

PLANETARY DYNAMOS: PS SOLUTIONS

$$\text{I (a)} \quad 2\rho \frac{\Omega}{r} \times \underline{u} \approx \underline{J} \times \underline{B}$$

$$\Rightarrow 2\rho\Omega u \sim \frac{1}{\mu_0} B^2/L \quad (\text{used Ampere's Law})$$

$$\Rightarrow B \sim \left(2\rho\Omega\mu_0 uL \right)^{1/2}$$

$$\text{(b) using } Re_M = \sigma\mu_0 uL \Rightarrow$$

$$B \sim \left(2\rho\Omega Re_M / \sigma \right)^{1/2}$$

$$\text{(c) } Re_M = 1000$$

$$\sigma = 10^6 \text{ S/m}$$

$$\rho = 10^4 \text{ Kg/m}^3$$

$$\text{Earth: } \Omega = 7.3 \times 10^{-5} \text{ s}^{-1}, \text{ Mercury: } 1.2 \times 10^{-6} \text{ s}^{-1}$$

$$B_{\text{Earth}} \sim \left(2 \cdot 10^4 \cdot 7.3 \cdot 10^{-5} \cdot 10^3 / 10^6 \right)^{1/2}$$

$$\sim \left(1.5 \times 10^{-3} \right)^{1/2}$$

$$\sim 4 \times 10^{-2} \text{ T}$$

$$B_{\text{Mercury}} \sim \left(2 \cdot 10^4 \cdot 1.2 \times 10^{-6} \cdot 10^3 / 10^6 \right)^{1/2}$$

$$\sim \left(2.4 \times 10^{-5} \right)^{1/2}$$

$$\sim 5 \times 10^{-3} \text{ T}$$

$$\text{(d) For Earth } \Gamma_{\text{CMB}} \sim 0.5 \Gamma_{\text{SWF}}$$

$$\Rightarrow B_{\text{SWF}} \sim \left(\frac{\Gamma_{\text{CMB}}}{\Gamma_{\text{SWF}}} \right)^3 B_{\text{CMB}}$$

$$\sim (0.5)^3 \cdot (4 \times 10^{-2})$$

$$\sim 5 \times 10^{-3} \text{ T}$$

$$= 5 \times 10^6 \text{ nT}$$

$$\text{Actual } B_{\text{SWF}} \sim 6 \times 10^4 \text{ nT}$$

For Mercury: $r_{CMB} \sim 0.75 r_{swf}$

$$\Rightarrow B_{swf} \sim (0.75)^3 B_{CMB}$$

$$\sim (0.4) (5 \times 10^{-3})$$

$$\sim 2 \times 10^{-3} T$$

$$= 2 \times 10^6 \text{ nT}$$

actual $B_{swf} \sim 2 \times 10^2 \text{ nT}$

Discussion

Earth: ~ 2 orders of magnitude off
Mercury: ~ 4 orders of magnitude off.

BUT: We didn't take into account the fact that not all of the B in the core reaches the surface.

B in the core is made up of 2 parts:

$$\vec{B} = \vec{B}_T + \vec{B}_P$$

toroidal B poloidal B .

$$= \nabla \times T \hat{r} + \nabla \times \nabla \times P \hat{r}$$

- only B_p leaves the core.

