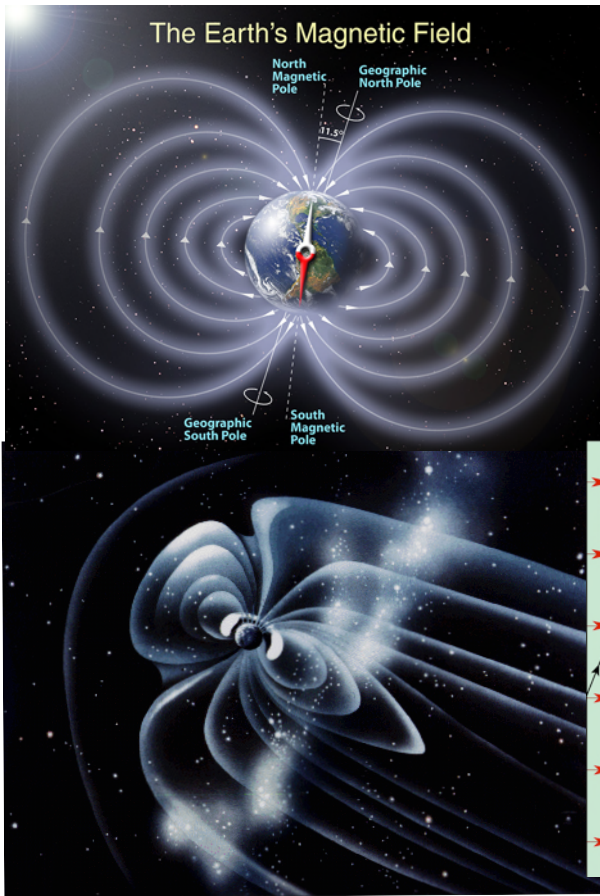
Fig 7.10 Earth Models



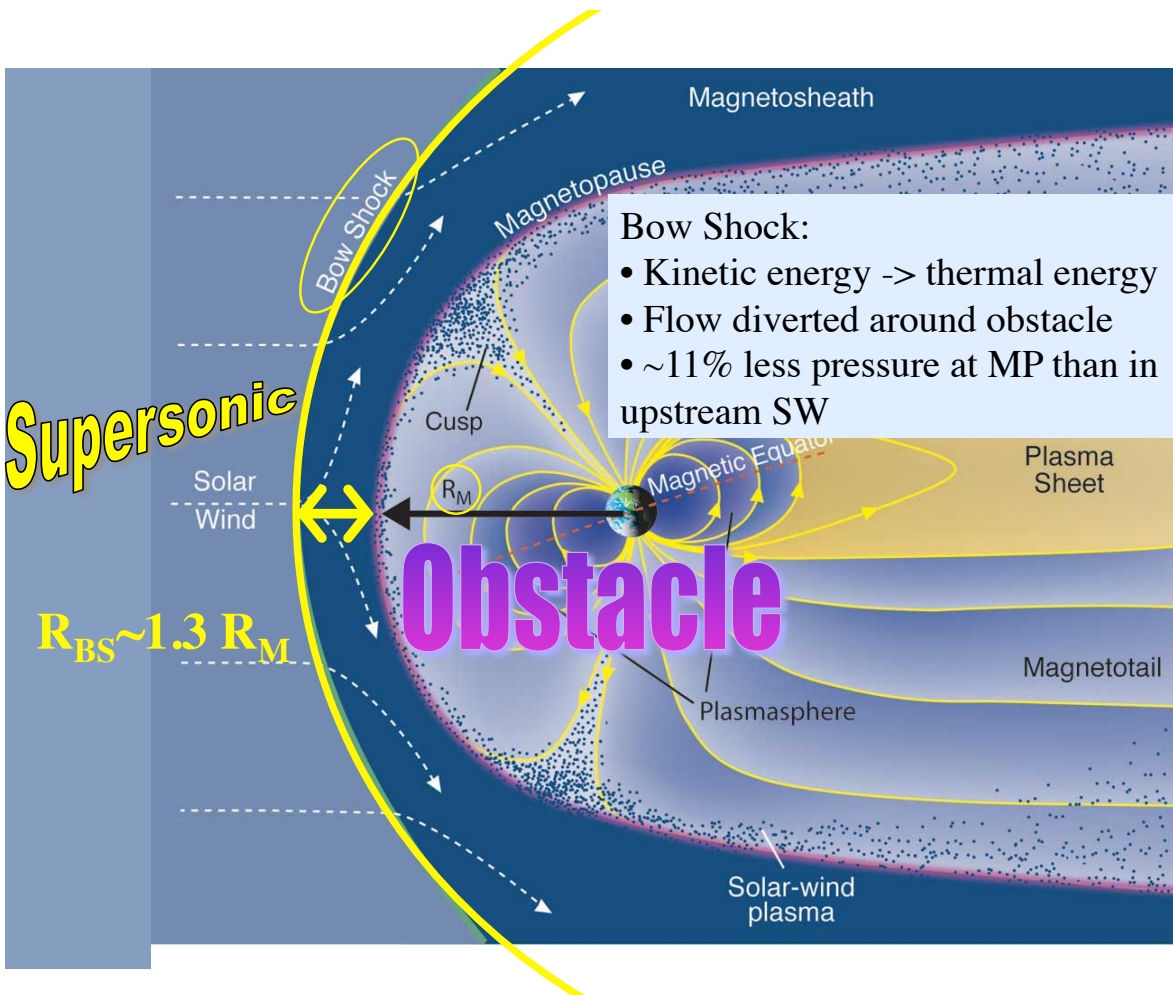
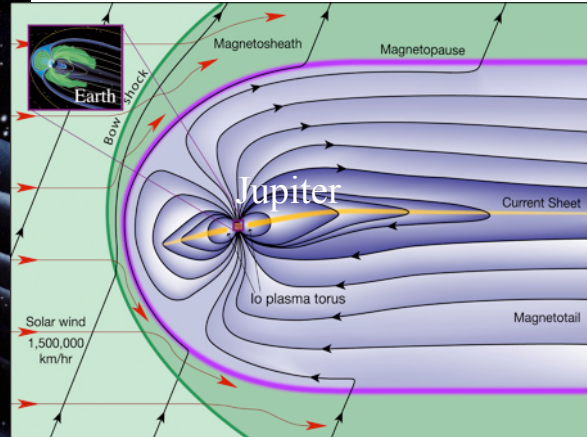
Core Heat Flux

Fig 7.11 Planets & Stars





Now we have magnetic fields.... what about magnetospheres?

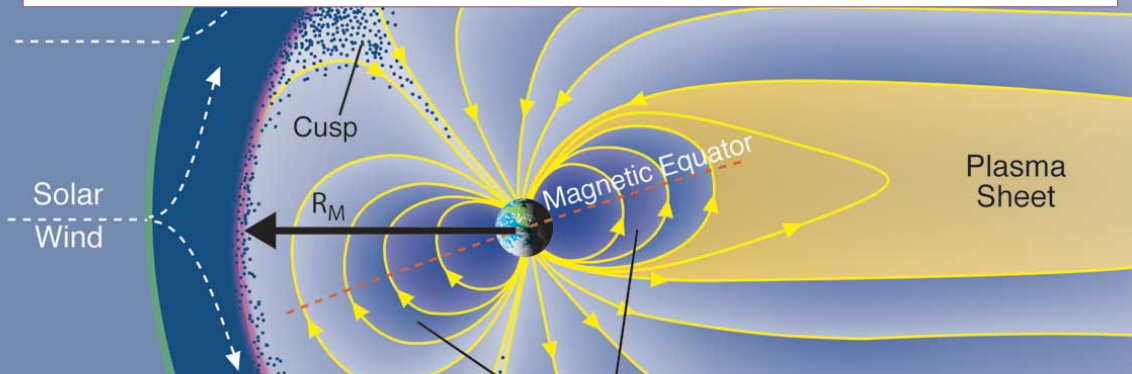


Bow Shock:

- Kinetic energy \rightarrow thermal energy
- Flow diverted around obstacle
- $\sim 11\%$ less pressure at MP than in upstream SW

Dipole Magnetic Field in Solar Wind

SW Ram Pressure \longleftrightarrow Magnetic Pressure

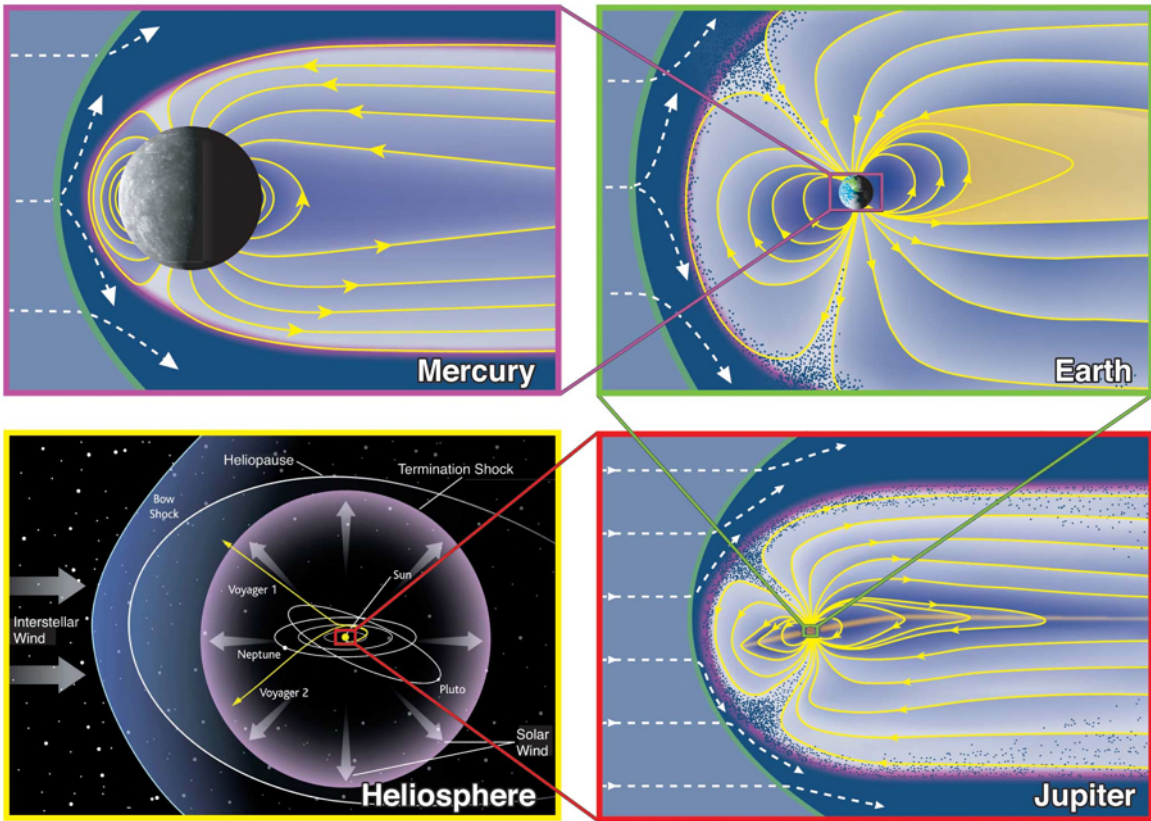


$$R_{MP} / R_{planet} = \left[B_o^2 / 2 \mu_o \rho_{sw} V_{sw}^2 \right]^{1/6}$$

Chapman-Ferraro Distance

$$R_{CF}/R_p \sim \{B_o^2 / 2 \mu_o \rho_{sw} V_{sw}^2\}^{1/6}$$

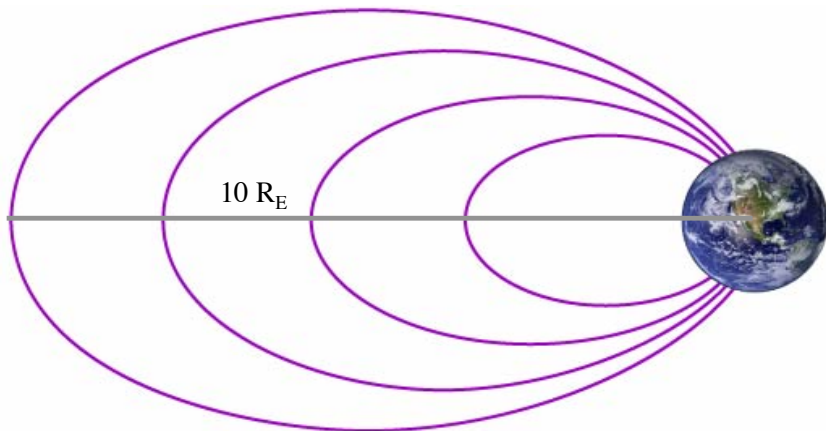
	Mercury	Earth	Jupiter	Saturn	Uranus	Neptune
B_o Gauss	.003	.31	4.28	.22	.23	.14
R_{CF} Calc.	1.4 R_M	10 R_E	42 R_J	19 R_S	25 R_U	24 R_N
R_M Obs.	1.4-1.6 R_M	8-12 R_E	60-90 R_J	16-22 R_S	18 R_U	23-26 R_N



Earth ~ Dipole

$$R_{mp} \sim (\rho V^2)^{-1/6}$$

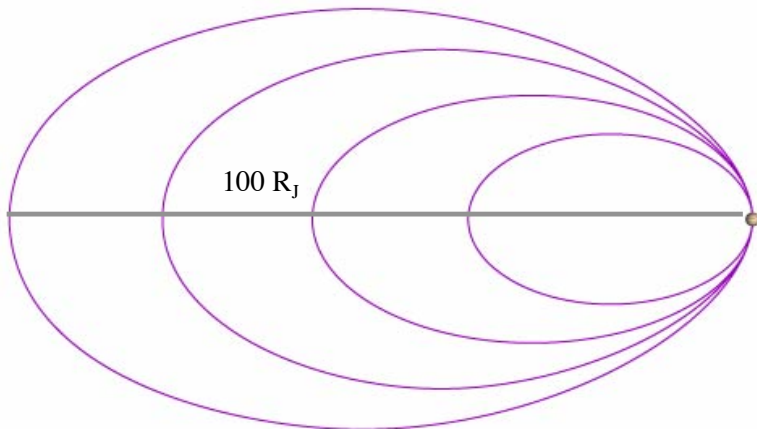
→ solar wind ρV^2



Jupiter

$$R_{mp} \sim (\rho V^2)^{-1/3}$$

→ solar wind ρV^2



Earth ~ Dipole

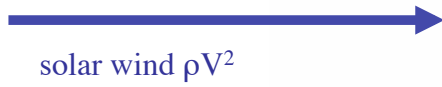
$$R_{mp} \rightarrow 0.7 R_{mp}$$



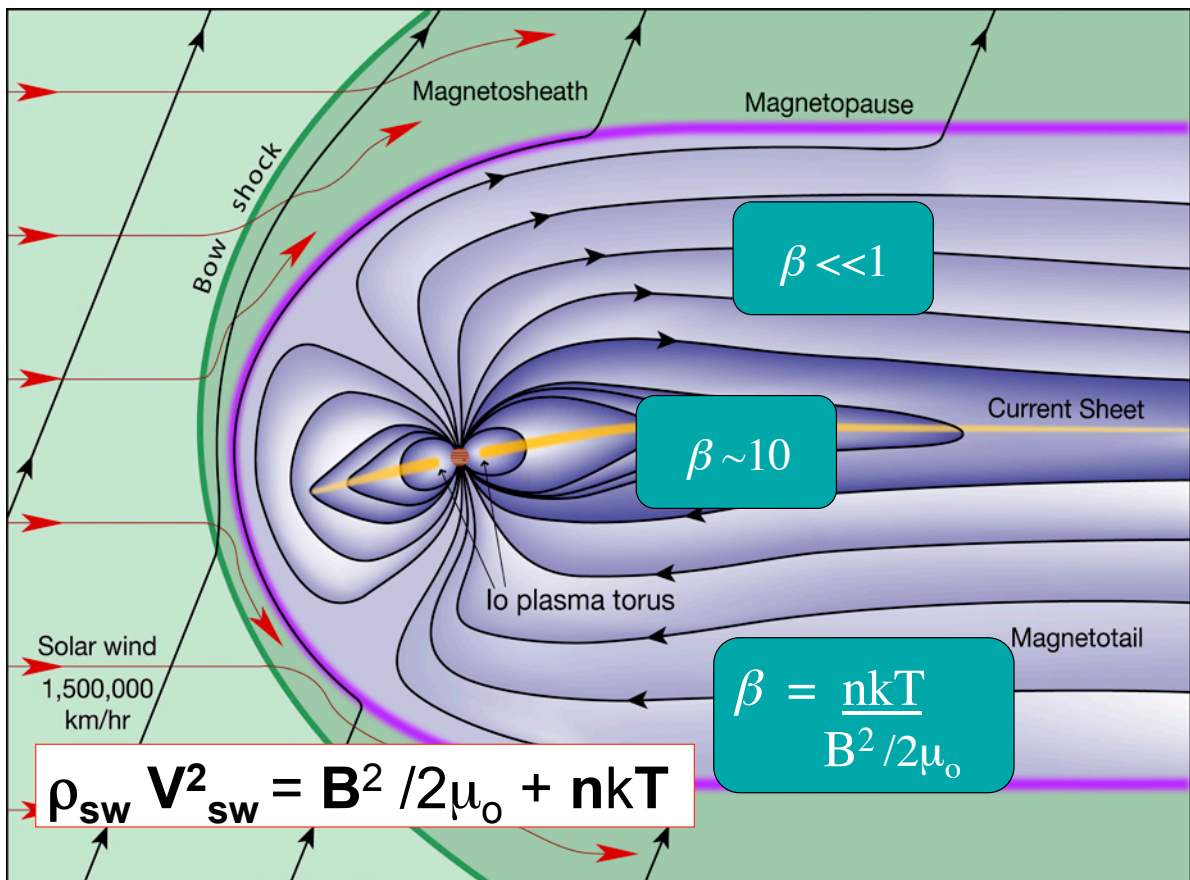
x10 Solar wind pressure

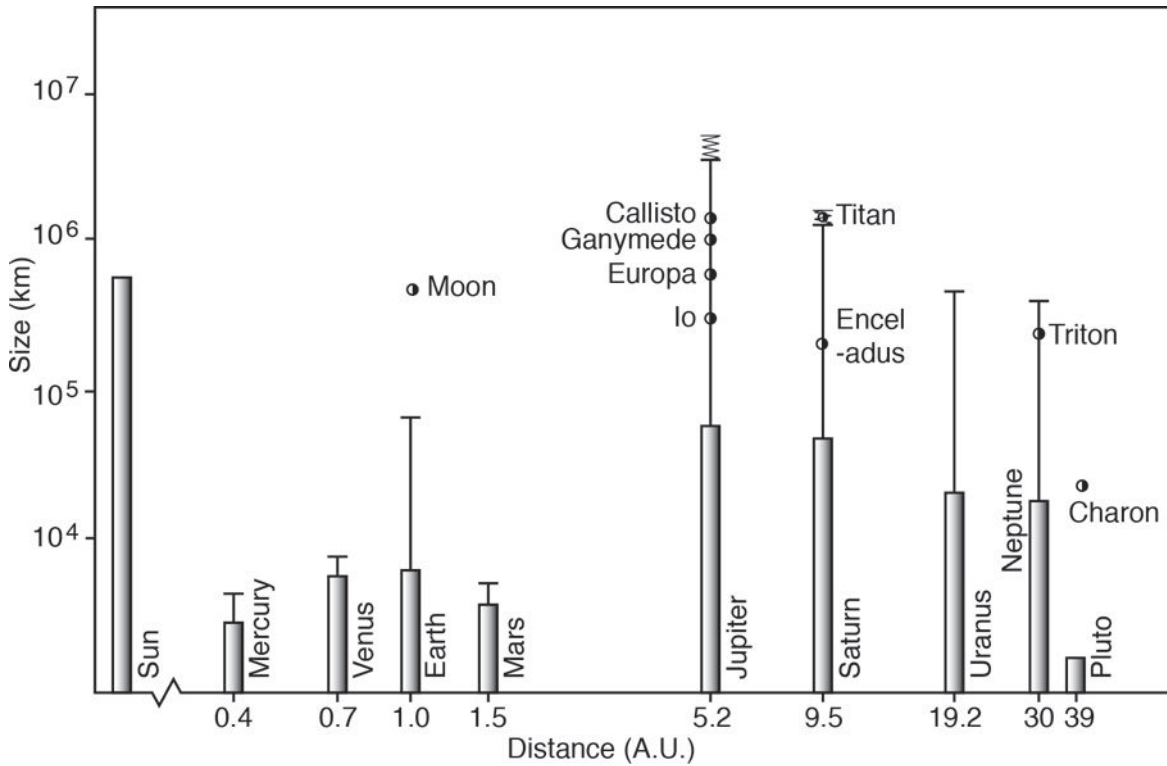
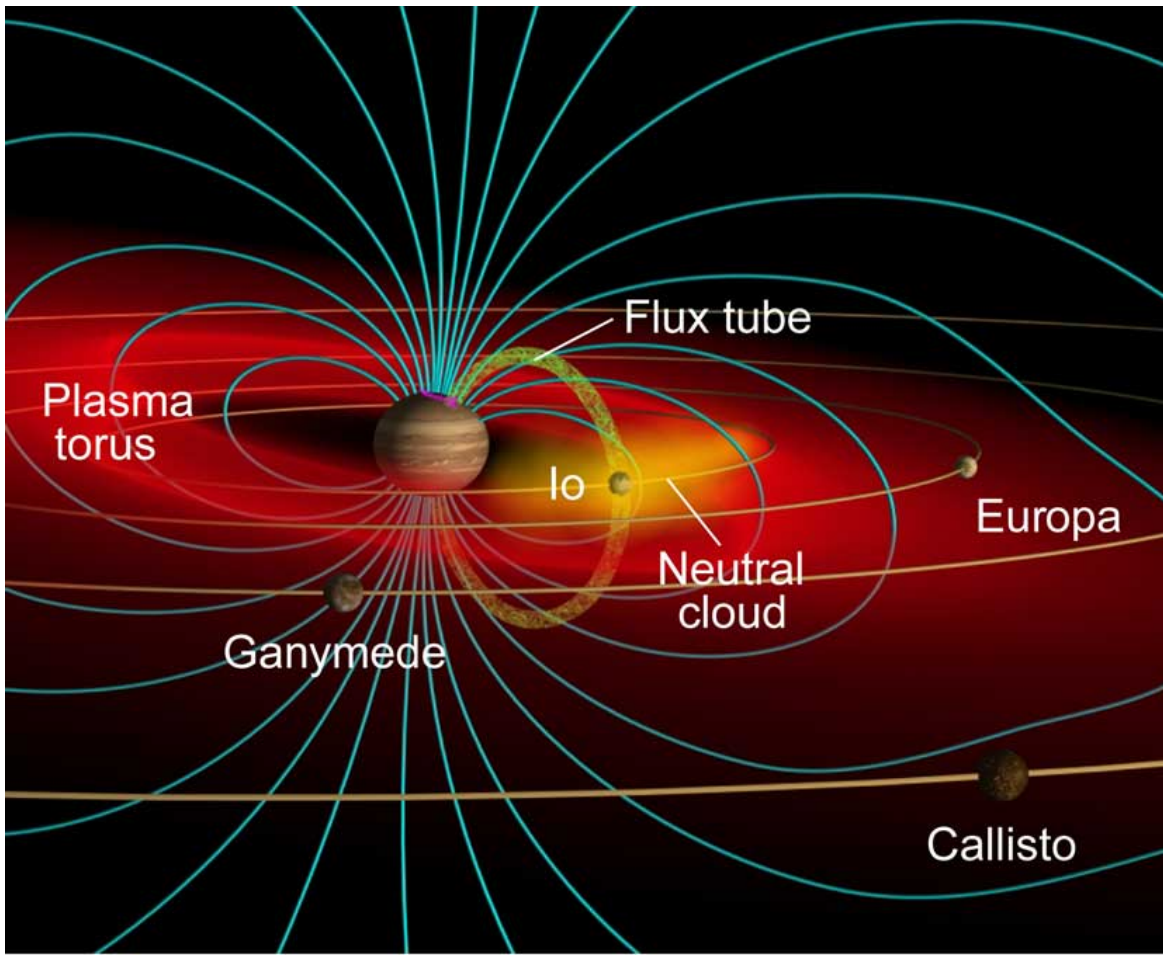
Jupiter

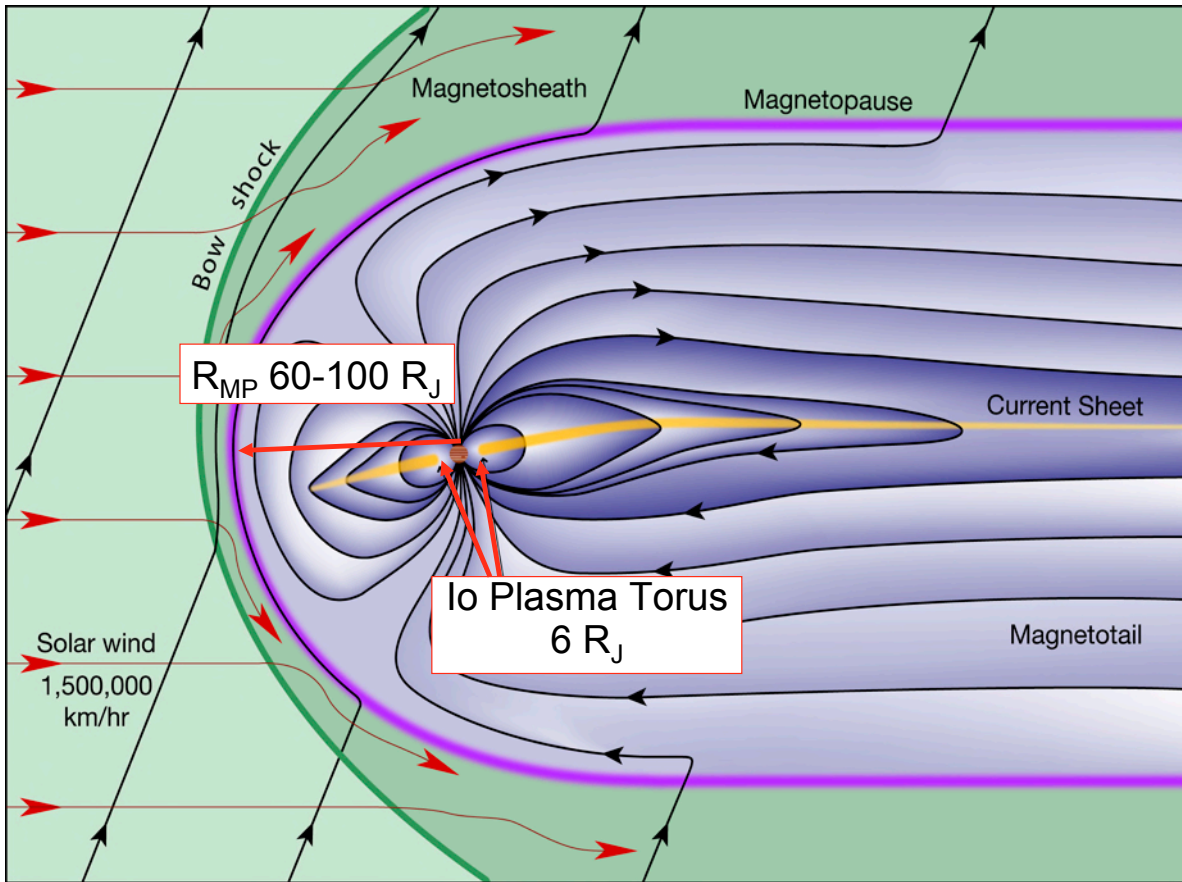
$$R_{mp} \rightarrow 0.5 R_{mp}$$



Factor ~10 variations in solar wind pressure at 5 AU -
> observed 100-50 R_J size of dayside magnetosphere







Solar wind interaction with planets – Vol. I Ch. 13 also Bagenal 2011

	Mercury	Venus	Earth	Mars	Jupiter	Saturn	Uranus	Neptune	Pluto
Distance from Sun, a_p (A.U.) ^a	0.39	0.72	1 ^b	1.52	5.2	9.5	19	30	40
Solar Wind Density ^b (cm ⁻³)	53	14	7	3	0.2	0.07	0.02	0.006	0.003
IMF strength ^c (nT)	41	14	8	5	1	0.6	0.3	0.2	0.1
IMF azimuth angle ^c	23°	38°	45°	57°	80°	84°	87°	88°	88°
Radius, R_p (km)	2,439	6,051	6,373	3,394	71,400	60,268	25,600	24,765	1,170 (±33)
Sidereal spin period (day)	58.6	-243	0.9973	1.026	0.41	0.44	-0.72	0.67	-6.39
Magnetic Moment ^d (M _E)	3-6 x 10 ⁻⁴	<10 ⁻⁵	1	<10 ⁻⁵	20,000	600	50	25	?
Surface Magnetic Field ^e B_0 (nT)	250-290	-	30,600	-	430,000	21,400	22,800	14,200	?
R_{CF} ^f (R_p)	1.6R _M	-	10 R _E	-	46 R _J	20 R _S	25 R _U	24 R _N	?
Observed R_{MP} (R_p)	1.5 R _M	-	8-12R _E	-	63-92 R _J	22-27 R _S	18 R _U	23-26 R _N	?

^a 1 A.U. = 1.5×10^8 km

^b The number density of the solar wind fluctuates by about a factor of 5 about typical values of $n_{sw} \sim 7$ (cm⁻³) / a_p^2]. The mass density of the solar wind is $\rho_{sw} = 1.04 n_{sw}$ (amu cm⁻³)

^c Mean values for the interplanetary magnetic field (IMF) in units of nano-Tesla with spherical components B_r , B_θ , B_ϕ . The azimuth angle is $\tan^{-1}(B_\phi/B_r)$. The radial component of the IMF, B_r , decreases as $1/a_p^2$ while the transverse component, B_ϕ , increases with distance.

^d M_{Earth} = 7.9×10^{25} Gauss cm³ = 7.9×10^{15} Tesla m³

^e Magnitude of dipole (see text for references).

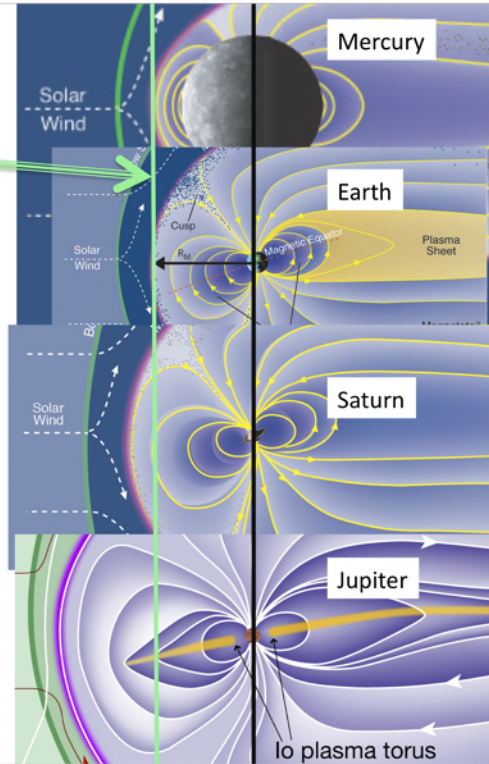
^f R_{CF} is calculated using $R_{CF} = \xi (B_0^2 / 2\mu_0 \rho V_{sw}^2)^{1/6}$ for typical solar wind conditions of ρ_{sw} given above and $V_{sw} \sim 400$ km s⁻¹ and ξ an empirical factor of ~ 1.4 to match Earth observations (Walker and Russell 1995).