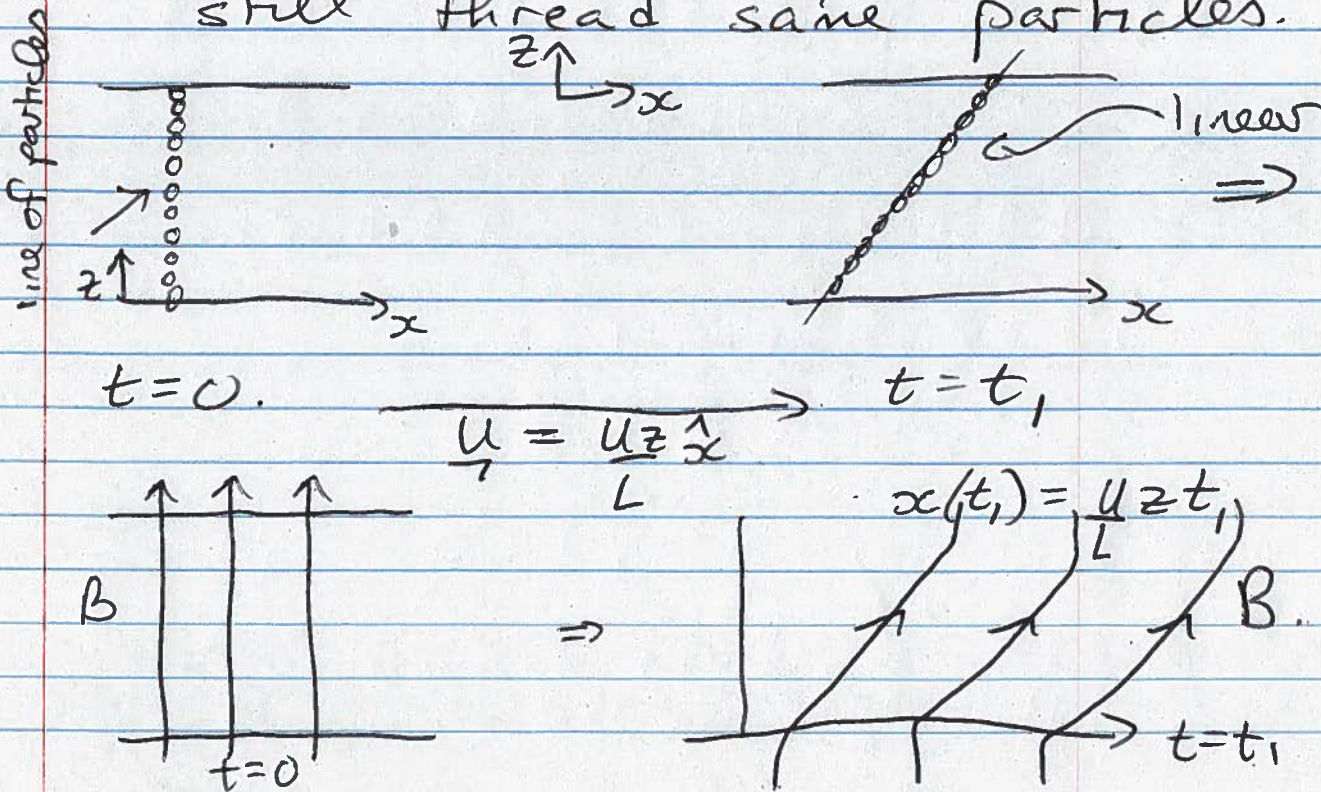
- $B_p \sim 0.1 B_T$ (this might gain an order of magnitude in our estimates. Maybe OK for Earth, Mercury still ~ 3 o.o.m. off.

[2] (a) $\frac{\partial T}{\partial t} = \alpha \nabla^2 T$
 $\frac{T}{\tau} \sim \alpha \frac{T}{L^2}$
 $\Rightarrow \tau \sim \frac{L^2}{\alpha}$

(b) $\frac{\tau_1}{\tau_2} = \left(\frac{L_1}{L_2}\right)^2 = \left(\frac{10^2}{10^3}\right)^2 = 10^{-2}$

\Rightarrow 100 km body cools $\sim 100\times$ faster.

[3] (a) key point: If fluid is perfectly conducting, then magnetic fields frozen into fluid. So take a vertical profile of fluid parcels @ $t=0$ & see where they are @ a later time. B field must still thread same particles.



$$\begin{aligned}
 (b) \quad \frac{\partial \vec{B}}{\partial t} &= \nabla \times (\vec{u} \times \vec{B}) \\
 &= \nabla \times \left(\frac{u}{L} z \hat{x} \times B_0 \hat{z} \right) \\
 &= \nabla \times \left(-\frac{u}{L} z B_0 \hat{y} \right) \\
 &= \frac{u B_0}{L} \hat{x}
 \end{aligned}$$

$$(c) \quad \frac{\partial B_i}{\partial t} = \frac{u B_0}{L} \hat{x}$$

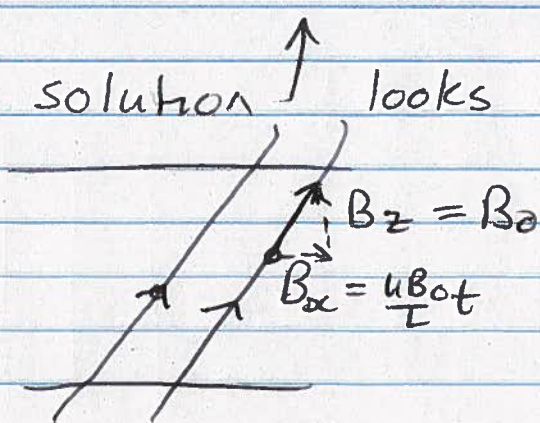
$$\Rightarrow \vec{B}_i(t) = \frac{u B_0}{L} t \hat{x}$$

$$\Rightarrow \vec{B}_{\text{tot}}(t) = \frac{u B_0}{L} t \hat{x} + B_0 \hat{z}$$

original field.

$$\begin{aligned}
 (d) \quad \nabla \times \left(+\frac{u}{L} z \hat{x} \times \frac{u B_0 t}{L} \hat{x} \right) \\
 = 0 \quad \Rightarrow \text{no new field generation.}
 \end{aligned}$$

$$\Rightarrow \vec{B}_{\text{tot}}(t) = \frac{u B_0}{L} t \hat{x} + B_0 \hat{z}$$

(e) solution  looks like (a) sketch.