Magnetospheres scaled by stand-off distance of dipole field

	M/M _E	MP _{Dipole}	MP _{mean}	MP _{Range}
Mercury	~8x10 ⁻³	1.4 R _M	1.4 R _M	
Earth	1	10 R _E	10 R _E	
Saturn	600	20 R _S	24 R _S	22-27* R _S
Jupiter	20,000	46 R _J	75 R _J	63-92# R _J

Inflated magnetospheres of Jupiter & Saturn due to HOT PLASMAS

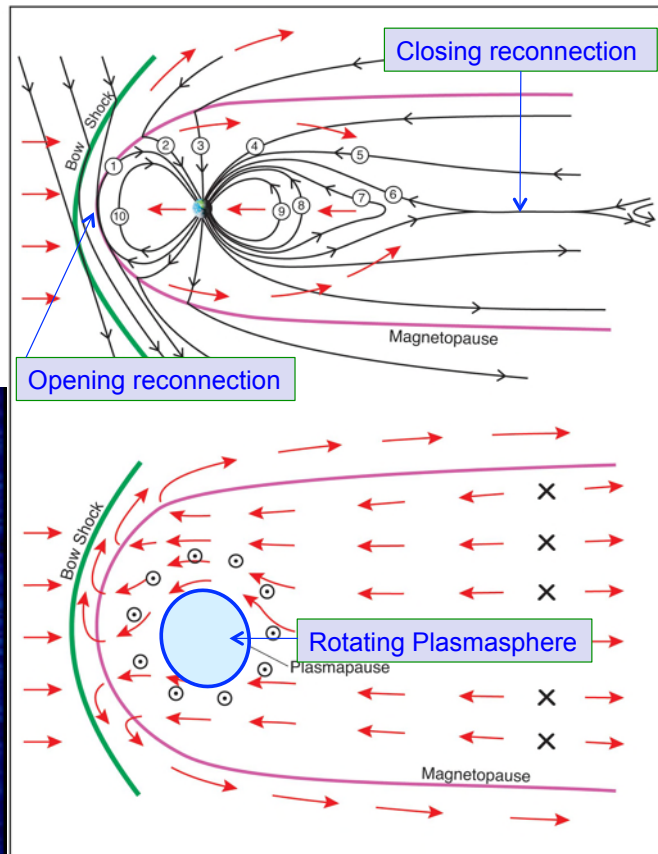
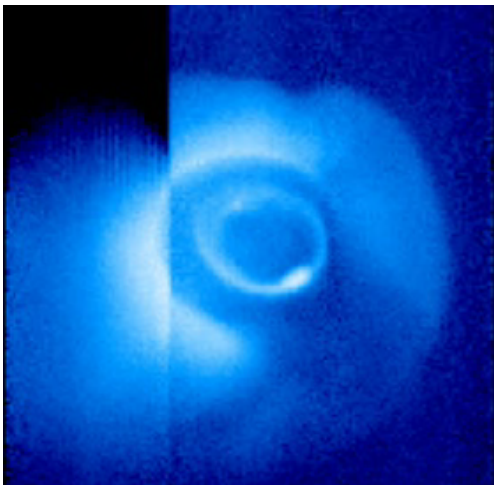
Note bimodal average locations
 * Achilles et al. 2008 # Joy et al. 2002



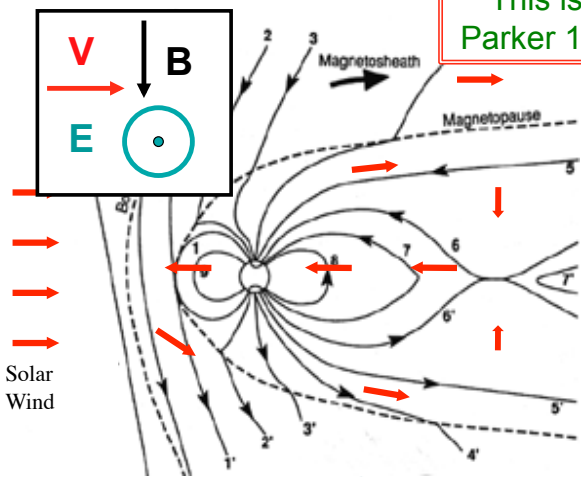
Lecture 4
 Vol. I Ch. 10

Dungey Cycle

Dynamics at Earth driven by the solar wind coupling the Sun's magnetic field to the Earth's field



This is the conventional E-J approach. See Parker 1996; Vasyliunas 2005 for B-V approach



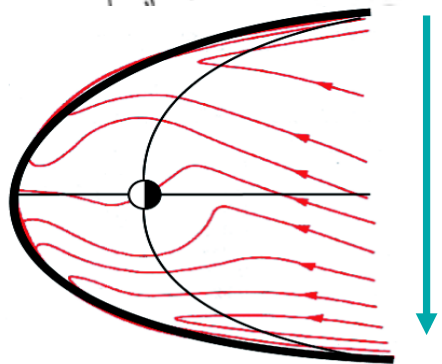
The Dungey Cycle
Solar wind driven magnetospheric convection*

$$\mathbf{E}_{\text{convection}} = -\zeta \mathbf{V}_{\text{SW}} \times \mathbf{B}_{\text{SW}}$$

$\zeta \sim$ efficiency of reconnection
 $\sim 10\text{-}20\%$

crude approximation!!

$$\mathbf{E}_{\text{conv}} \sim \text{constant in m'sphere}$$

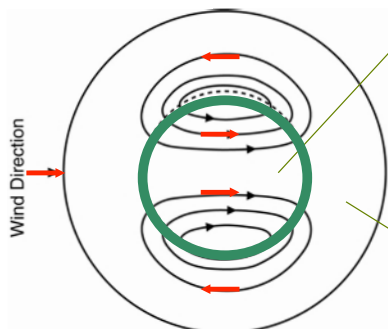
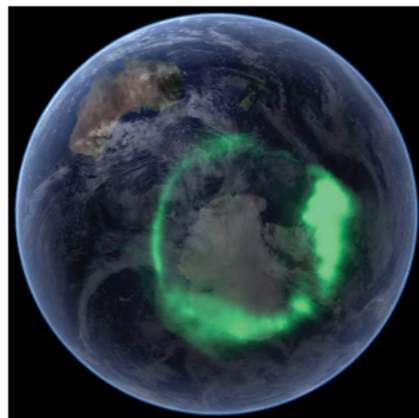
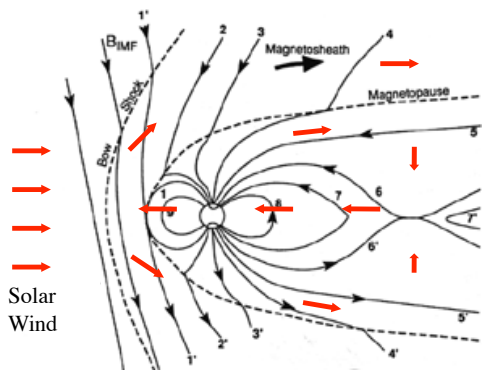


$$\mathbf{V}_{\text{convection}}$$

$$\sim \zeta V_{\text{SW}} (R/R_{\text{MP}})^3$$

(where 3 power assumes a dipole - in reality, the flow is not uniform and the power somewhat less)

(*strictly speaking not convection but advection or circulation)

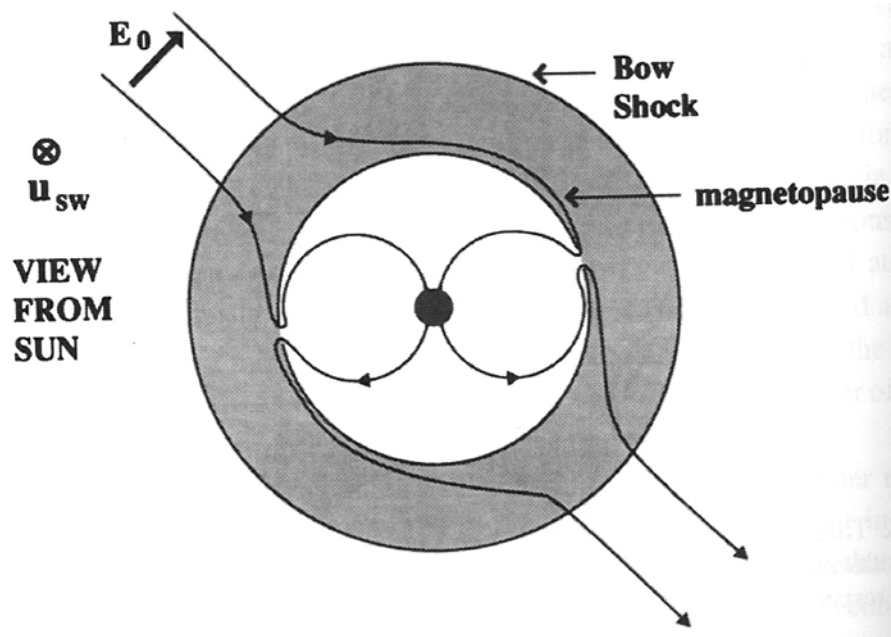


Connected to solar wind

Closed magnetic field

Polar view

Reality = Messy & 3D



Dynamics

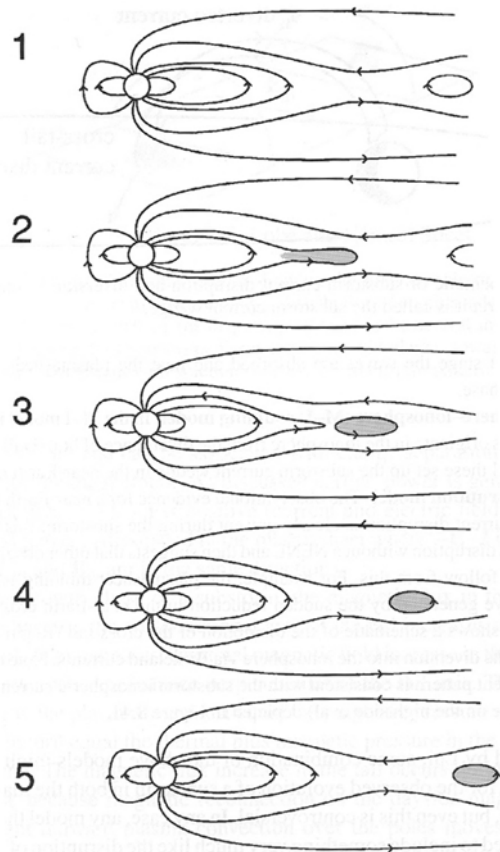
Dayside magnetopause

- Response to B_{sw} direction
- Solar wind ram pressure

Tail Reconnection

- Depends on recent history of dayside reconnection and state of plasmasheet

Space Weather!



$$\mathbf{V}_{CO} \sim \boldsymbol{\Omega} \times \mathbf{R}$$

$$\mathbf{V}_{\text{convection}}$$

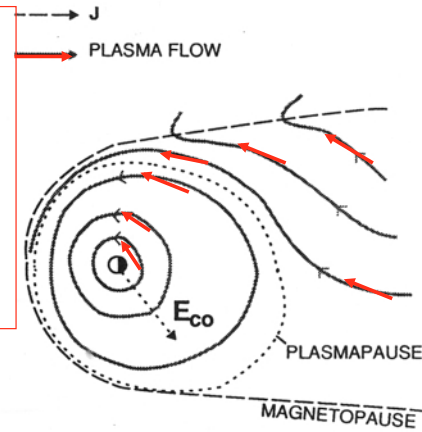
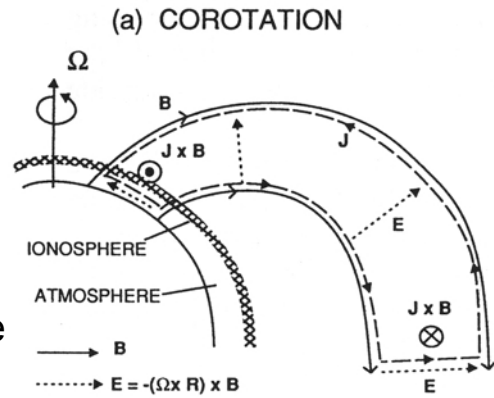
$$\sim \zeta V_{SW} (R/R_{MP})^3$$

Fraction of planetary magnetosphere that is rotation dominated is...

$$R_{pp}/R_{MP} \sim [r_p R_{MP} \Omega / \zeta V_{SW}]^{1/2} \propto \Omega^{1/2} \mu^{1/6} / (\rho_{SW})^{1/12} V_{SW}^{2/3}$$

Where r_p = planetary radius

μ = magnetic moment of planet $B_0 R_p^3$



Magnetospheric Dynamics – Vol. I Ch. 13 also Bagenal 2011

	Mercury	Earth	Jupiter	Saturn	Uranus	Neptune
R_{MP}^a (km)	4000	6.5×10^4	6×10^6	1×10^6	6×10^5	6×10^5
V_{SW}^b (km/s)	370	390	420	430	450	460
t_{N-T}^c (s)	10 s	3 min	4 hr	45 min	20 min	20 min
R_T^d (R_p)	3	20	170	40	50	50
R_T^d (km)	8000	1.3×10^5	1.2×10^7	2.3×10^6	1.3×10^6	1.2×10^6
$V_{rec,1}^e$ (km/s)	40	22	16	16	16	16
$V_{rec,2}^f$ (km/s)	37	39	42	43	45	46
t_{rec}^g (s)	3 min	1 hr	80 hr	15 hr	8 hr	7 hr
d_X^h (R_p)	30	200	1700	400	500	500
$V_{co}/V_{rec,2}^i$	4×10^{-5}	0.04	8	1.3	0.4	0.4
R_{pp}^j (R_p)	0.03	6.7	350	95	70	70

^a Sub-solar magnetopause distance.

^b $V_{SW} = 387 (a_p/a_E)^{0.05}$ (km/s) from Belcher et al. (1993)

^c Solar wind nose-terminator time: $t_{N-T} = R_{MP} / V_{SW}$

^d Radius of cross section of magnetotail, approximated as $R_T = 2R_{MP}$.

^e Reconnection speed assuming 20% reconnection efficiency and $v_{rec} \sim 0.2 v_{SW} B_{SW} / B_{MP}$ km/s (e.g. Kivelson 2007)

^f Reconnection speed assuming 10% reconnection efficiency and $v_{rec} \sim 0.1 v_{SW}$ km/s

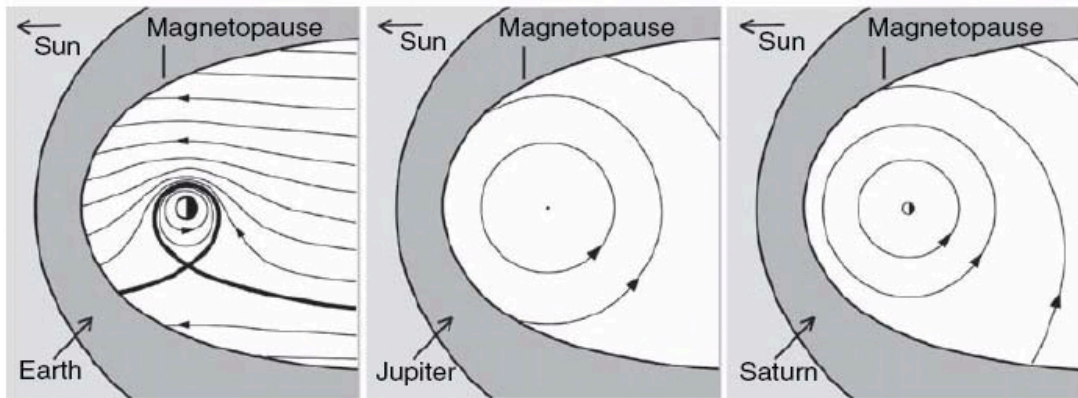
^g Reconnection time $t_{rec} = R_T / v_{rec,2}$ (s)

^h Distance to X-line $d_X \sim v_{SW} t_{rec}$

ⁱ Assumes rotation speed at the magnetopause is ~30% of rigid corotation

^j Distance to plasmopause, where corotation is comparable to reconnection flow (e.g. Kivelson 2007)

Solar-wind vs. Rotation-dominated magnetospheres



$$R_{\text{plasmopause}} / R_{\text{Planet}} =$$

6.7

350

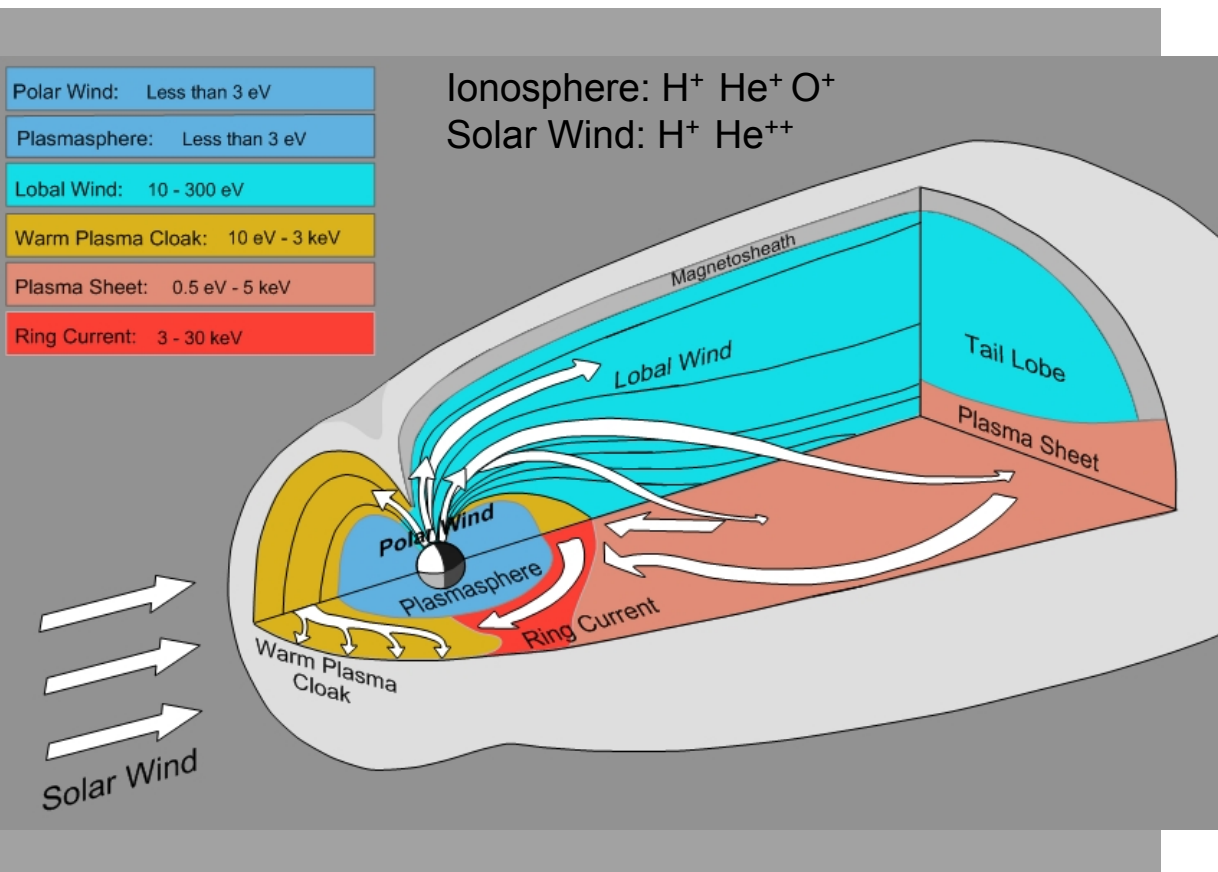
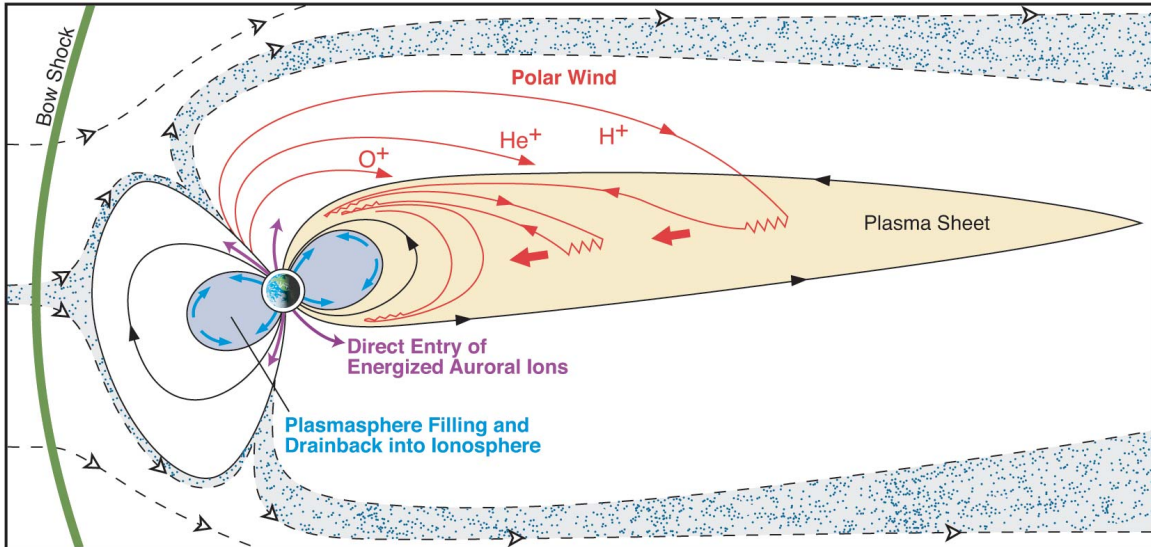
95

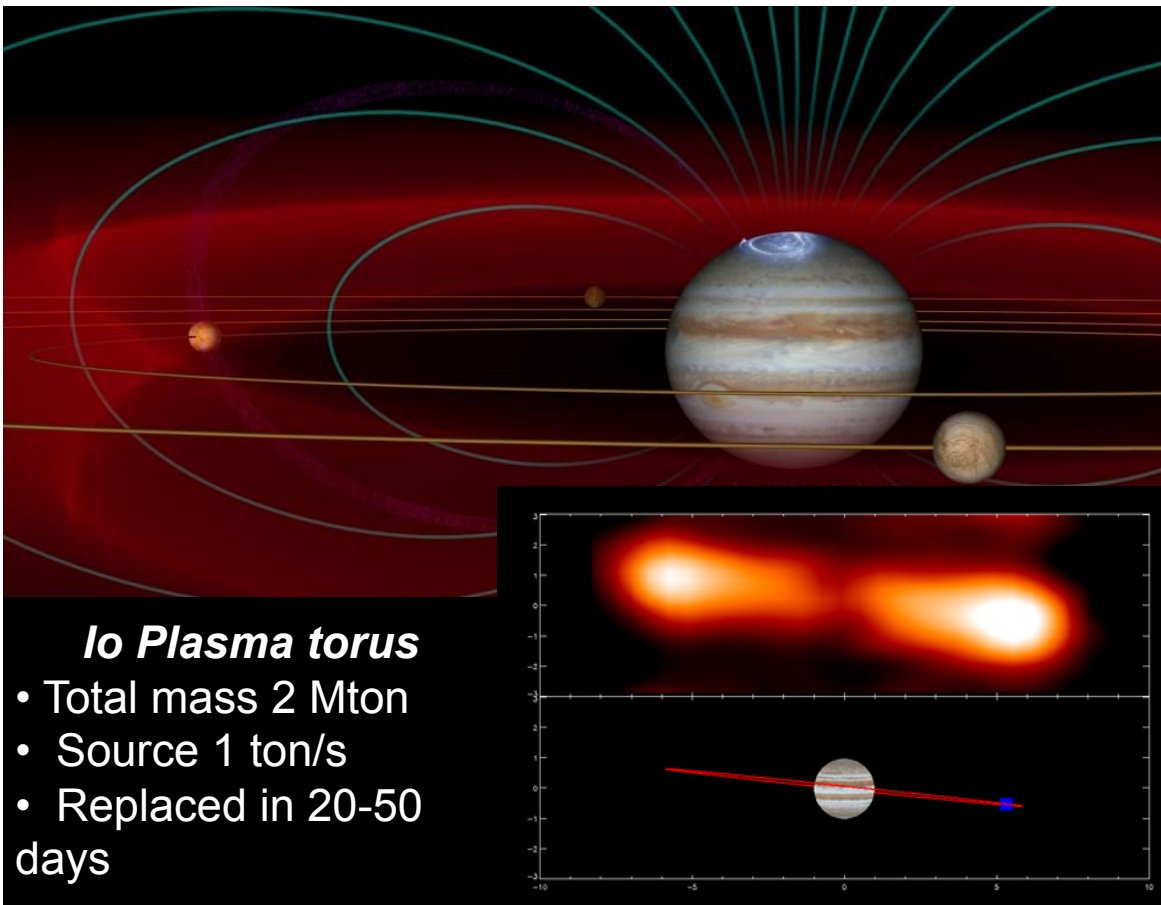
Plasma Sources

	Mercury	Earth	Jupiter	Saturn	Uranus	Neptune
N_{max} cm^{-3}	~1	1-4000	>3000	~100	~3	~2
Composition	H^+ Solar Wind	O^+ H^+ Ionosphere	O^+ S^+ I^+ Io	O^+ H_2O^+ H^+ Enceladus	H^+ Ionosphere	H^+ N^+ Triton Ionosphere
Source kg / s	?	5	700-1200	~20 70-700?	~0.02	~0.2

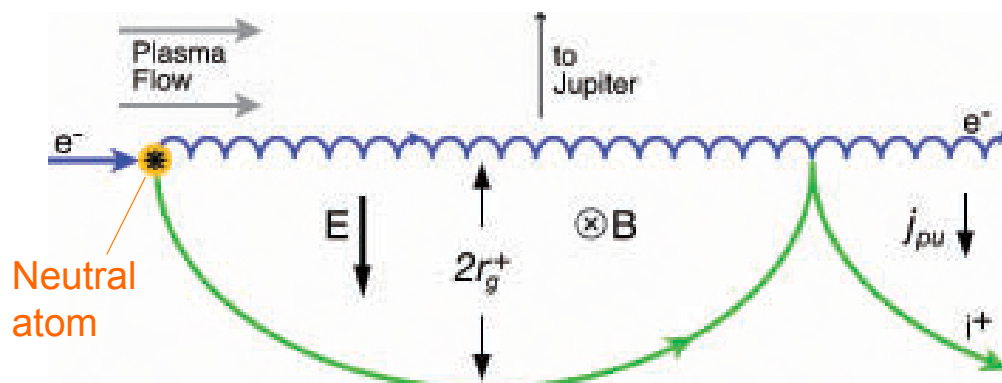
Sources of Plasma:

Solar Wind + ionosphere mixed (over the poles) into magnetotail and convected sunward



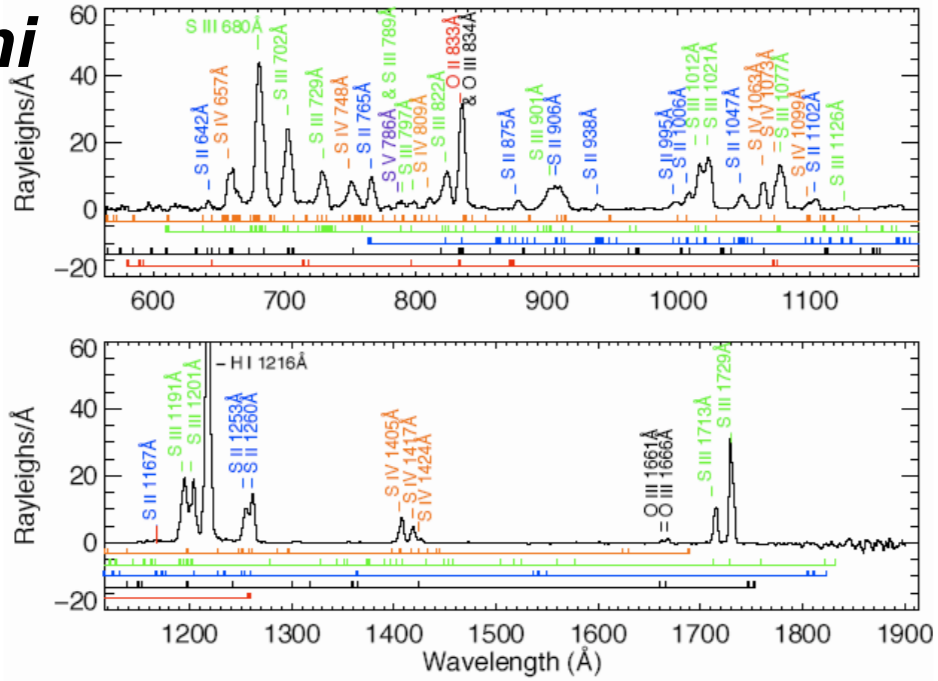


Ion Pick Up

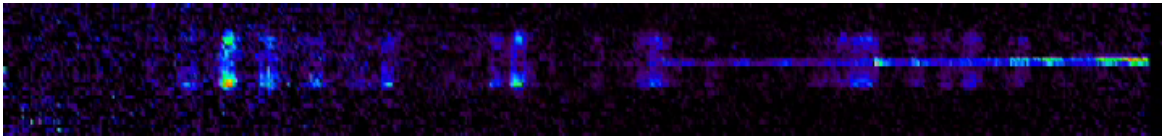


The magnetic field couples the plasma to the spinning planet

Cassini UVIS



Andrew
Steffl



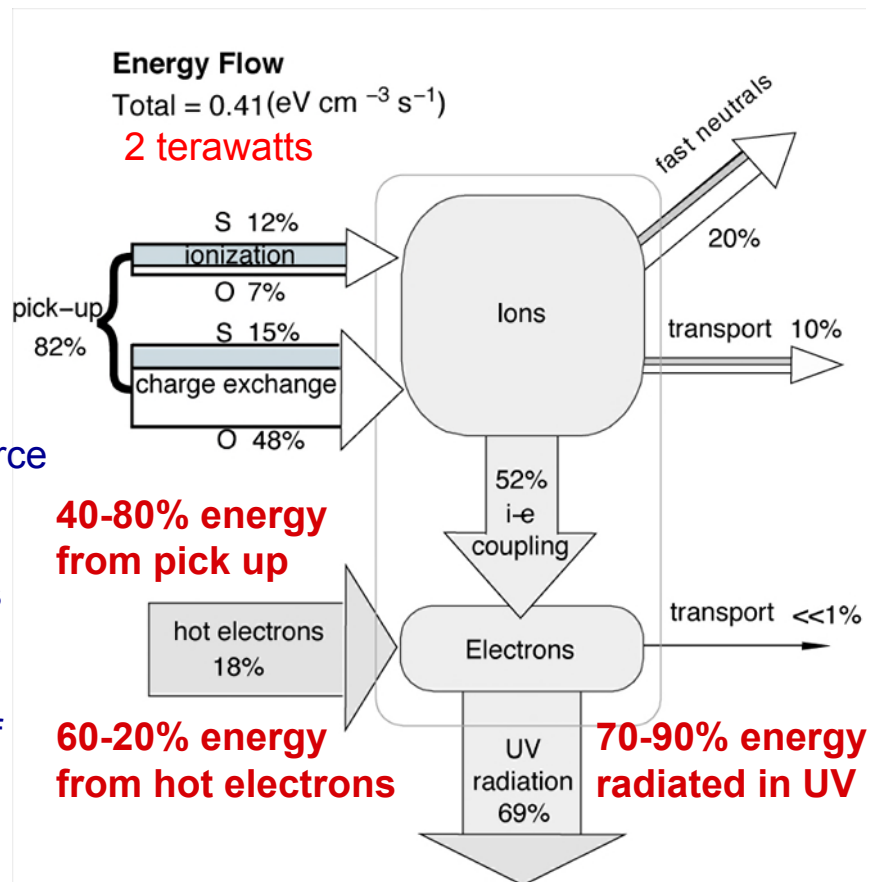
Energy Crisis!

- Total thermal energy $\sim 6 \times 10^{17}$ J

- Ion pick-up source takes 4 days

- UV power cools electrons ~ 7 hrs

- Extra source of hot electrons



Jupiter's 3 Types of Aurora

Steady Main
Auroral Oval

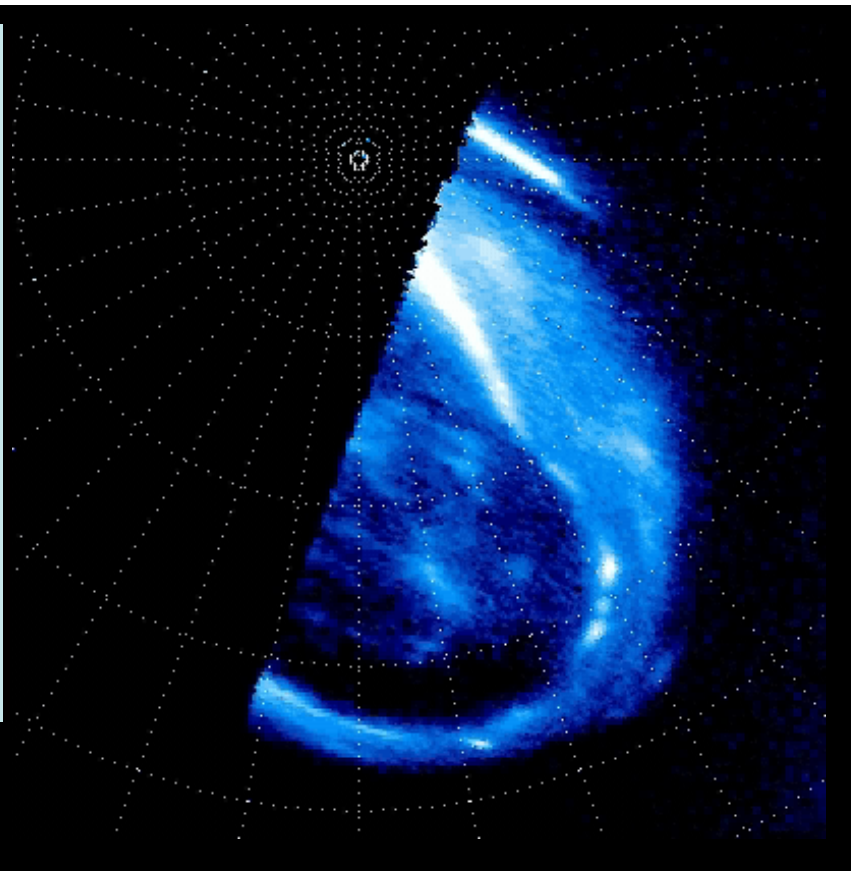
Variable
Polar Aurora

Aurora associated with moons

*Jupiter's
Aurora -
The Movie*

*Fixed
magnetic
co-
ordinates
rotating
with Jupiter*

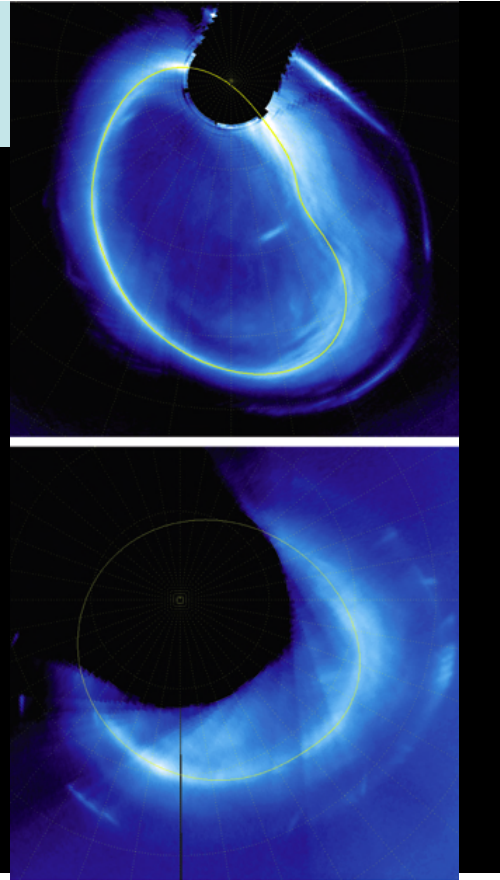
*Clarke et al.
Grodent et al.
HST*



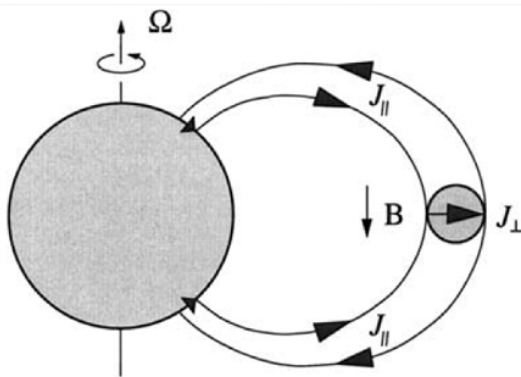
Main Aurora

- Shape constant, fixed in magnetic co-ordinates
- Magnetic anomaly in north
- Steady intensity
- $\sim 1^\circ$ Narrow

Clarke et al., Grodent et al. HST

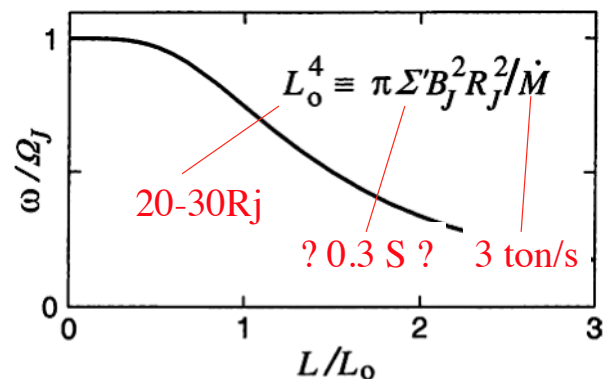


The aurora is the signature of Jupiter's attempt to spin up its magnetosphere



Hill 1979

- Outward transport of Iogenic plasma
- J_{\parallel} transfers load to ionosphere
- Transfer of angular momentum limited by ionospheric conductivity Σ

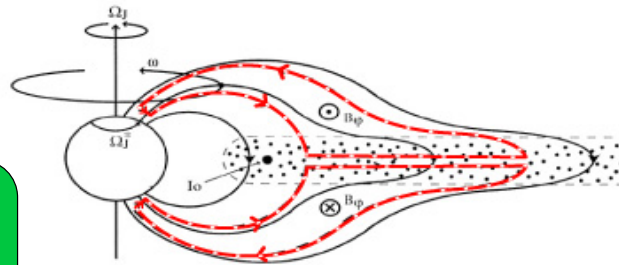
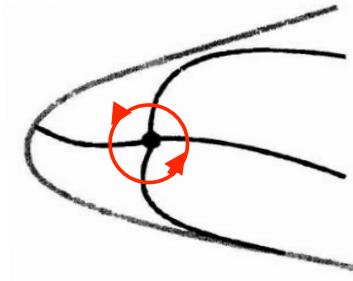


N.B. L_0 has units

Coupling the Plasma to the Flywheel

- As plasma from Io moves outwards its rotation decreases (conservation of angular momentum)
- Sub-corotating plasma pulls back the magnetic field
- $\text{Curl } \mathbf{B} \rightarrow$ radial current J_r
- $J_r \times \mathbf{B}$ force enforces rotation

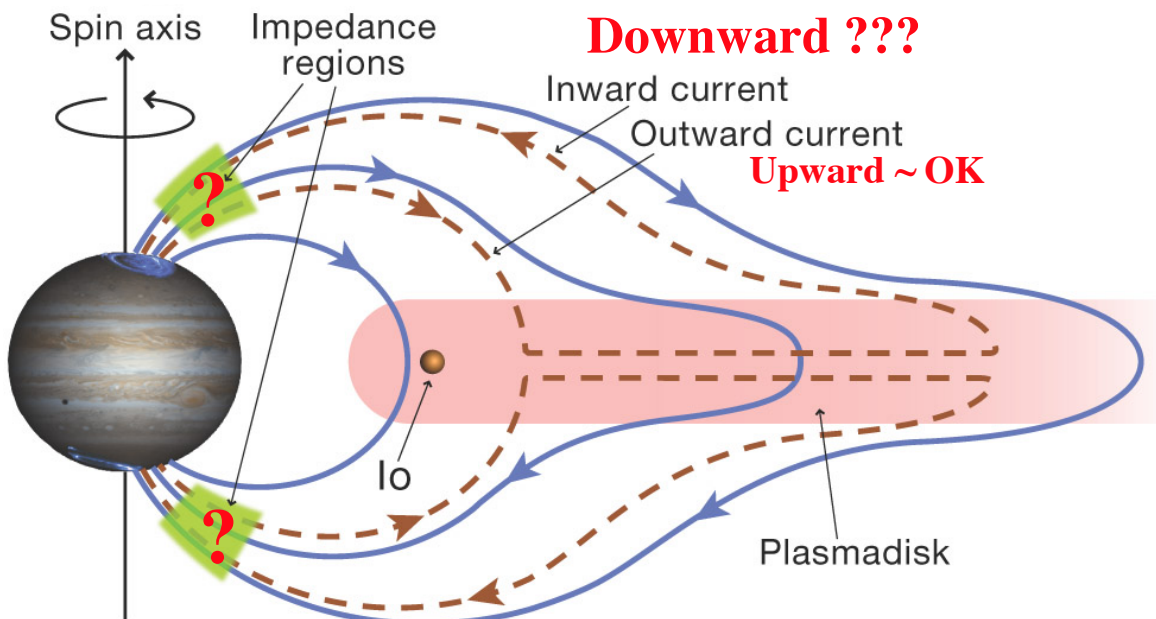
Khurana 2001



Cowley & Bunce 2001

Field-aligned currents couple magnetosphere to Jupiter's rotation

The aurora is the signature of Jupiter's attempt to spin up its magnetosphere



Where is the clutch slipping?

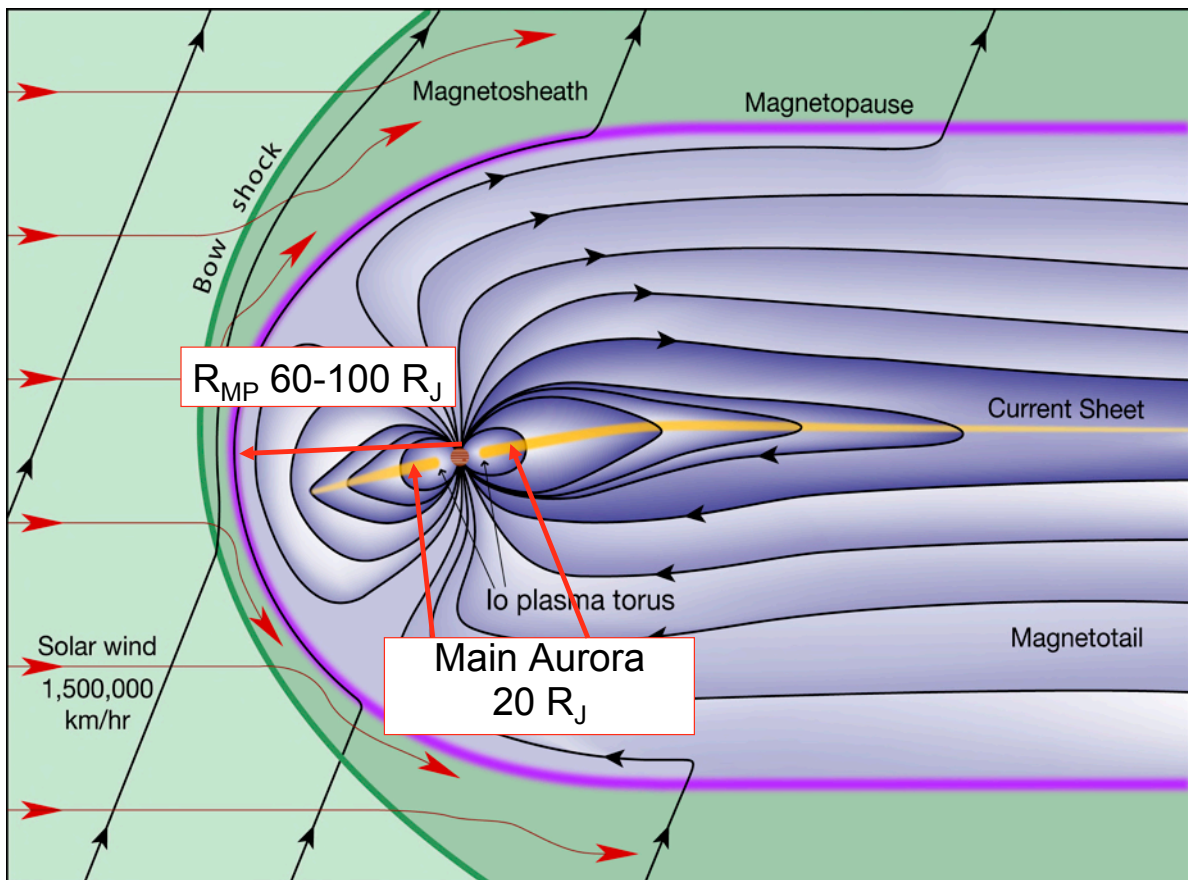
Mass loading



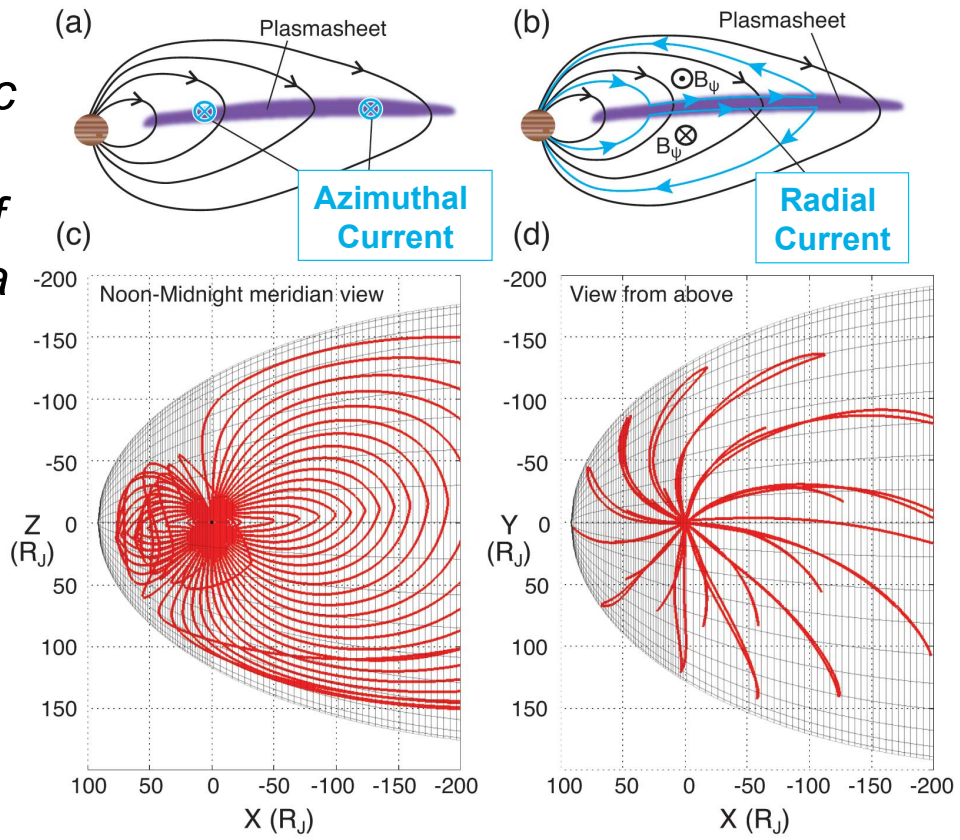
A - Between deep and upper atmosphere?

B - Between upper atmosphere and ionosphere?

C - Lack of current-carriers in magnetosphere $\rightarrow E_{\parallel}$?



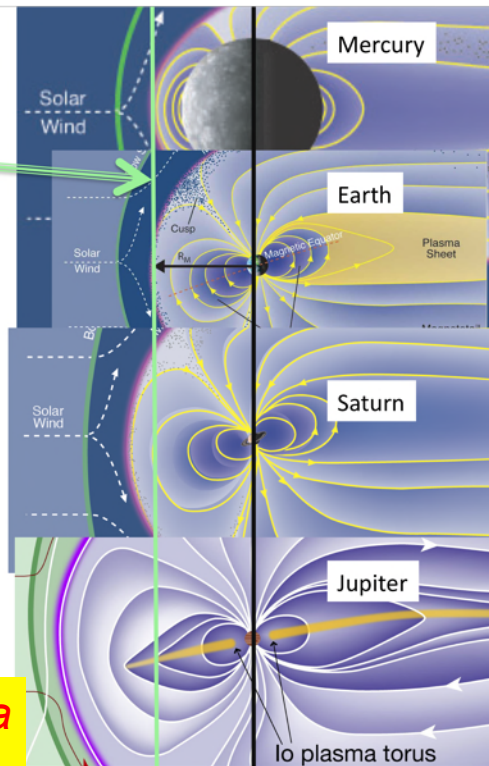
Magnetic field model of Khurana



Magnetospheres scaled by stand-off distance of dipole field

	M/M_E	MP_{Dipole}	MP_{mean}	MP_{Range}
Mercury	$\sim 8 \times 10^{-3}$	$1.4 R_M$	$1.4 R_M$	
Earth	1	$10 R_E$	$10 R_E$	
Saturn	600	$20 R_S$	$24 R_S$	$22-27^* R_S$
Jupiter	20,000	$46 R_J$	$75 R_J$	$63-92^{\#} R_J$

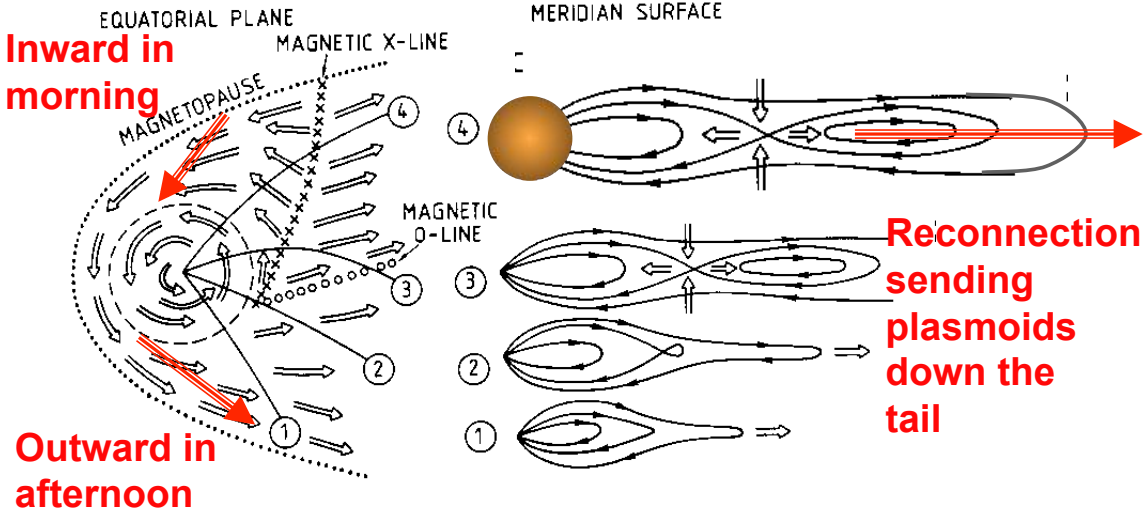
Inflated magnetospheres of Jupiter & Saturn due to HOT PLASMAS



What process heats the plasma as it moves outwards???

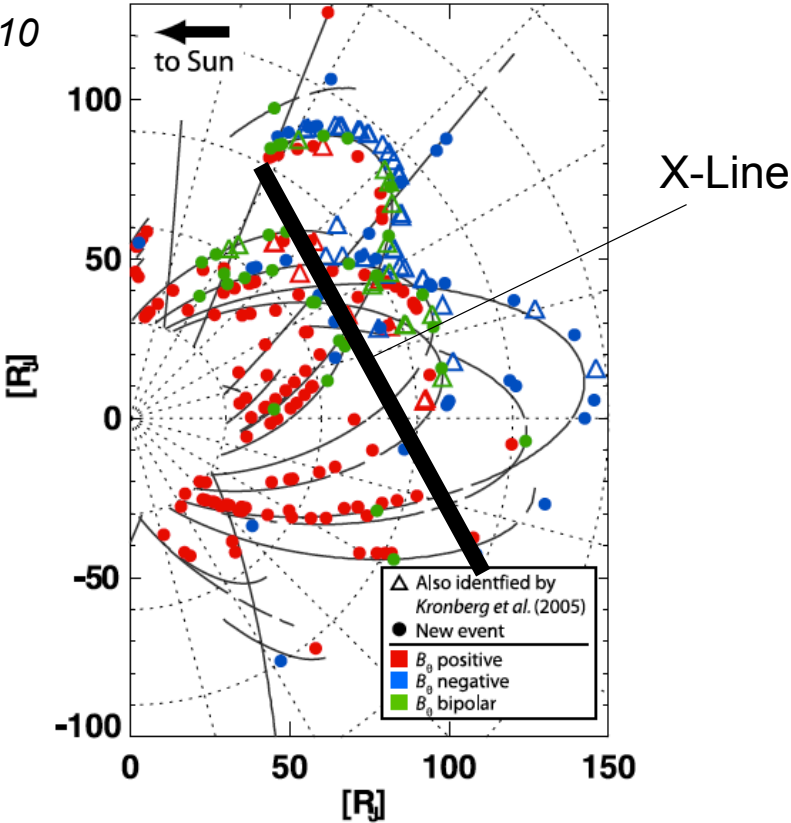
Vasyliunas Cycle

Vasyliunas
 Cowley et al.
 Southwood & Kivelson



Vogt et al. 2010

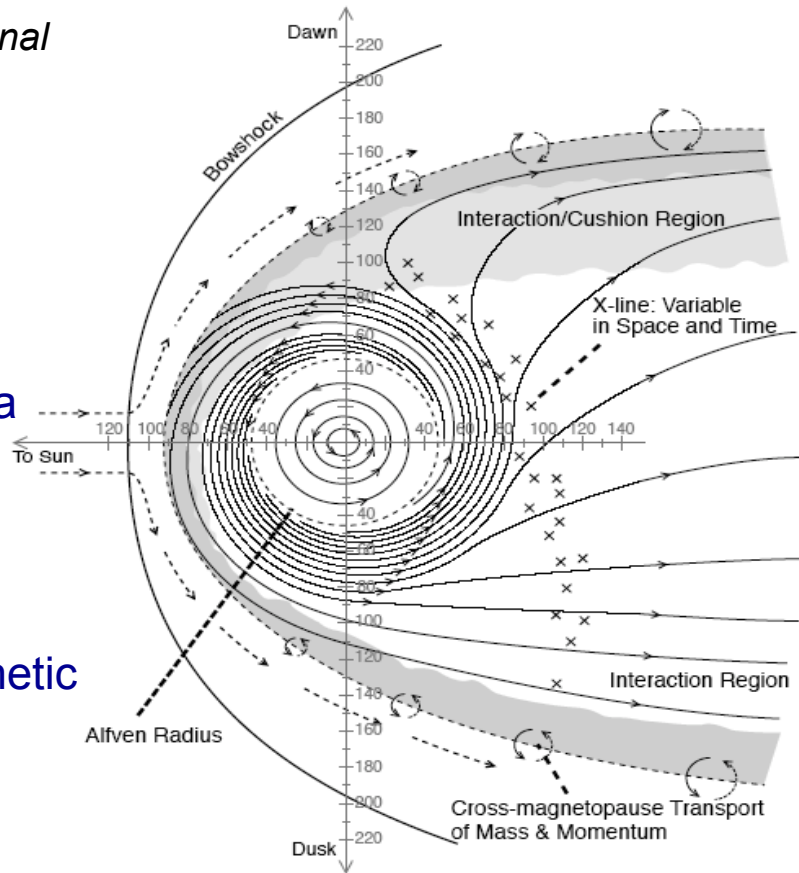
Observations of plasmoid events in Galileo data



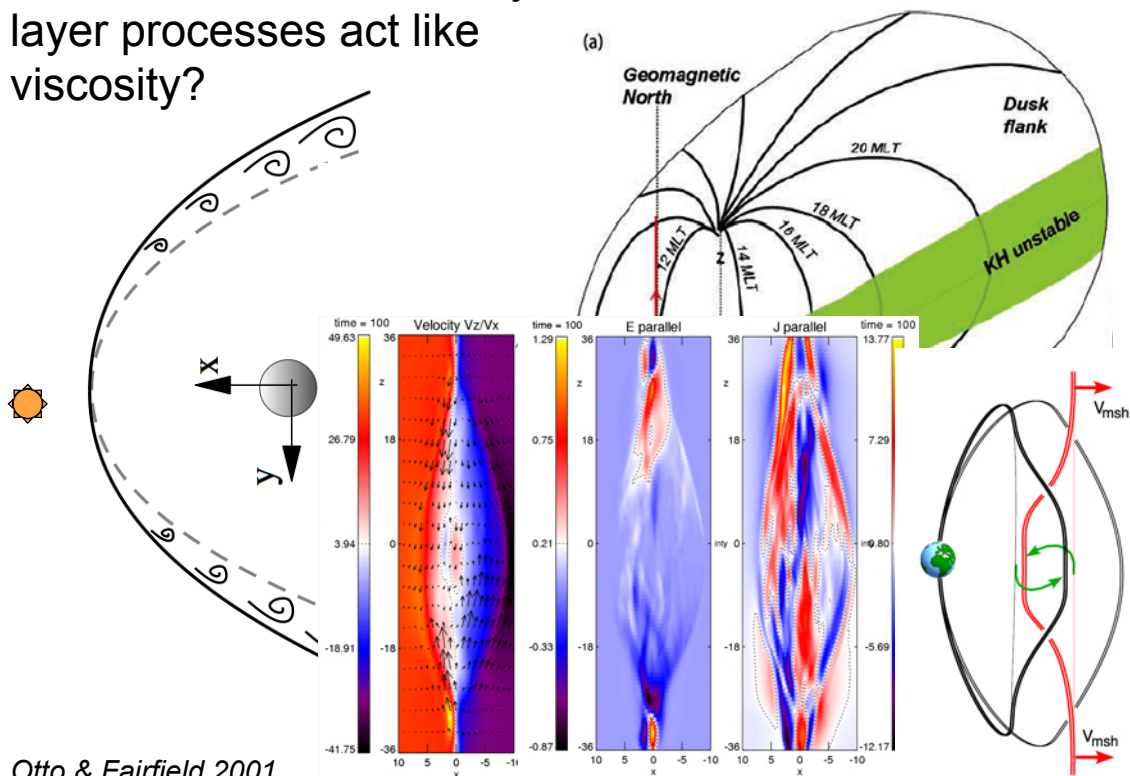
Delamere & Bagenal
(2011)

Solar wind
interaction:

- More of a plasma-plasma interaction
- Less of an interaction between magnetic fields



Can small-scale boundary-layer processes act like viscosity?

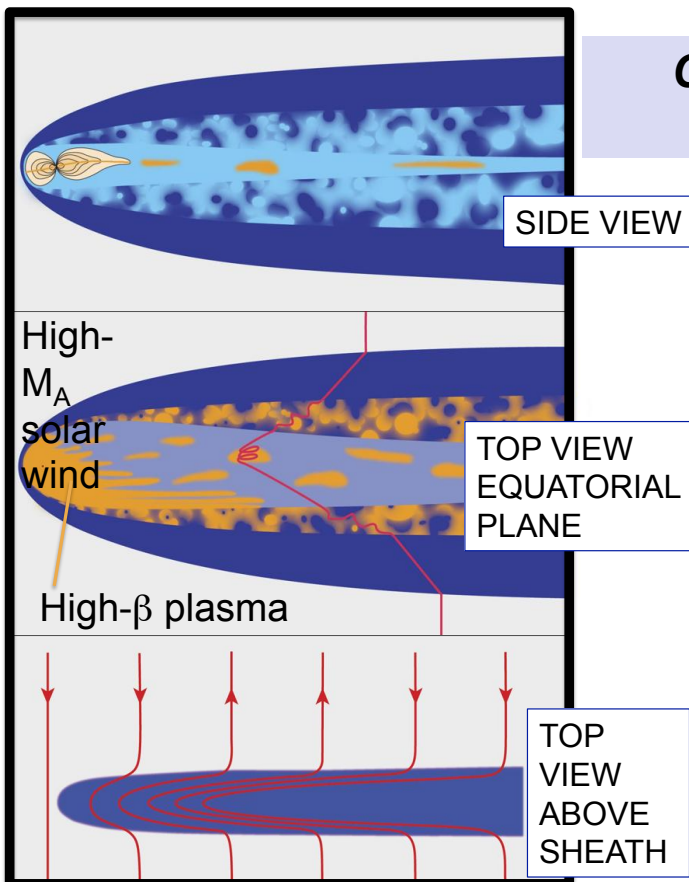
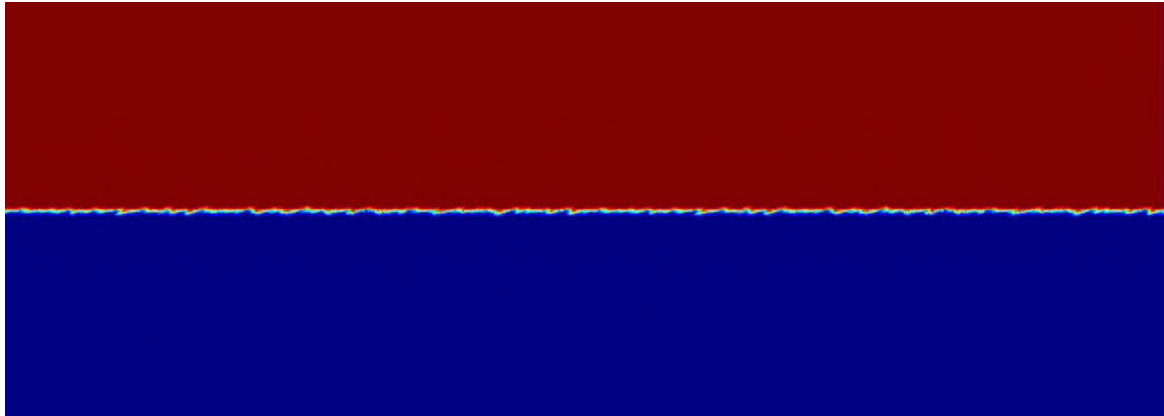
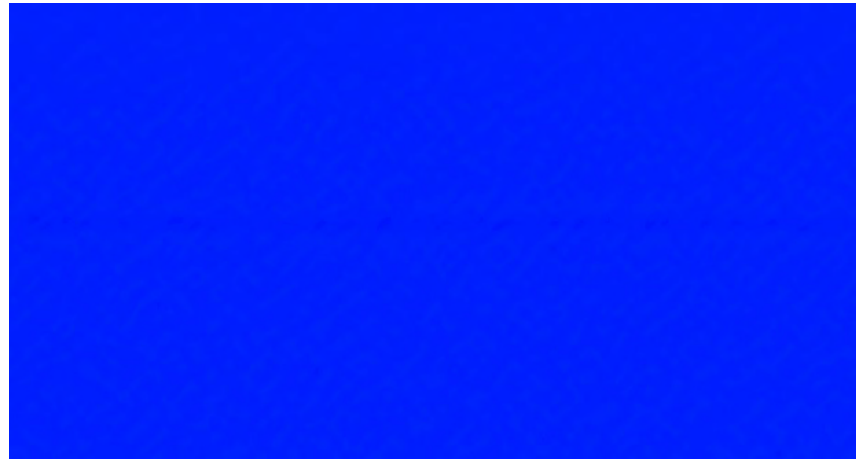


Otto & Fairfield 2001

Hybrid simulations of Kelvin-Helmholtz instability
 - ions=particles
 - electrons=fluid

Peter Delamere, CU

Heavy Light



Could Jupiter be a Colossal Comet?

- Plasma-plasma interaction with magnetic field playing less of a role than at Earth
- Solar wind hung up on the boundary layers
- Venus- or comet-like rather than field-controlled terrestrial tail.



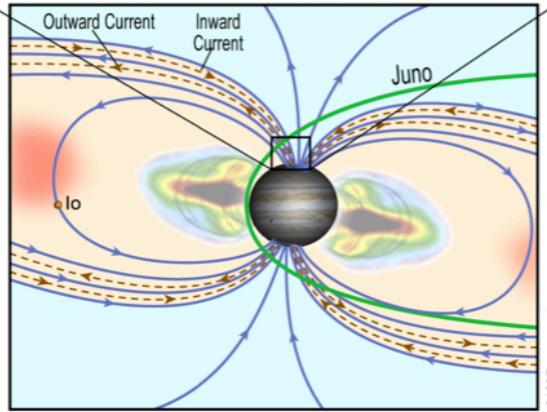
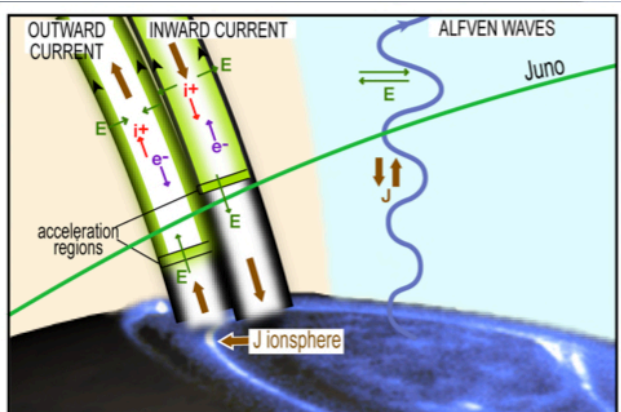
Polar Magnetosphere

Juno passes directly through auroral field lines

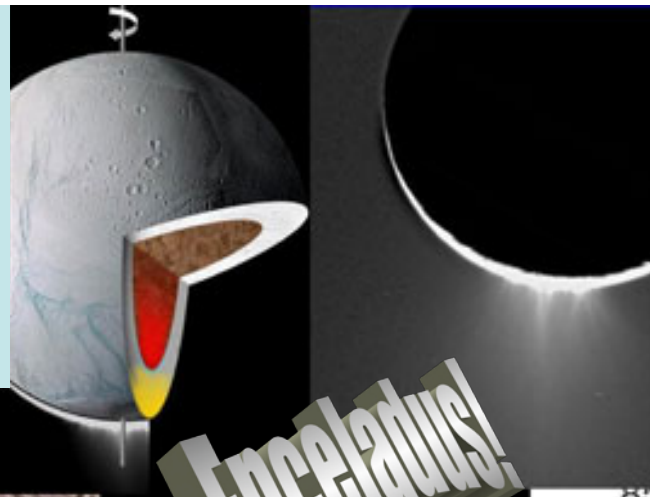
Measures particles precipitating into atmosphere creating aurora

Plasma/radio waves reveal processes responsible for particle acceleration

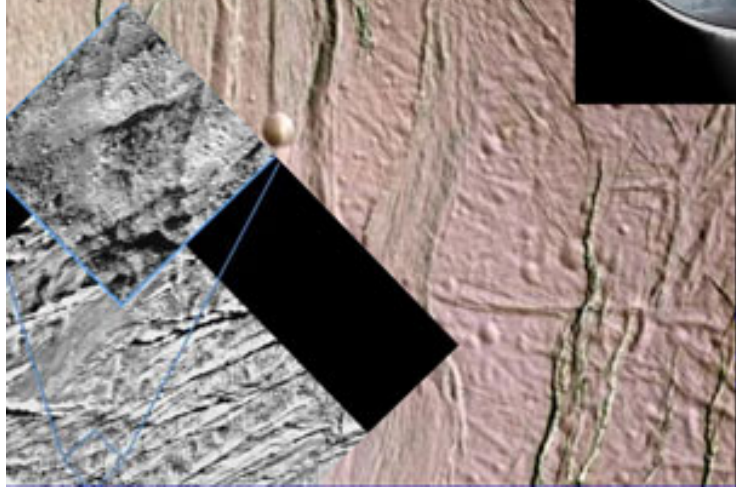
UV & IR images provides context for *in-situ* observations



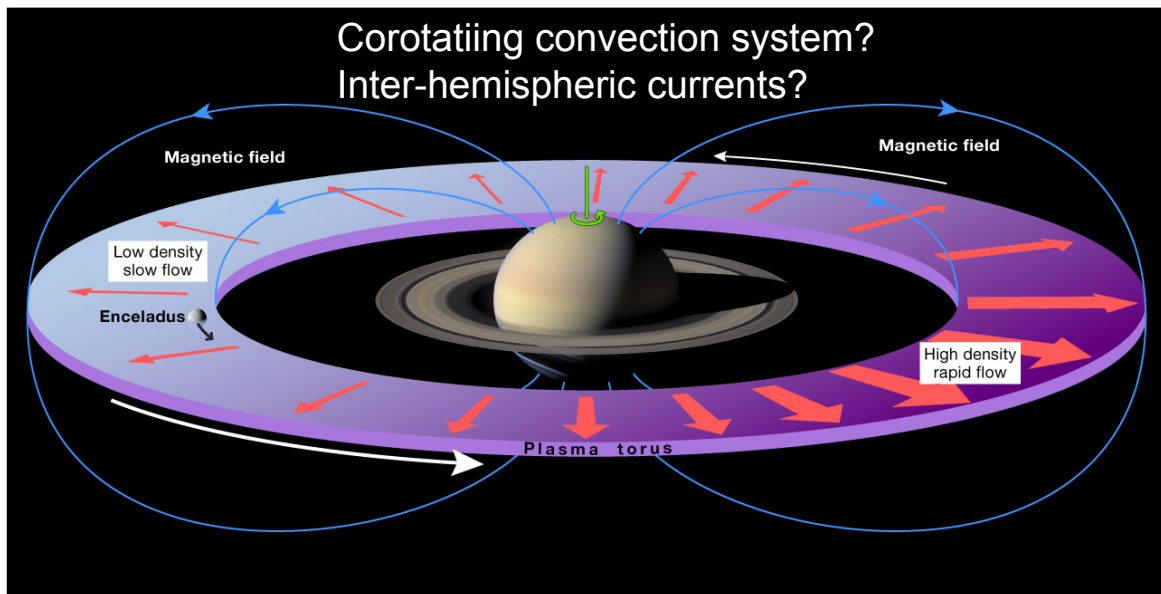
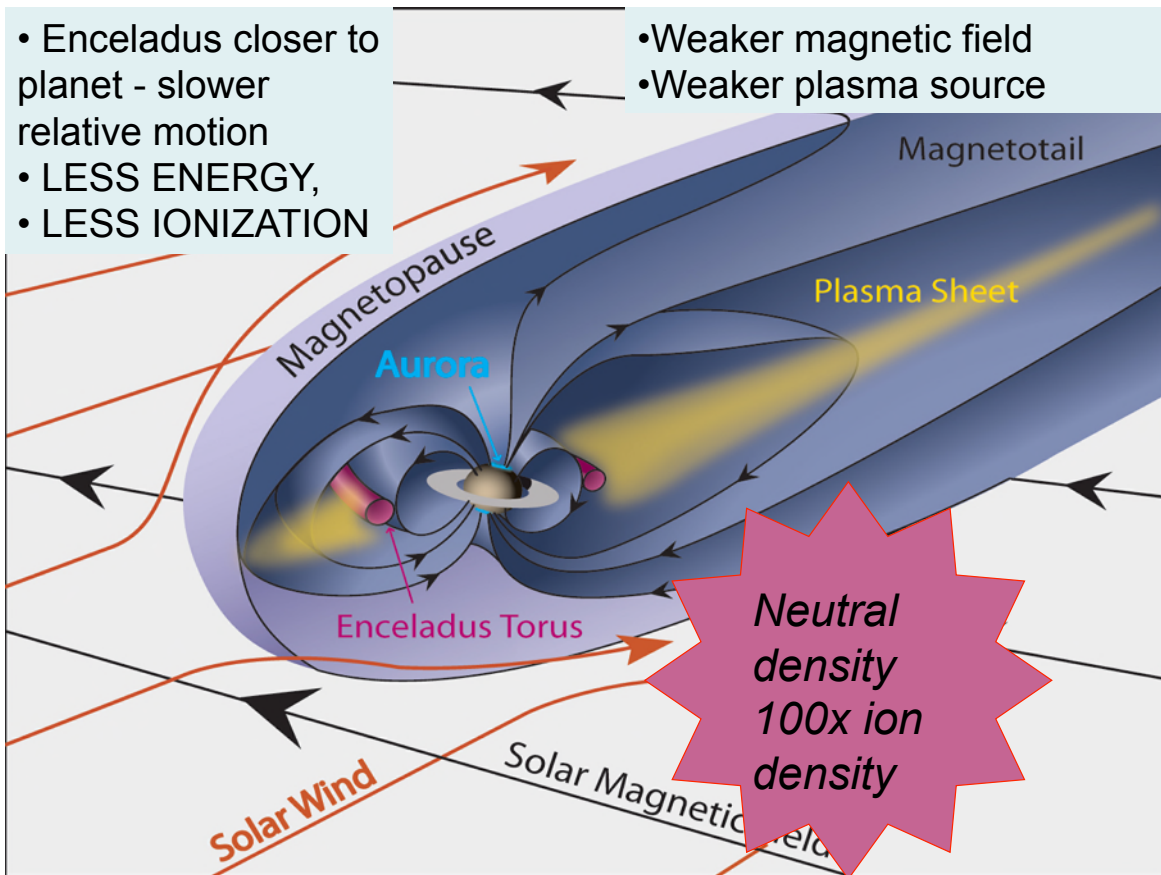
Saturn:
Strong satellite & ring sputtering, weak ionization



Enceladus!



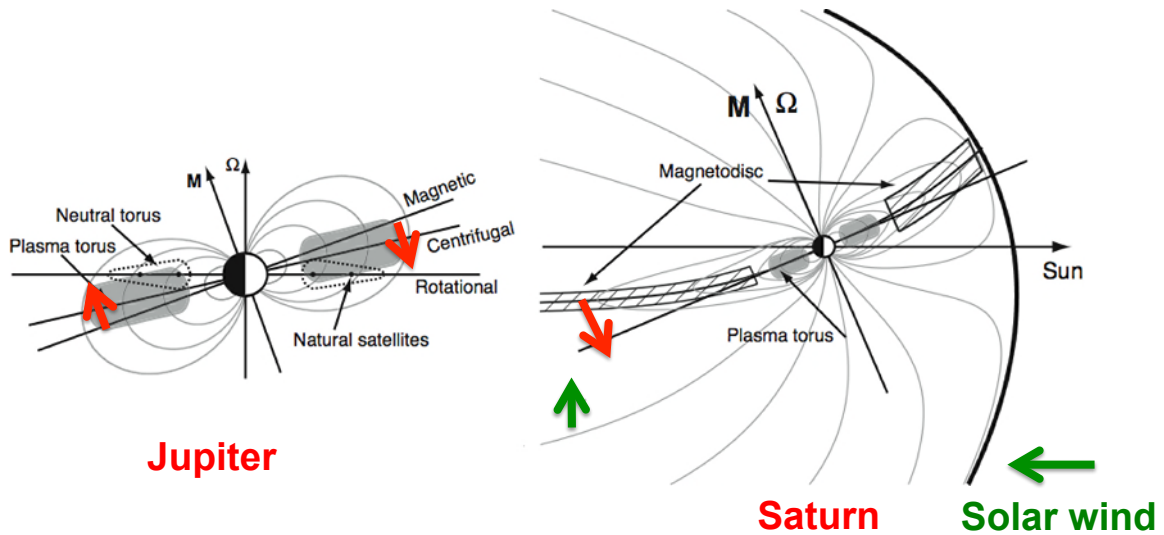
T, K
85
80
75
70
65



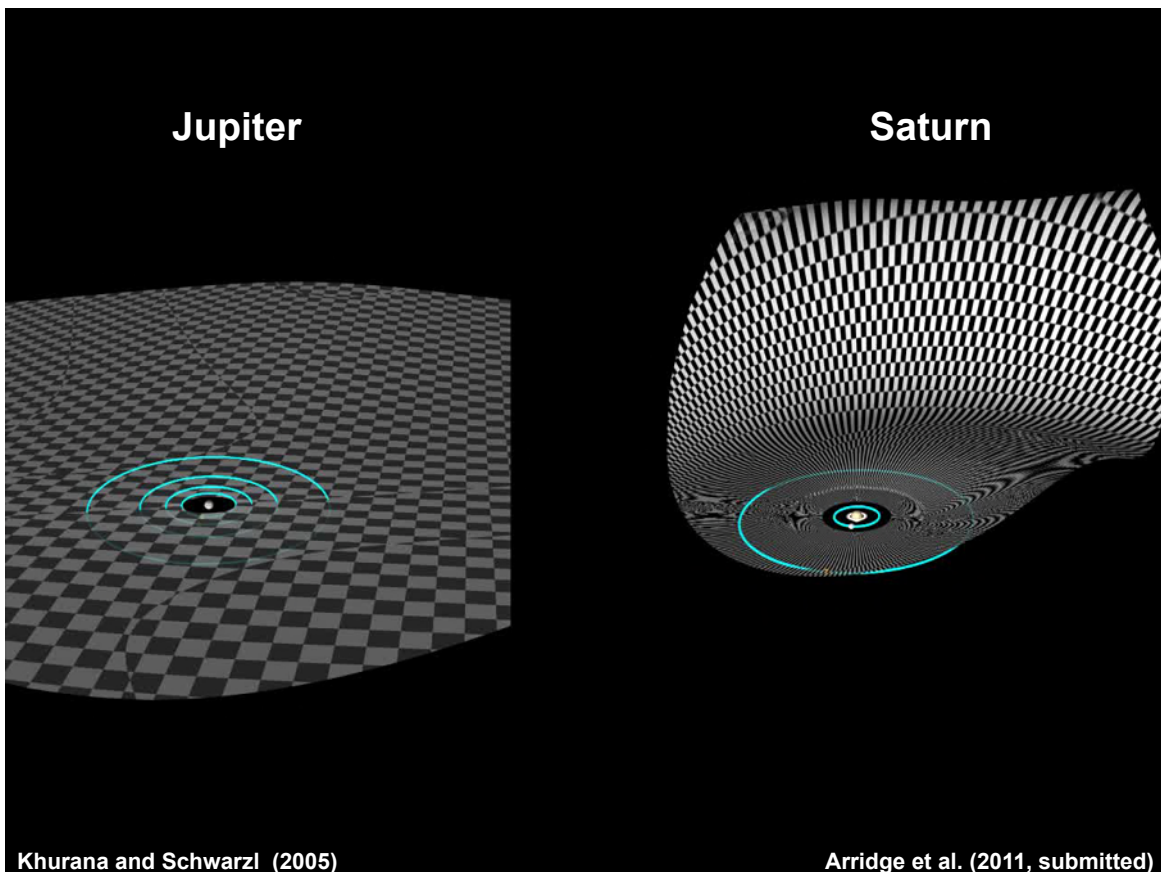
- Saturn's magnetic field is very symmetric – why are there periodic variations?
- Variable rotation rate?? Changes over “season”
- Current system between poles modulating ionization? Slowly changing with illumination of ionosphere / thermospheric winds??
- North & South ionospheres rotate at different rates??

Plasma sheet shape: forcing

- Global shape of the current and plasma sheet is determined by:
 - Diurnal motion (dipole tilt) / other periodic mechanisms.
 - Centrifugal forcing on plasma offset from the rotational equator.
 - Stresses imposed on the magnetosphere from the solar wind.



Arridge et al. (2011)



Khurana and Schwarzl (2005)

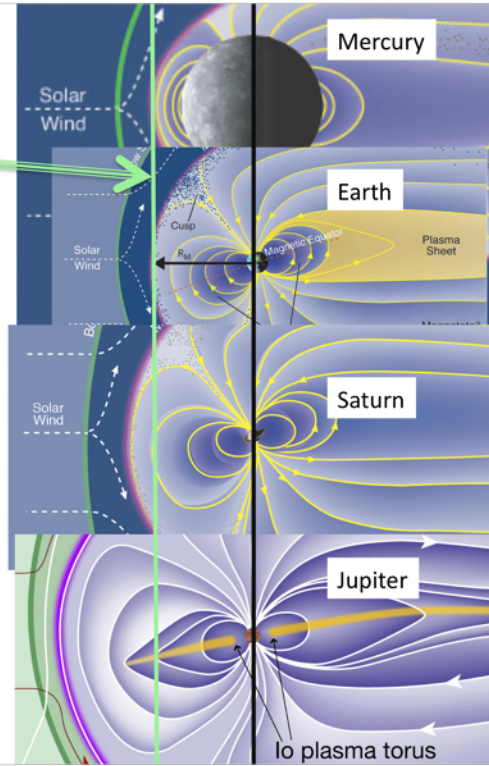
Arridge et al. (2011, submitted)

Magnetospheres scaled by stand-off distance of dipole field

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Jupiter	20,000	46 R _J	75 R _J	63-92# R _J

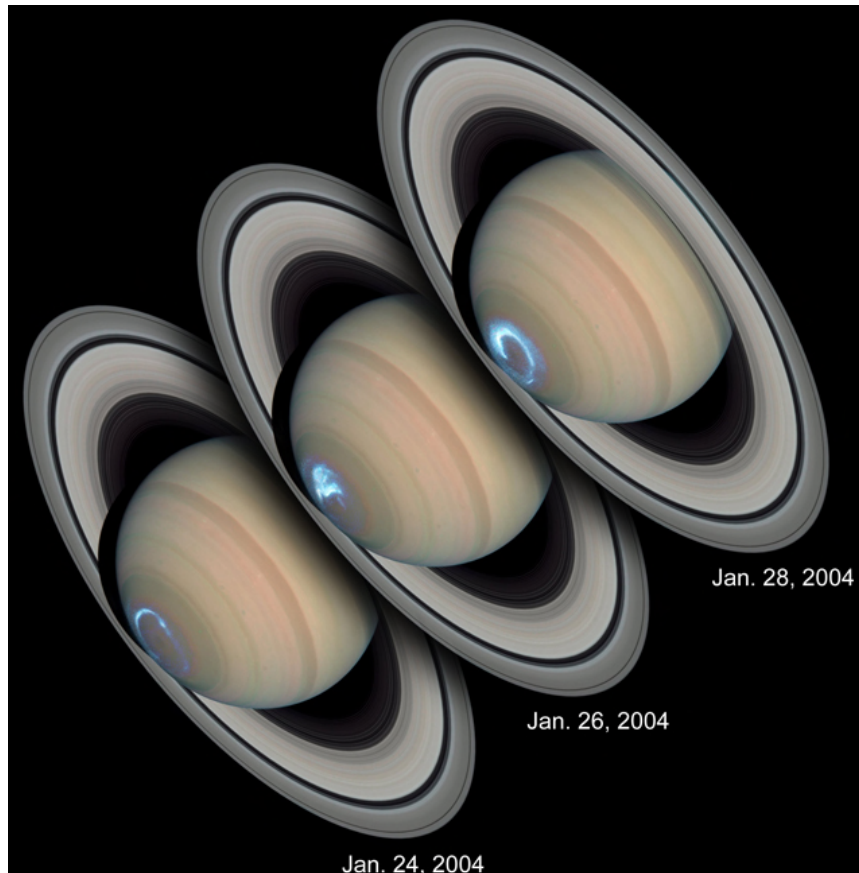
Inflated magnetospheres of Jupiter & Saturn due to HOT PLASMAS

Note bimodal average locations
 * Achilles et al. 2008 # Joy et al. 2002



Saturn's aurora

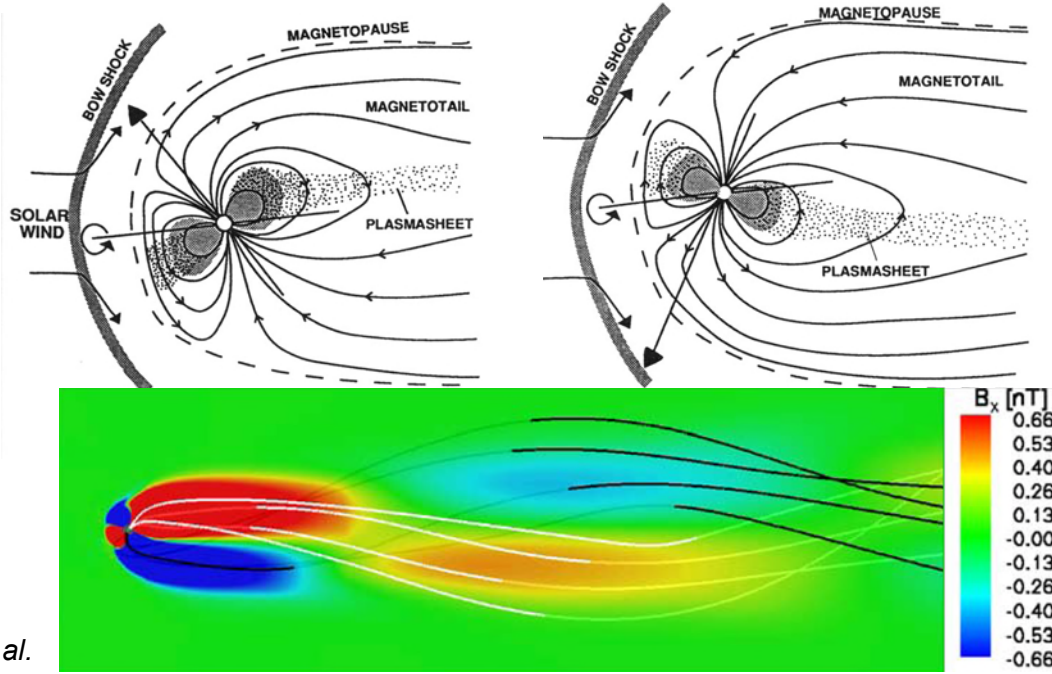
- strongly modulated by the solar wind
- open-closed boundary



Clarke et al.

Uranus

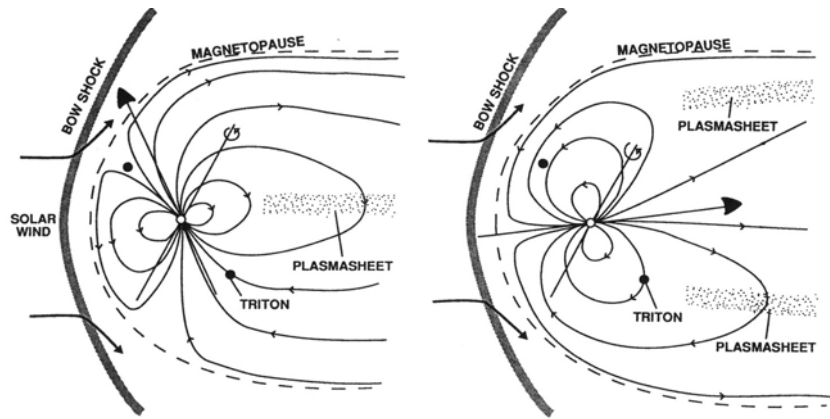
- Highly asymmetric,
- Highly non-dipolar
- Complex transport (SW + rotation)
- Multiple plasma sources (ionosphere + solar wind + satellites)



Toth et al.

Neptune

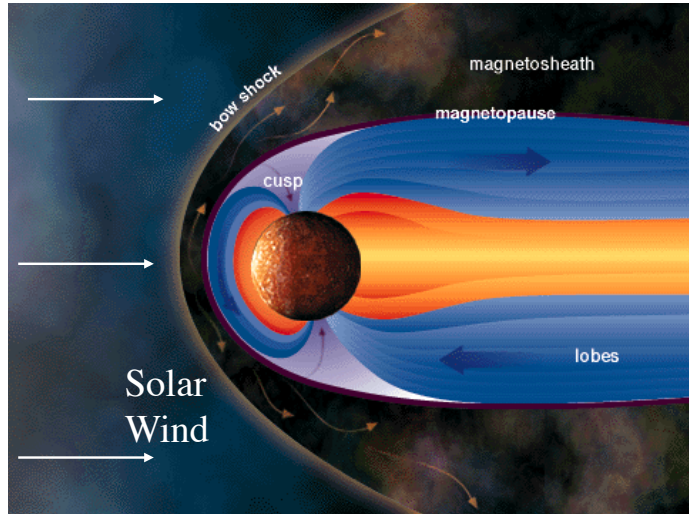
Similarly complex as Uranus



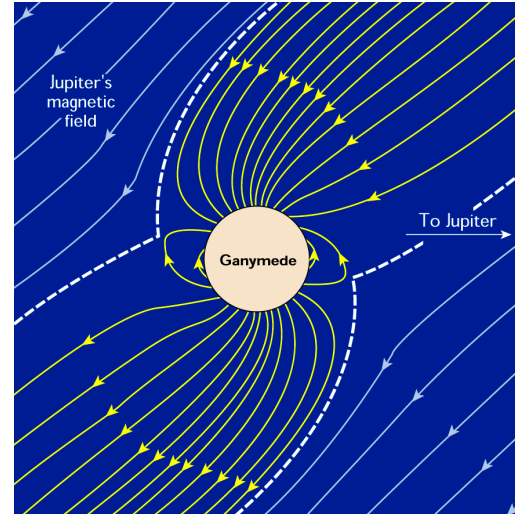
Zieger et al.

Mercury & Ganymede

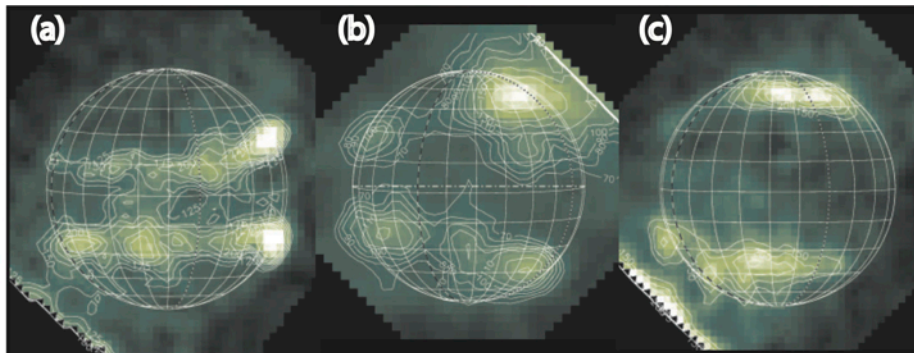
Mercury - Magnetic field detected by *Mariner 10* in 1974



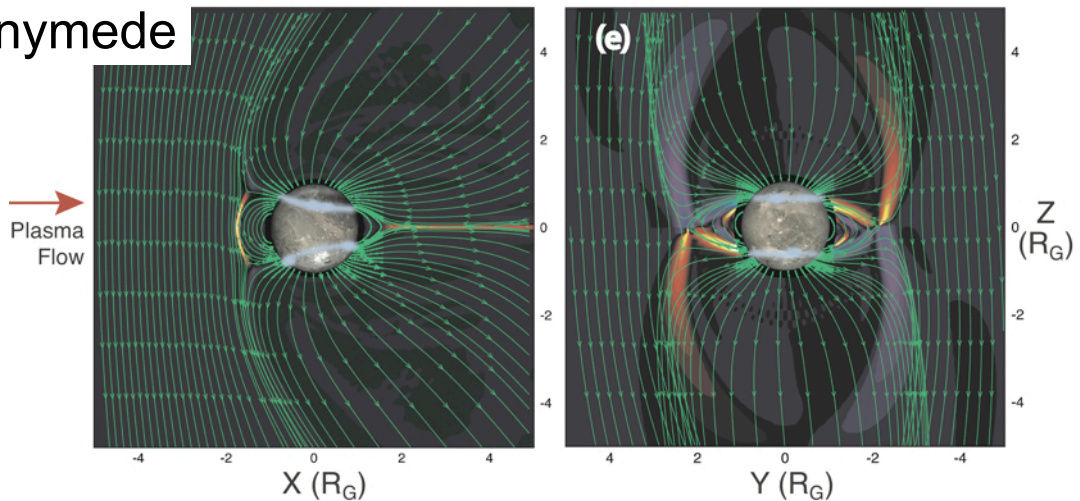
Ganymede - Magnetic field detected by *Galileo* in 1996

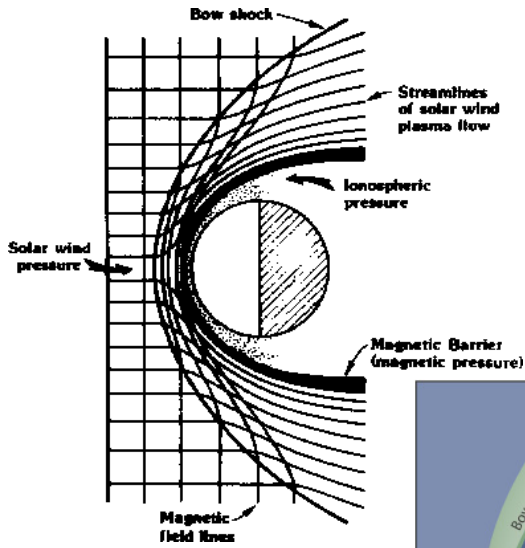


$B_{\text{surface}} \sim 1/100 \text{ Earth}$ Diameter of Earth



Ganymede





Pressure Balance:

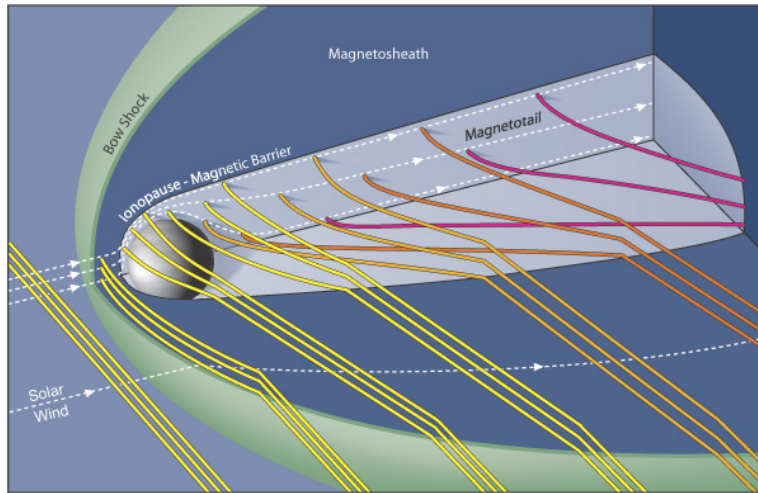
Solar wind ram pressure

-> compressed IMF

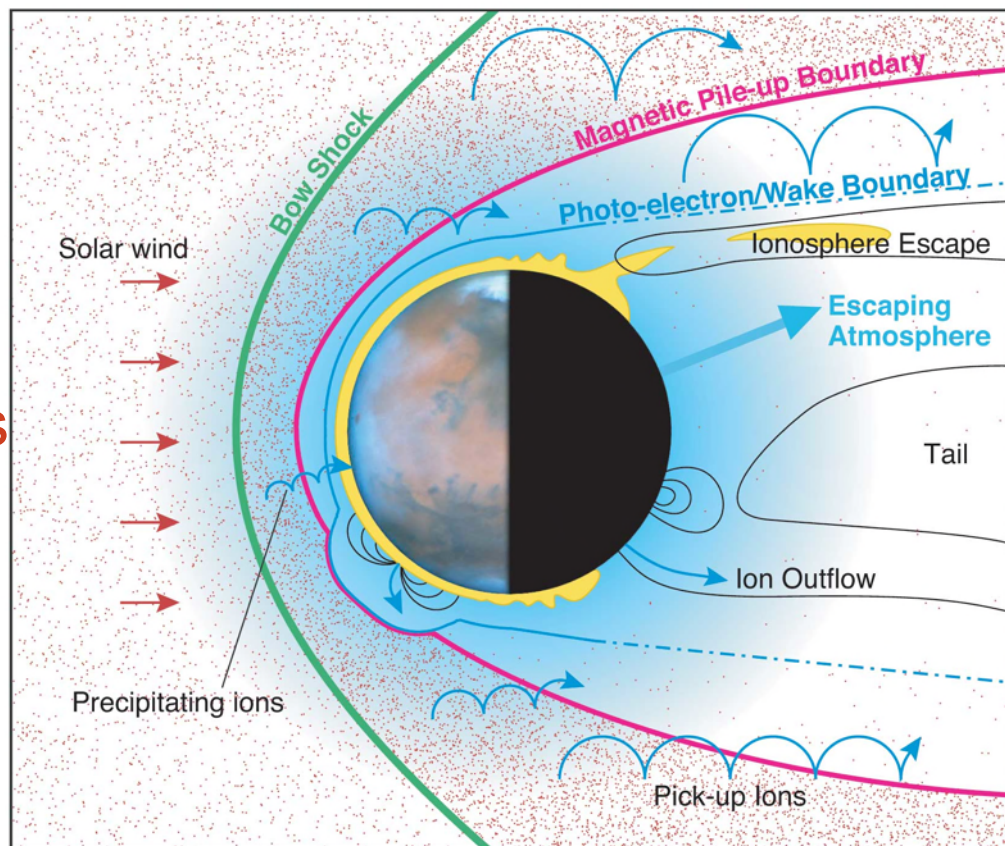
-> magnetic pressure

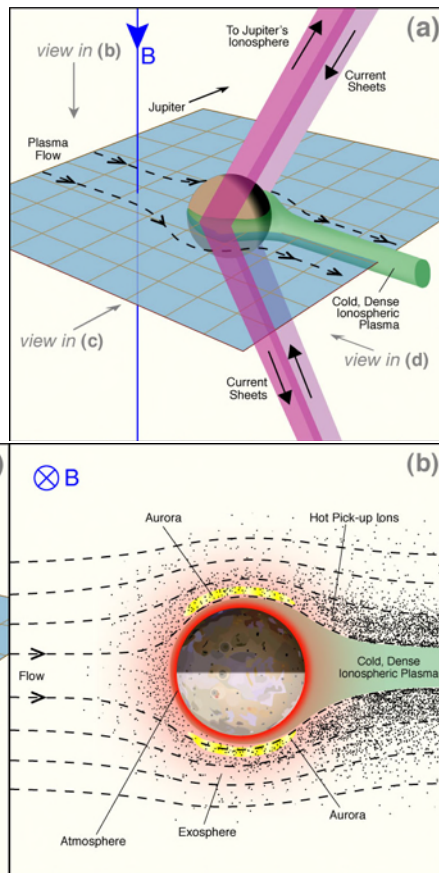
vs. ionospheric thermal pressure

VENUS



Mars

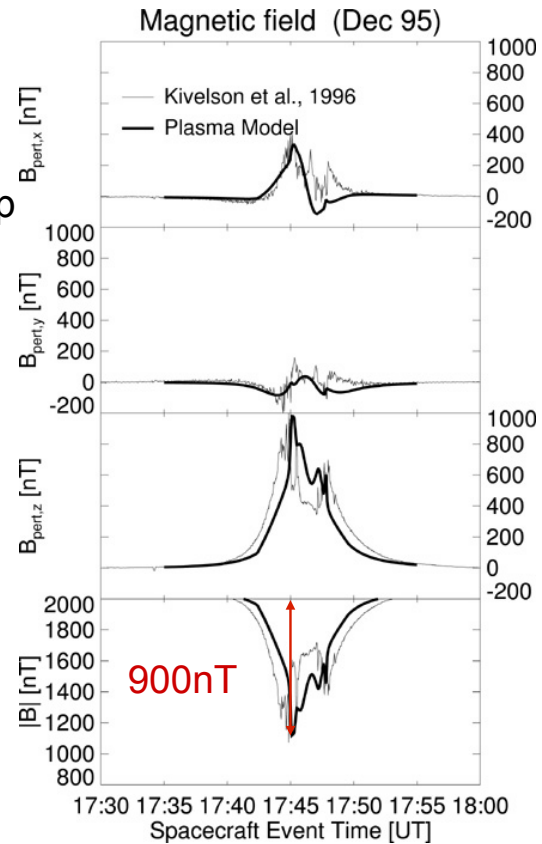




Current
= 3 MAmp

I_0

$\Delta B/B$
= 0.45



Summary

- Diverse planetary magnetic fields & magnetospheres
- Dynamo primarily requires region of liquid conducting material that is convecting – generally limited by heat flow in core
- Earth, Mercury, Ganymede magnetospheres driven by reconnection
- Jupiter & Saturn driven by rotation & internal sources of plasma
- Uranus & Neptune are complex – *need to be explored!*